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Spatial modelling of gully erosion using Evidential Belief Function, Logistic Regression and a new ensemble EBF–LR algorithm

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Spatial modelling of gully erosion using Evidential Belief Function, Logistic Regression and a new ensemble EBF–LR algorithm

Short title: Spatial modelling of gully erosion

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Abstract

This study aims to assess gully erosion sensitivity and delineate gully erosion-prone areas in Toroud Watershed, Semnan Province, Iran. Two different methods, namely, logistic regression and evidential belief function, were evaluated, and a new ensemble method was proposed using the combination of both methods. We initially created a gully erosion inventory map using different resources, including early reports, Google Earth images and Global Positioning Systems (GPS)-aided field surveys. We subsequently split this information randomly and selected 70% (90) of the gullies for calibration and 30% (38) for validation. The method was constructed using a combination of morphometric and thematic predictors that include 16 conditioning parameters. We also assessed the following: i) potential multicollinearity issues using tolerance (TOL) and variance inflation factor indices and ii) covariate effects using LR coefficients and EBF class weights. Results show that land

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2
3 33 use/land cover, lithology and distance to roads dominate the method with the greatest effect
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5 34 on gully occurrences. We produced three sensitivity maps and evaluated their predictive
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7 35 power through area under the curve (AUC) and seed cell area index (SCAI) analyses. AUC
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9 36 results revealed that the ensemble method presented a considerably higher performance
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11 37 (AUC = 0.909) than the individual LR (0.802) and EBF (0.821) methods. Similarly, SCAI
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13 38 displayed a constant decrease from the ensemble to single methods. The resulted gully
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15 39 erosion susceptibility map could be used by decision-makers and local managers for soil
16
17 40 conservation, and for minimizing damages in development activities including construction
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19 41 of infrastructures such as roads and the route of gas and electricity transmission lines.

22 42 **Keywords:** Gully erosion; Logistic Regression; Evidential Belief Function; Ensemble
23
24 43 method; GIS; Iran

26 44 **Introduction**

28 45 Water-related soil erosion is one of the most important natural hazards in arid and semi-arid
29
30 46 regions. Apart from the land degradation, loss of soil resources, reduction of soil fertility,
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32 47 desertification and destruction of human infrastructure (on site impacts), with the deposition
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34 48 of materials in the canals and downstream slopes, impact on surface water resources and
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36 49 quality, and also impose economic and ecological costs to societies and therefore has a
37
38 50 negative impact on their sustainability development (off-site impacts) (Ayele et al., 2016).
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40 51 Global estimates show that approximately 6 million hectares of arable land lose their fertility
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42 52 annually due to soil erosion (Worker, 2004). Iran ranks second worldwide in terms of the
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44 53 volume of soil erosion (Najafi, 2005). The annual amount of soil losses in Iran is 2 to 2.5
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46 54 billion tons, which is equivalent to 8% of the global soil erosion (Najafi, 2005). This amount
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48 55 is significant, considering that Iran's share of land area is 1.1% of the world's land area.
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50 56 Conditions are extremely alarming that in the draft law on soil conservation and water
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52 57 management, more than half the area of Iran (88 million hectares) is declared in a critical

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3 58 state of erosion per hectare (Najafi, 2005). Gully erosion is the main destructive type of
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5 59 water-induced soil erosion worldwide (Pourghasemi et al., 2017; Rahmati et al., 2017).
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7 60 According to various definitions, a gully is a deep channel with steep slope edges caused by
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9 61 soil erosion and has a cross section that is greater than 0.304 m (FAO, 1965; Wang et al.,
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11 62 2016). The gully erosion, as a degradation process, with the destruction of surface and subsoil
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13 63 horizons leads to the massive deformation of the land surface and often causes extreme land
14
15 64 degradation. Development of gullies is considered as a catalytic agent of destruction of
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17 65 agricultural land, residential areas, and infrastructures. Therefore, this phenomenon impose
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19 66 extensive damages to humankind (Mekonnen et al., 2017; Pulley et al., 2018).
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21
22 67 These results have driven numerous investigations over the last decades, with the aim to
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24 68 develop methods for evaluating gully erosion processes and sensitivity mapping (Shellberg et
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26 69 al., 2016; Rahmati et al., 2017). The latter is a prerequisite for decision makers and land use
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28 70 planners. Geographical Information Systems (GIS) is considered a basic analysis tool for
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30 71 gully erosion sensitivity mapping (GESM) because it effectively manipulates spatial data
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32 72 (McCloskey et al., 2016). Moreover, recent advancement in remote sensing and modelling
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34 73 techniques contribute to significant improvements in GESM (Shellberg et al., 2016). Gully
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36 74 erosion is a compound of different physiographic factors (Dube et al., 2014) that interplay at
37
38 75 small and large scales. Therefore, the community has focused its efforts on GIS-based
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40 76 techniques for GESM, including weights-of-evidence (WoE; Rahmati et al., 2016), frequency
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42 77 ratio (FR; Conforti et al., 2011; Umer et al., 2014), logistic regression (LR; Conoscenti et al.,
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44 78 2014), information amount (Conforti et al., 2011), the analytical hierarchy process (AHP;
45
46 79 Zakerinejad and Maerker, 2014), multivariate adaptive regression splines (MARS: Gómez-
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48 80 Gutiérrez et al., 2015), artificial neural network (ANN; Rahmati et al., 2017), support vector
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50 81 machine (SVM; Pourghasemi et al., 2017), classification and regression tree (CART; Marker
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3 82 et al., 2011), random forest (RF; Kuhnert et al., 2010), and ensemble of ANN- SVM and
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5 83 maximum entropy (ME), data mining methods (Pourghasemi et al., 2017) .
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7 84 Generally, statistical methods can obtain accurate results because they can use different
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9 85 datasets (Yilmaz, 2009). Moreover, they commonly involve a catchment to regional scales
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11 86 for gully erosion studies. Such methods derive functional relations between gully erosion
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13 87 occurrences and a set of conditioning factors. In the present study, two statistical methods,
14
15 88 namely, evidential belief function (EBF) and LR methods, were applied. Literature review
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17 89 showed that LR and EBF methods have been successfully used for landslide assessment
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19 90 (Pourghasemi et al., 2013; Pham et al., 2016; Wang et al., 2016; Althuwaynee et al., 2012;
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21 91 Bui et al., 2017; Chen et al., 2017). Additionally, machine learning approaches are commonly
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23 92 used in the literature (Oh and Pradhan, 2011; Bui et al., 2012; Althuwaynee et al., 2014;
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25 93 Tehrany et al., 2015).
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28 94 In compared to other statistical approaches, LR methods allow selecting the variables step by
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30 95 step with the least amount of contingent error rates for the spatiotemporal prediction of the
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32 96 future events outside the training area (Brenning, 2005). LR can include independent
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34 97 variables of categorical or continuous nature (Brenning, 2005). It characterises the type of
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36 98 relationship between gully erosion occurrence and conditioning factors, which can be positive
37
38 99 or negative (Lee, 2005). Meanwhile, EBF methods compute the correlation between gully
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40 100 erosion and conditioning factors.
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43 101 On the basis of the literature review, despite the high capability of the EBF and LR methods
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45 102 in hazard mapping, no comprehensive study on the combination of these methods for GESM
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47 103 exists. The main purpose of this research is to apply an ensemble bivariate (EBF) and
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49 104 multivariate (LR) models as a new approach in Toroud Watershed to identify the risk areas of
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51 105 gully erosion. The results would contribute to the sustainable development in this area and
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53 106 minimise soil and economic losses.
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107 **Material and methods**

108 **Study area**

109 The Toroud Watershed is covering about 228.85 km², lies between the latitudes of
110 35°18'15"N and 35°33'33"N, and the longitudes of 54°56'33"N and 55°09'02"N in
111 northeastern of Semnan Province, Iran. (Fig. 1). The geomorphology of the study area is
112 controlled by geological, hydrological and structural setting. The northern, western, and
113 eastern sectors of the watershed are located on the cones. Slopes in this region have convex
114 profiles and are often highly dissected by V-shaped valleys. On the other hand, the central,
115 southern, southeastern and southwestern sectors are located on the gentle slopes which are
116 characterized by flat and concave profiles. The average value of elevation and slope in
117 Toroud watershed are 799.5 m and 1.52° respectively. According to the climatic
118 classification in Iran (IDWRM, 2013), the study area has an arid climate (average annual
119 rainfall of 43.8 mm), and rainfall in this area varies from 33.17 mm to 93.83 mm constituting
120 80% of the rainfall which mainly falls in December and January (IRIMO, 2014). The average
121 annual temperature of the study area is 23.4, varying from 50 degrees in summer to -5
122 degrees in winter. The study area is mainly covered with deserts and poor pasture. Its
123 lithology is mainly composed of Qft2 (low-level piedmont fan and valley terrace deposits),
124 Qft1 (high-level piedmont fan and valley terrace deposits), PIQc (fluvial conglomerate,
125 piedmont conglomerate and sandstone), Qsf (salt flat) and Murm (light - red to brown marl
126 and gypsiferous marl with sandstone intercalations) (GSI, 1997).

127 The study area is characterized by poorly evolved soils (Entisols and Aridisols; USDA, 2006).
128 Because of the high concentration of gullies is near roads and agricultural lands, therefore the
129 development of gullies causes massive economic damage to the inhabitants. The study area is
130 selected because of two main reasons: 1) a highway is crossing across this region connecting
131 between the two major provinces of Iran (Isfahan and Khorasan Razavi); 2) the area is highly

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3 132 susceptible to gully erosion due to complex geological formations with the presence of
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5 133 gypsum and salt leading to high evaporation.
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9 135 **Fig. 1** Study area.
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13 137 **Methodology**

15 138 As shown in Fig. 2, this study is conducted in four stages as follows: (1) preparation of data
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17 139 (Meyer and Martinez-Casasnovas, 1999), (2) application of EBF method and determination
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19 140 of the relationship between gully erosion occurrences and conditioning factors (Al Abadi and
20
21 141 Al Ali, 2017), (3) utilisation of the LR method and determination of the effect of conditioning
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23 142 factors (Martinez-Casasnovas et al. 2004), (4) GESM of the study area using individual and
24
25 143 combined methods and (5) evaluation of the capability and robustness of the combined
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27 144 method in comparison with the individual EBF and LR methods in terms of receiver
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29 145 operating characteristic (ROC)–area under the curve (AUC) and seed cell area index (SCAI)
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31 146 (Luca et al., 2011).
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37 148 **Fig. 2** Flowchart of the methodology used.
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41 150 **Preparation of dependent and independent variables**

43 151 **Gully erosion inventory map (dependent variable)**

45 152 The geomorphological community has widely used the hypothesis that events in the past are
46
47 153 important for understanding the future (Cama et al., 2017) by using statistical methods and
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49 154 predict various phenomena. This assumption postulates that unless significant environmental
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51 155 changes occur, we can investigate previous events, derive functional relations with respect to
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53 156 a given set of covariates and predict potential future occurrences. The gully erosion literature
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3 157 has also used this hypothesis to study the sensitivity of a given area (Pourghasemi et al.,
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5 158 2017). However, for GESM, the primary requirement includes the data collection or
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7 159 inventory that represents the dependent variable of any predictive method (Pourghasemi et
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9 160 al., 2017).

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13 162 In the present work, early reports, Google Earth images (dated 22/05/2017) and
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15 163 comprehensive field surveys were used to determine the locations of gully erosion and
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17 164 construct an accurate gully erosion inventory map. A total of 128 gullies in the study area
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19 165 were digitised (Fig. 1). Out of the 128 gully erosion occurrences, 90 cases (70%) were
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21 166 randomly selected to calibrate our methods, and the remaining 38 gullies were used for
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23 167 validation (Cama et al., 2017). Some samples of gullies are shown in Fig. 3.

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26 168 **Fig. 3** Field photographs showing identified gullies in the study area.

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29 30 170 **Gully erosion conditioning factors (independent variables)**

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32 171 The selection of suitable gully erosion conditioning factors (independent variables) is a main
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34 172 step in the modelling of phenomena, such as gully erosion (Rahmati et al., 2017).

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36 173 Sixteen geo-environmental parameters were selected on the basis of the relevant literature
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38 174 (Rahmati et al., 2017) and the opinions of academics and natural resources experts. Out of
39
40 175 these 16 factors, the geomorphometric ones were derived from a DEM with horizontal
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42 176 resolution of 30 m using the ASTER GDEM elevation data together with ArcGIS 10.5 and
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44 177 SAGA-GIS software.

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47 178 For clarity, we report the manner in which we computed the less common predictors adopted
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49 179 in this study (i.e., stream length (LS), topography wetness index (TWI) and stream power
50
51 180 index (SPI)) as follows (Moore et al., 1991):

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$$LS = (A_S/22.13)^{0.6} \times (\sin\beta/0.0896)^{1.3}, \quad (1)$$

$$TWI = \ln(A_S/\tan\beta), \quad (2)$$

$$SPI = A_S \times \tan\beta. \quad (3)$$

Where A_S is the special area of the basin (m^2/m) and β is slope in degrees.

Landsat 8 OLI images were selected to extract the Normalized Difference Vegetation Index (NDVI) and land use/land cover (LU/LC) (Fig 4a) using ENVI 5.1 software. The NDVI amount was extracted in ArcGIS 10.5 software.

The lithological unit map (Fig 4b) was obtained from the geological map in 1:100,000-scale (GSI, 1997) and were classified into eight classes on the basis of their lithological characteristics and prior knowledge of their influence on gully erosion occurrence. Table 1 shows the lithological characteristics of Toroud Watershed. A soil type map (Fig 4c) was obtained from the Semnan Agricultural and Natural Resources Research Centre and classified into three classes.

Distance to roads (Fig 4d) was calculated from the topographic map with a scale of 1:50,000. The mean annual rainfall layer was prepared in ArcGIS10.5 environment based on 30 years (1984–2014) of precipitation data of the Toroud, Razveh, Moalleman and Hosseinan Stations from the Iran Meteorological Organisation (IRIMO, 2014).

The classes of each gully erosion conditioning factor are shown in Table 2. The pixel size of all layers is the same with that of the digital elevation method (Dube et al., 2014).

Fig. 4 Some examples of gully erosion conditioning factors, a) LU/LC, b) lithology, c) soil type, and d) distance to road.

Table 1 Description of geological units in the study area.

Table 2 Overview of factors used for GESM.

206 Multicollinearity analysis of gully erosion conditioning factors

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3 207 When applying any linear statistical method, the multicollinearity of data may hinder the
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5 208 reliability and interpretation of results (Pradhan and Seenii, 2017, 2018; Pradhan et al., 2017).
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7 209 Multicollinearity is defined as the linear dependency that links two or more independent
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9 210 variables in a dataset. Thus, multicollinearity must be checked even for gully erosion
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11 211 sensitivity methods. Several approaches may be adopted to assess multicollinearity. For
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13 212 example, the least absolute shrinkage and selection operator (Camilo et al., 2017) is popularly
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15 213 used to penalise the number of covariates in each method, thereby removing the redundant
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17 214 information carried by collinear predictors. Alternatively, an example can be found in the
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19 215 literature where the variance inflation factor (VIF) is used to recognise the presence of
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21 216 multicollinearity in data (Cama et al., 2017). Here, we opted for VIF, which is the reciprocal
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23 217 of TOL, and calculated as $1 - R^2$, where R^2 is obtained by regressing each variable with
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25 218 respect to the remaining variables in the multivariate regression (Holloway et al., 2017). A
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27 219 $TOL \leq 0.1$ or a $VIF \geq 10$ indicates serious multicollinearity (Guo-Liang et al., 2017).
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33 221 **LR method**

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37 223 LR is a multivariate statistical method that corresponds to the generalised linear method
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39 224 (GLM) when the distribution to be fitted corresponds to Bernoulli distribution (Hemasinghe
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41 225 et al., 2018). This distribution describes the probabilities of a binary outcome; thus, it is
42
43 226 extremely convenient for predicting the presence or absence of gully erosion. LR can derive
44
45 227 multivariate relationships between gullies and a set of predictors, assuming linearity between
46
47 228 the target and explanatory variables, the latter being continuous and categorical in nature
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49 229 (Lee, 2005). These relationships can be used in an additive equation to produce the
50
51 230 probability of gully occurrence, thereby generating GESMs (Zhou et al., 2018). GLMs,
52
53 231 particularly LR, are quantitative methods that determine the influence of each independent
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variable through the coefficients and anti-logarithm of the coefficients (Lee and Sambath, 2006), which offers geomorphologists a chance to infer which environmental properties may dominate the rise of gully erosion and perform remediation to mitigate such phenomenon.

The LR is expressed as follows:

$$P = \frac{1}{1+e^{-z}} \quad (5)$$

where P is the possibility of gully erosion occurrence, denoted by 0 to 1; and Z denotes the gully erosion conditioning parameters and presumed as a linear composition of the conditioning factors, X_i ($i = 1, 2, \dots, n$), as shown as follows:

$$Z = \text{Logit}(P) = \ln \frac{p}{1-p} = C_0 + C_1 X_1 + \dots + C_n X_n, \quad (6)$$

where C_0 is the constant coefficient of the method; and X_1, X_2, \dots, X_n are the coefficients of independent variables C_1, C_2, C_n , respectively.

EBF

EBF methods have been successfully used as a data-based approach for potential assessment of mineral deposits and landslide and groundwater sensitivity mapping (Mogaji et al. 2014).

Data-driven EBF methods can compute the weight of each class of conditioning factors using the relationship between classes and phenomenon occurrence (Wang et al., 2016). This method has four functions, including belief (Bel), disbelief (Dis), uncertainty (Unc) and plausibility (Pls; Mogaji et al. 2014). These functions are calculated as follows (Park et al., 2011):

$$(Bel) = \left[\frac{[\lambda(T_p)_{A_{ij}}]}{\sum[\lambda(T_p)_{A_{ij}}]} \right], \quad (7)$$

$$[\lambda(T_p)_{A_{ij}}] = \left[\frac{N(F \cap A_{ij})}{N(F)} \right] / \left[(N(A_{ij}) - N(F \cap A_{ij})) / [N(P) - N(F)] \right] \quad (8)$$

$$(Dis) = \frac{[\lambda(\bar{T}_p)_{A_{ij}}]}{\sum[\lambda(\bar{T}_p)_{A_{ij}}]}, \quad (9)$$

$$[\lambda(\bar{T}_p)_{A_{ij}}] = \left[\frac{N(F \cap A_{ij})}{N(F)} \right] / \left[(N(P) - N(F) - N(A_{ij}) + N(F \cap A_{ij})) / [N(P) - N(F)] \right], \quad (10)$$

$$(Unc) = [1 - (Bel) - (Dis)], \quad (11)$$

$$(Pls) = [1 - (Dis)], \quad (12)$$

where $N(F \cap A_{ij})$ is the aggregation of gullies occurring in A_{ij} , $N(F)$ is the total aggregation of the entire gullies in the study area, $N(A_{ij})$ is the aggregation of pixels in A_{ij} and $N(P)$ is the aggregation of pixels in the entire study area.

268 Validation of GESMs

269 A suitable validation is necessary to produce a reliable GESM for any area (Rahmati et al.,
270 2016). Various validation strategies have been introduced in the literature, among which the
271 ROC curve and its integral (known as AUC) have been widely adopted in the
272 geomorphological community (Pourghasemi et al., 2017) together with Success or prediction
273 rate curve (Frattini et al., 2010), and cell area index (Süzen and Doyuran, 2004).

274 In this study, we combined AUC and SCAI to provide a comprehensive assessment of the
275 validation performance (Hong et al., 2018; Thai Pham et al., 2018; Lee et al., 2018;
276 Pourghasemi and Rahmati, 2018).

277 AUC is one of the most useful and efficient methods for predicting and determining the
278 accuracy of models (Gómez-Gutiérrez et al., 2015). In fact, this curve is considered a
279 graphical representation of the true prediction of occurrence and non-occurrence of a

phenomenon (Dube et al., 2014). AUC represents the predictive amount of the system by describing its capability to estimate the occurrence (the presence of gully) and non-occurrence (the absence of gully) of events accurately. If a method cannot estimate the occurrence of gullies better than the probable or random point, then the AUC is 0.5 and therefore the least accurate; if the AUC is equal to 1, then the method is the most accurate. AUC values can be classified as follows: 0.5–0.6, poor; >0.6–0.7, average; >0.7–0.8, good; >0.8–0.9, very good; and >0.9–1, excellent (Yesilnacar, 2005). The SCAI validation technique was proposed by Suzen and Doyuran (2004). SCAI is calculated by dividing the percentage of pixels of the specific gully erosion sensitivity class by the percentage of the pixels of existing gullies in the specific gully sensitivity zone. SCAI shows the density of gullies among the gully sensitivity zones (Pawluszek and Borkowski, 2017). In the SCAI indicator, the high and very high susceptibility classes have very small SCAI values, and low and very low susceptibility classes have higher SCAI values (Suzen and Doyuran, 2004).

293 **Results**

294 **Multicollinearity test**

295 Before applying the EBF and LR methods, we checked for potential linear dependencies
296 between the pairs of covariates to avoid multicollinearity among variables. The values of
297 TOL and VIF of the 16 conditioning factors were calculated to detect multicollinearity (Table
298 3). From Table 3, the maximum VIF and minimum TOL were 3.452 and 0.290, respectively.
299 The results indicated that no multicollinearity existed between the conditioning factors.

300 **LR method**

301 In this study, we calculated the regression coefficients of the gully erosion conditioning
302 factors using IDRISI software. The coefficients of the gully erosion conditioning factors
303 using the LR method are shown in Fig. 5. According to LR results, parameters of LU/LC,
304 lithology, distance to road and soil type with scores 0.692, 0.492, 0.385, and 0.355

305 respectively, have had the greatest effect on the gully erosion. This results are in line with the
 306 findings of (Conoscenti et al. 2014). On the contrary, factors of slope aspect, profile
 307 curvature, TWI, and drainage density with scores -0.098, -0.0541, -0.054, and -0.026
 308 respectively, have had little impact on the gully erosion. Rest of the parameters such as slope
 309 degree, plan curvature, SPI, NDVI, rainfall, distance to stream, convergence index and LS are
 310 ranked from fifth to twelfth.

311 The z value is computed as follows:

$$\begin{aligned}
 Z = & -11.7494 + (-0.098 \times \text{slope aspect}) + (0.492 \times \text{lithology}) \\
 & + (0.086 \times \text{slope degree}) + (0.082 \times \text{plan curvature}) \\
 & + (0.01 \times \text{distance to stream}) + (0.043 \times \text{NDVI}) + (0.044 \times \text{SPI}) \\
 & + (-0.054 \times \text{TWI}) + (-0.024 \times \text{LS}) + (0.355 \times \text{soil type}) \\
 & + (-0.026 \times \text{drainage density}) + (-0.013 \times \text{convergence index}) \\
 & + (0.693 \times \text{land use}) + (-0.054 \times \text{profile curvature}) \\
 & + (0.022 \times \text{rainfall}) + (0.385 \times \text{distance to road}). \quad (13)
 \end{aligned}$$

313
 314 According to the z value, the P values were computed using Eq. 5. The P values varied from
 315 0.0032 to 0.074. The subsequent sensitivity amounts were classified into five classes (Fig. 6a)
 316 according to the natural breaks method (Guo-Liang et al., 2017). The resulting map indicated
 317 that the low class had the largest area (30.85%), followed by the moderate (18.18%), very
 318 low (17.56%), high (17.26%) and very high (16.12%) classes. In addition, the largest
 319 percentage (47.15%) of the total gullies fell in the very high sensitivity zones, followed by
 320 the moderate (27.15%), high (13.11%), low (11.65%) and very low (0.917%) zones (Table
 321 5).

322 **Fig. 5** Coefficients of the conditioning factors in the LR model.

1
2
3 323 **Table 3** Multicollinearity test among conditioning factors.
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5
6 325 **EBF method**
7

8 326 The results of the EBF method are shown in Table 4. The correlation rate between the
9
10 327 conditioning factors and the gully erosion occurrence is related on the Bel amounts. The
11
12 328 absence of Bel in each class indicates no contribution to gully erosion (Pradhan et al., 2014;
13
14 329 Zeinivand and Ghorbani Nejad, 2018). As shown in Table. 4, the slope degree in the class of
15
16 330 $<1^\circ$ has the most Bel amount (0.306). This result indicates the utmost effect on gully erosion
17
18 331 occurrences. The slope aspect suggests that the SE and F classes have the largest Bel amounts
19
20 332 (0.218 and 0.204, respectively); therefore, gully erosion in these classes is high, which is
21
22 333 consistent with the findings of Rahmati et al. (2016). On the basis of the profile curvature, the
23
24 334 largest degree of belief amounts (0.227) is related to the class of -0.001 to -0.0005 .
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27

28 335 **Table. 4** Spatial relationship between gully erosion occurrence and conditioning factors using
29 336 EBF model.
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32
33 338 In the case of plane curvature, the results show that flat areas with the largest degree of Bel
34
35 339 (0.367) contribute a higher probability of gully erosion than the concave (0.338) and convex
36
37 340 (0.293) areas. On the basis of NDVI results, classes of 0.039–0.13 and >0.39 with Bel
38
39 341 amounts of 0.639 and 0.363, respectively, have a strong relationship with gully occurrence,
40
41 342 whereas the class of >0.13 with Bel amount of 0 has no influence on gully erosion. This
42
43 343 result is consistent with that of Gomez-Gutierrez et al. (2009). On the basis of the relationship
44
45 344 between gully occurrence and LU/LC among the various land uses, desert areas have the
46
47 345 largest Bel amount (0.772) and therefore has the largest gully sensitivity. By contrast,
48
49 346 irrigated lands have no correlation with gully erosion because of vegetation cover. These
50
51 347 results are similar to those of Conoscenti et al. (2014). In the case of lithology, the results
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53 348 indicated that high levels of piedmont fan and valley terrace deposits (Qft1) with maximum
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349 Bel degrees (0.276) show the upmost sensitivity to gully erosion. For the distances to road
 350 and river, an inverse relationship exists between the sensitivity of area to gully erosion and
 351 Bel degree. In other words, the amount of Bel decreases with the increase in distance of the
 352 factors. This result agrees with that of Shellberg et al. (2016), which suggests that
 353 anthropogenic linear features in the landscape, such as roads, contribute to flow concentration
 354 and thus contribute to gullies. Moreover, Conoscenti et al. (2014) statistically correlated
 355 gullies to a river network, which we investigated in the present study, including drainage
 356 density. The Bel amounts for drainage effects show that gully erosion sensitivity is high in
 357 the class of 2.19–2.54 km/km²; particularly, a low density indicates a low Bel amount. The
 358 analysis of Bel for the relationship between gully erosion occurrence and TWI shows that the
 359 class of –0.84–0.94 with the largest Bel amount (0.284) has a strong correlation with gully
 360 erosion. According to the rainfall results, the class of 80–85 mm has the largest amount of
 361 Bel (0.410). In the case of SPI, the results indicate that the class of 7,000–11,000 has a high
 362 sensitivity to gully erosion, and the class of >11,000 with the lowest Bel amount (0.055) has
 363 the weakest correlation with gully occurrence. According to the Bel degree of the LS factor, a
 364 direct relationship exists between sensitivity to gully occurrence and Bel degree. Therefore,
 365 the probability of gully erosion increases proportionally with LS. The interpretation of the
 366 role of soil type in the method indicates that the Entisols/Aridisols class with Bel amount of
 367 0.934 has a strong and positive relationship with gully occurrence. These results indicate that
 368 these soils are prone to erosion. On the basis of the convergence index, this parameter
 369 exhibits a diverse relationship with gully erosion and with increasing convergence. After the
 370 computation of the spatial relationship between gully erosion and conditioning factors, the
 371 GESM via the EBF method was constructed using Eq. (14).

372 $ESM =$

373 $(slope\ aspect_{Bel}) + (lithology_{Bel}) + (slope\ degree_{Bel}) + (plan\ curvature_{Bel}) +$

$$\begin{aligned}
 & (distance\ to\ stream_{Bel}) + (NDVI_{Bel}) + (SPI_{Bel}) + (TWI_{Bel}) + (LS_{Bel}) + \\
 & (soil\ type_{Bel}) + (drainage\ density_{Bel}) + (convergence\ index_{Bel}) + (land\ use_{Bel}) + \\
 & (profile\ curvature_{Bel}) + (rainfall_{Bel}) + (distance\ to\ road_{Bel}) \quad (14)
 \end{aligned}$$

The result of the EBF method ranges from 1.850 to 6.750. Then, the resultant GESM was divided into five classes from very low to very high (Fig. 6b; Wang et al., 2016, Arabameri et al., 2017). The results of classification indicate that 19.37% and 20.48% of the study area are in the very low and low sensitivity zones, respectively, and 19.83%, 19.97% and 20.35% fall in the moderate, high and very high sensitivity zones, respectively. Moreover, 49.81% and 27.20% of the total gullies fall in the very high and high sensitivity zones, respectively; whereas moderate, low and very low susceptible zones are 15.16%, 6.25% and 1.56% of the gullies, respectively (Table5).

Fig. 6 GESM using LR, EBF and ensemble models.

Table 5 Area of susceptibility classes and SCAI.

Ensemble method

The combination of models compounds the results of individual methods into a combined method to increase the precision of prediction capability and has therefore received considerable interest from researchers (Guo-Liang et al., 2017). In the present work, LR and EBF methods were combined for the GESM of Toroud Watershed to overcome their individual disadvantages. For this purpose, the coefficients obtained by LR were multiplied by the weight obtained in the EBF method using Eq. 15. Finally, the sensitivity of landscape to gully erosion was computed in ArcGIS10.5 using the Weighted Sum Tools.

$$\begin{aligned}
 GESM = & (-0.098 \times slope\ aspect_{Bel}) + (0.492 \times lithology_{Bel}) + (0.086 \times \\
 & slope\ degree_{Bel}) + (0.082 \times plan\ curvature_{Bel}) + (0.01 \times distance\ to\ stream_{Bel}) +
 \end{aligned}$$

$$\begin{aligned}
& (0.043 \times NDVI_{Bel}) + (0.044 \times SPI_{Bel}) + (-0.054 \times TWI_{Bel}) + (-0.024 \times LS_{Bel}) + \\
& (0.355 \times soil\ type_{Bel}) + (-0.026 \times drainage\ density_{Bel}) + \\
& (-0.013 \times convergence\ index_{Bel}) + (0.693 \times land\ use_{Bel}) + \\
& (-0.054 \times profile\ curvature_{Bel}) + (0.022 \times rainfall_{Bel}) + \\
& (0.385 \times distance\ to\ road_{Bel}) \qquad \qquad \qquad (15)
\end{aligned}$$

The resultant map was classified into five classes similar to the EBF and LR methods (Fig. 7c). The results indicate that from the total area (227.97 km²) of Toroud Watershed, 19.63% (44.76 km²) belong to the very low sensitivity class, 20.23% (46.11 km²) to the low class, 20.14% (45.92 km²) to the moderate class, 19.70% (44.91 km²) to the high class and 20.28% (46.24 km²) to the very high sensitivity class. Generally, from the total gully erosion area (2.72 km²), 0.82% (0.022 km²) fall in the very low sensitivity class, followed by 3.67% (0.1 km²), 10.93% (0.293 km²), 20.68% (0.562 km²) and 63.87% (1.737 km²) that fall in the low, moderate, high and very high sensitivity classes, respectively.

414 **Validation of the three methods**

415 The results of validation using the AUC and SCAI methods are shown in Tables 5 and 6.
 416 According to the AUC results, the combined method has a higher accuracy of 0.909 (90.09%)
 417 than the individual statistical methods (EBF 0.821 (82.1%) and LR 0.802 (80.2%)).

418 In addition, the amount of SCAI in the three methods gradually decrease from very low to
 419 very high sensitivity zones; consequently, the values of SCAI decreased from 12.48, 10.53
 420 and 17.39 in the very low sensitivity class to 0.316, 0.430 and 0.348 in the very high
 421 sensitivity class in the LR, EBF and ensemble methods, respectively.

422 **Table 6** AUC values of three models.

423 **Discussion**

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3 424 Due to the shortcomings and limitations of each of the quantitative (data-driven and
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5 425 knowledge-based) methods, scientists proposed and developed ensemble methods in order to
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7 426 overcome their disadvantages and increase their efficiency (Tehrany et al., 2014). In this
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9 427 study, two types of bivariate and multivariate data-driven methods, namely, EBF and LR, and
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11 428 their ensemble were applied to produce GESMs. Given that the results of data-driven
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13 429 methods are obtained from data, the input data should be reliable. In this study, early reports,
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15 430 Google Earth images and comprehensive field surveys by GPS were used to produce the
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17 431 gully erosion inventory, from which all the analyses were performed. Bivariate statistical
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19 432 methods (such as EBF) can perform a quantified prediction of sensitivity using the factor
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21 433 class weighted amounts acquired according to the distribution of events (Tehrany et al., 2013;
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23 434 Guo-liang et al., 2017). The main advantage of EBF is that, unlike other bivariate models,
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25 435 EBF supports a series of mass functions including belief, disbelief, uncertainty, and
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27 436 plausibility. Therefore, the results can adequately show quantitative relationships between
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29 437 gully occurrences and conditioning factors by modeling the degree of uncertainty.
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33 438 The main disadvantage of EBF method is its incapability to compute the weight of
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35 439 conditioning factors. EBF has been used in various research ranging from landslide
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37 440 sensitivity to groundwater potential mapping (Park et al., 2014; Wang et al., 2016; Zeinivand
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39 441 and Ghorbani Nejad, 2018). Wang et al. (2016) stated that the EBF method with AUC =
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41 442 80.09 is highly capable of predicting areas prone to landslide. Meanwhile, the main
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43 443 advantage of multivariate methods, such as LR, is its capability to evaluate the relationship
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45 444 between an occurrence and the conditioning factors. Thus, such methods enable the
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47 445 assessment of the significance and the removal of causative factors (Pourghasemi et al., 2013;
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49 446 Guo-liang et al., 2017; Raja et al., 2017).

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52 447 Pourghasemi et al. (2013) showed that binary LR is highly capable of identifying areas prone
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54 448 to landslide compared with other methods. Raja et al. (2017) stated that LR has a high
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3 449 prediction capability. The results of the LR method indicate that the factors of LU/LC,
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5 450 lithology and distance to road with coefficients of 0.693, 0.492 and 0.385, respectively, have
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7 451 the highest effects on gully occurrence. These results are consistent with those of Belayneh et
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9 452 al. (2014), Conoscenti et al. (2014), Shellberg et al. (2016) and McCloskey et al. (2016). The
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11 453 comparison of the ensemble method with the individual LR and EBF methods shows that the
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13 454 combined method has a higher prediction accuracy than the individual methods.

15 455 **Conclusion**

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18 456 This study was carried out in order to not only investigate the capability of an ensemble
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20 457 model, EBF-LR, to predict the GESM, but also compare its capability with standalone EBF
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22 458 and LR models. Over the years, researchers and natural resource managers around the world
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24 459 have been working on various types of models to assess GESM. In this research, a new
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26 460 scientific methodology framework is proposed using a combination of bivariate (evidential
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28 461 belief function, EBF) and multivariate (Logistic regression, LR) methods implemented in a
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30 462 geographical information system (GIS) to predict gully erosion-prone areas. For this purpose,
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32 463 16 gully erosion conditioning factors and 90 gully locations (training dataset) are used for
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34 464 modelling and GESM. Subsequently, 38 gully locations (validation dataset) are used for
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36 465 validation of GESMs. The resultant maps can be used for preventive measures and to reduce
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38 466 possible damages caused by them.

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41 467 In addition, the importance of all gully erosion conditioning factors was investigated based on
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43 468 all modelling approaches. The result of LR method indicates that factors of LU/LC, lithology
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45 469 and distance to road have the greatest effect on gully occurrence in the study area. The results
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47 470 of validation using AUC and SCAI indicators confirm that the proposed integrated method
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49 471 has a higher accuracy than the individual EBF and LR methods. The results show that the
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51 472 areas with very high susceptibility to gully erosion are mainly distributed in the central part
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53 473 of the study area and gully occurrence are highly predicted near the roads and rivers, flat
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3 474 topography, sparse vegetation and susceptible lithology to erosion. Unfortunately, despite the
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5 475 large dispersion of gullies and the erosion activity in the study area, no measures have been
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7 476 considered to control the growth of gullies and the risk of losses from erosion, and the most
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9 477 destructive type of erosion has not been seriously investigated by authorities. The
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11 478 revitalisation of vegetation, which increases surface roughness, improves soil, increases
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13 479 organic matter and decreases runoff; the management of human activities to prevent and
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15 480 reduce their destructive effects, especially on slopes; and the prevention of grazing in areas
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17 481 sensitive to gully erosion are proposed to reduce the rate of erosion in the study area.

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20 482 Due to the higher accuracy of the GESM using a combined approach, planners and decision-
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22 483 makers can use it to carry out developmental projects such as road construction and
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24 484 electricity and gas transmission lines in order to prevent possible damages.

25 26 485 **Conflict of Interest**

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28
29 486 Authors declare that there is no conflict of interest in regarding the publication of this paper.

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714 **TABLE 1** Description of geological units in the study area (GSI, 1997)

Class	Lithology	Age	AGE_ERA
Qft2	Low level piedmont fan and valley terrace deposits	Quaternary	CENOZOIC
PIQc	Fluvial conglomerate, Piedmont conglomerate and sandstone.	Pliocene- Quaternary	CENOZOIC
Qft1	High level piedmont fan and valley terrace deposits	Quaternary	CENOZOIC
Murm	Light - red to brown marl and gypsiferous marl with sandstone intercalations	Miocene	CENOZOIC
Mur	Red marl, gypsiferous marl, sandstone and conglomerate (Upper red Fm.)	Miocene	CENOZOIC
Qsf	Salt flat	Quaternary	CENOZOIC
Edav	Dacitic to Andesitic volcanic	Eocene	CENOZOIC
E1c	Pale-red, polygenic conglomerate and sandstone	Paleocene-Eocene	CENOZOIC

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748 **TABLE 2** Overview of factors used for Gully Erosion Susceptibility Mapping.

Factor	Range		Classes	Method
	min	max		
Slope	0.00	18.158	1. (<1°), 2- (1-2°), 3. (2-3°), 4. (3-5°), 5. (>5°)	Natural break
Aspect	-1	360	1. Flat (-1°), 2. North (337.5-360°, 0-22.5°), 3. Northeast (22.5- 67.5°), 4. East (67.5-112.5°), 5. Southeast (112.5-157.5°), 6. South (157.5-202.5°), 7. Southwest (202.5-247.5°), 8. West (247.4-292.5°), and 9. Northwest (292.5-337.5°)	Equal interval
Plan curvature	-0.01	0.009	1. Concave (< 0), 2. Flat (0), 3. Convex (> 0)	Natural break
Profile curvature	-0.01	0.013	1. (-0.01 - -0.001), 2. (-0.001- -0.0005), 3. (-0.0005-0.0005), 4. (0.0005- 0.001), 5. (0.001-0.013)	Natural break
LS	0	457.37	1. (<50m), 2. (50-150m), 3. (150-250m), 4. (250-350m), 5. (>350m)	Natural break
NDVI	-0.112	0.403	1. (<0.039), 2. (0.039 - 0.1), 3. (>0.130)	Natural break
TWI	-3.22	6.907	1. (< -0.84), 2. (-0.84 - 0.94), 3. (0.94 - 4), 4. (>4)	Natural break
SPI	0.713	24294	1. (<300), 2. (300 - 1500), 3. (1500 - 3000), 4. (3000 - 7000), 5. (7000 - 11000), 6. (>11000)	Natural break
Drainage density	1.190	3.246	1. (<1.86Km/Km2), 2. (1.86 - 2.19 Km/Km2), 3. (2.19 - 2.54K m/Km2), 4. (> 2.54 Km/Km2)	Natural break
Distance to road	0	13684.7	1. (<500m), 2. (500 - 1000m), 3. (1000 - 1500m), 4. (1500 - 2000m), 5. (>2000m)	Natural break
Distance to river	0	636.39	1. (<100m), 2. (100 - 200m), 3. (200 - 300m), 4. (>300m)	Natural break
LU/LC	-	-	1. (Irrigated lands), 2. (Desert), 3. (Pasture), 4. (salty lands)	Supervised classification
Lithology	-	-	1. (Qft2), 2. (PlQc), 3. (Qft1), 4. (Murm), 5. (Mur), 6. (Qsf), 7. (Edav), 8. (E1c)	Lithological units
Rainfall	73.17	93.70	1. (< 77.60 mm), 2. (77.60 - 81.70mm), 3. (81.70 - 85.41mm), 4. (85.41 - 88.79mm), 5. (>88.79mm)	Natural break
Soil type	-	-	1. (Entisols/Aridisols), 2. (Salty flats), 3. (Rocky lands)	Supervised classification
Convergence	-100	100	1. (< -30), 2. (-30 - -2), 3. (-2 - 30), 4. (> 30)	Natural break

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758 **TABLE 3** Multicollinearity test among conditioning factors.

Factors	Multi-collinearity test		Factors	Multi-collinearity test	
	Tolerance	VIF		Tolerance	VIF
Lithology	0.632	1.582	Plan curvature	0.773	1.294
LU/LC	0.483	2.070	NDVI	0.398	2.511
Soil type	0.353	2.829	LS	0.718	1.394
Convergence	0.784	1.275	Rainfall	0.723	1.382
Drainage density	0.290	3.452	Slope degree	0.583	1.716
Distance to road	0.296	3.382	SPI	0.616	1.623
Distance to stream	0.903	1.107	TWI	0.818	1.222
Profile curvature	0.837	1.194	Aspect	0.918	1.089

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764 **TABLE 4** Spatial relationship between gully occurrence and conditioning factors

Factors	Classes	Pixel of domain		Pixel of gullies		EBF			
		N	%	N	%	BEL	DIS	UNC	PLS
Slope (degree)	< 1	114459	45.02	393	55.51	0.31	0.16	0.53	0.84
	1 - 2	91337	35.93	226	31.92	0.22	0.21	0.56	0.79
	2 - 3	32337	12.72	59	8.33	0.16	0.21	0.62	0.79
	3 - 5	12799	5.03	25	3.53	0.17	0.21	0.62	0.79
	> 5	3297	1.30	5	0.71	0.14	0.20	0.66	0.80
Slope Aspect	F	15826	6.25	66	9.59	0.20	0.14	0.66	0.86
	N	9417	3.72	16	2.33	0.08	0.15	0.77	0.85
	NE	25783	10.17	49	7.12	0.09	0.15	0.76	0.85
	E	46880	18.50	105	15.26	0.11	0.15	0.74	0.85
	SE	48387	19.09	216	31.40	0.22	0.12	0.66	0.88
	S	36258	14.31	145	21.08	0.20	0.13	0.67	0.87
	SW	32005	12.63	62	9.01	0.09	0.15	0.75	0.85
	W	20605	8.13	23	3.34	0.05	0.15	0.79	0.85
	NW	15135	5.97	21	3.05	0.07	0.15	0.78	0.85
Plan Curvature (100/m)	concave	51659	20.32	141	19.92	0.34	0.34	0.32	0.66
	flat	148111	58.26	438	61.86	0.37	0.31	0.32	0.69
	convex	54458	21.42	129	18.22	0.29	0.35	0.35	0.65
Profile Curvature (100/m)	-0.01 - -0.001	7735	3.04	21	2.95	0.21	0.20	0.59	0.80
	-0.001 - -0.0005	71188	28.00	209	29.35	0.23	0.20	0.58	0.80
	-0.0005 - 0.0005	99643	39.19	283	39.75	0.22	0.20	0.58	0.80
	0.0005 - 0.001	68560	26.97	187	26.26	0.21	0.20	0.59	0.80
	> 0.001	7102	2.79	12	1.69	0.13	0.20	0.67	0.80
NDVI	> 0.039	107854	42.52	214	29.85	0.36	0.42	0.22	0.58
	0.039 - 0.130	144980	57.16	503	70.15	0.64	0.24	0.13	0.76
	>0.130	792	0.31	0	0	0	0.34	0.66	0.66
LU/LC	Irrigated lands	2349	0.93	0	0	0	0.25	0.75	0.75
	Desert	110530	43.62	502	72.97	0.77	0.12	0.11	0.88
	Pasture	138239	54.55	186	27.03	0.23	0.39	0.38	0.61
	salty lands	2288	0.90	0	0	0	0.25	0.75	0.75
	Qft2	97841	38.61	237	34.45	0.13	0.15	0.72	0.85
	PlQc	37999	15.00	136	19.77	0.19	0.13	0.67	0.87
	Qft1	28107	11.09	144	20.93	0.28	0.13	0.60	0.87
	Murm	35924	14.18	116	16.86	0.17	0.14	0.69	0.86
	Mur	13160	5.19	55	7.99	0.23	0.14	0.64	0.86
Lithology	Qsf	35824	14.14	0	0	0	0.17	0.83	0.83
	Edav	4333	1.71	0	0	0	0.14	0.86	0.86
	E1c	218	0.09	0	0	0	0.14	0.86	0.86
	< 500	23651	9.30	275	38.68	0.45	0.11	0.44	0.89
	500 - 1000 (m)	19078	7.50	131	18.42	0.26	0.15	0.59	0.85
	1000 - 1500 (m)	15096	5.94	65	9.14	0.16	0.16	0.68	0.84
	1500 - 2000 (m)	14317	5.63	31	4.36	0.08	0.17	0.75	0.83
	> 2000 (m)	182133	71.63	209	29.40	0.04	0.41	0.54	0.59
	< 100 (m)	112480	44.24	322	45.29	0.27	0.25	0.49	0.75
Distance to river (m)	100 - 200	74160	29.17	198	27.85	0.25	0.26	0.50	0.74
	200 - 300	48112	18.92	150	21.10	0.29	0.24	0.47	0.76
	>300	19523	7.68	41	5.77	0.20	0.26	0.55	0.74
Drainage density (km/km2)	< 1.86	29875	11.75	37	5.20	0.12	0.27	0.60	0.73
	1.86 - 2.19	94154	37.03	273	38.40	0.29	0.25	0.46	0.75
	2.19 - 2.54	76868	30.23	304	42.76	0.40	0.21	0.39	0.79
	> 2.54	53378	20.99	97	13.64	0.18	0.28	0.54	0.72

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Factors	Classes	Pixel of domain		Pixel of gullies		EBF			
		N	%	N	%	BEL	DIS	UNC	PLS
	< 75	52972	20.84	24	3.37	0.03	0.25	0.72	0.75
	75 - 80	38491	15.14	140	19.66	0.28	0.19	0.53	0.81
Rainfall (mm)	80 - 85	58041	22.83	314	44.10	0.41	0.15	0.44	0.85
	85 - 90	66082	25.99	219	30.76	0.25	0.19	0.56	0.81
	> 90	38643	15.20	15	2.11	0.03	0.23	0.74	0.77
	< -0.84	38462	15.13	111	15.68	0.25	0.25	0.50	0.75
TWI	-0.84 - 0.94	53947	21.22	174	24.58	0.28	0.24	0.48	0.76
	0.94 - 4	33355	13.12	87	12.29	0.23	0.25	0.52	0.75
	> 4	128465	50.53	336	47.46	0.23	0.26	0.51	0.74
	< 300	165290	65.02	428	60.45	0.15	0.19	0.66	0.81
	300 - 1500	34236	13.47	106	14.97	0.18	0.16	0.66	0.84
SPI	1500 - 3000	29230	11.50	95	13.42	0.19	0.16	0.65	0.84
	3000 - 7000	15491	6.09	35	4.94	0.13	0.17	0.70	0.83
	7000 - 11000	8927	3.51	43	6.07	0.28	0.16	0.56	0.84
	> 11000	1055	0.41	1	0.14	0.06	0.16	0.78	0.84
	< 50	70459	27.71	161	22.74	0.16	0.21	0.63	0.79
	50 - 150	35370	13.91	99	13.98	0.20	0.20	0.60	0.80
LS (m)	150 - 250	46128	18.14	130	18.36	0.20	0.20	0.60	0.80
	250 - 350	60273	23.71	179	25.28	0.21	0.20	0.59	0.80
	> 350	41999	16.52	139	19.63	0.23	0.19	0.57	0.81
	Entisols/Aridisols	216912	85.60	680	98.84	0.93	0.04	0.03	0.96
Soil type	Salty flats	372	0.15	0	0	0	0.45	0.55	0.55
	Rocky lands	36122	14.25	8	1.16	0.07	0.52	0.42	0.48
	< -30	30848	12.13	104	14.69	0.29	0.24	0.47	0.76
Cinvergence index (100/m)	-30 - -2	82064	32.28	256	36.16	0.27	0.23	0.50	0.77
	-2 - 30	104966	41.29	250	35.31	0.21	0.27	0.52	0.73
	> 30	36351	14.30	98	13.84	0.23	0.25	0.52	0.75

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788 **TABLE 5** Area of susceptibility classes and seed cell area index (SCAI).

Models	Classes	Total area (km ²)		Testing gullies (km ²)		Validation gullies (km ²)		Sum	SCAI
		Area	%	Area	%	Area	%		
LR	Very Low	40.20	17.57	0.03	1.41	0.00	0.00	1.41	12.49
	Low	70.62	30.86	0.31	17.58	0.01	0.00	17.58	1.76
	Moderate	41.62	18.19	0.40	22.36	0.34	0.09	22.45	0.81
	High	39.50	17.26	0.14	7.74	0.22	0.06	7.79	2.21
	Very High	36.91	16.13	0.91	50.91	0.38	0.10	51.01	0.32
EBF	Very Low	44.11	19.35	0.03	1.83	0.01	0.00	1.84	10.54
	Low	46.70	20.48	0.13	7.19	0.04	0.01	7.20	2.84
	Moderate	45.22	19.84	0.31	17.21	0.11	0.03	17.24	1.15
	High	45.54	19.98	0.47	26.66	0.27	0.07	26.73	0.75
	Very High	46.41	20.36	0.84	47.11	0.52	0.14	47.25	0.43
Ensemble	Very Low	44.77	19.64	0.02	1.13	0.00	0.00	1.13	17.39
	Low	46.12	20.23	0.09	4.94	0.01	0.00	4.94	4.10
	Moderate	45.92	20.14	0.27	14.95	0.03	0.01	14.96	1.35
	High	44.91	19.70	0.37	20.87	0.19	0.05	20.93	0.94
	Very High	46.25	20.29	1.03	58.11	0.71	0.19	58.30	0.35

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795 **TABLE 6** Area under the curve (AUC) value of three models.

Models	Area	Standard error	Asymptotic significant	Asymptotic 95% confidence interval	
				Lower bound	Upper bound
Ensemble	0.909	0.008	0.000	0.893	0.925
EBF	0.821	0.011	0.000	0.799	0.844
LR	0.802	0.012	0.000	0.779	0.825

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For Peer Review

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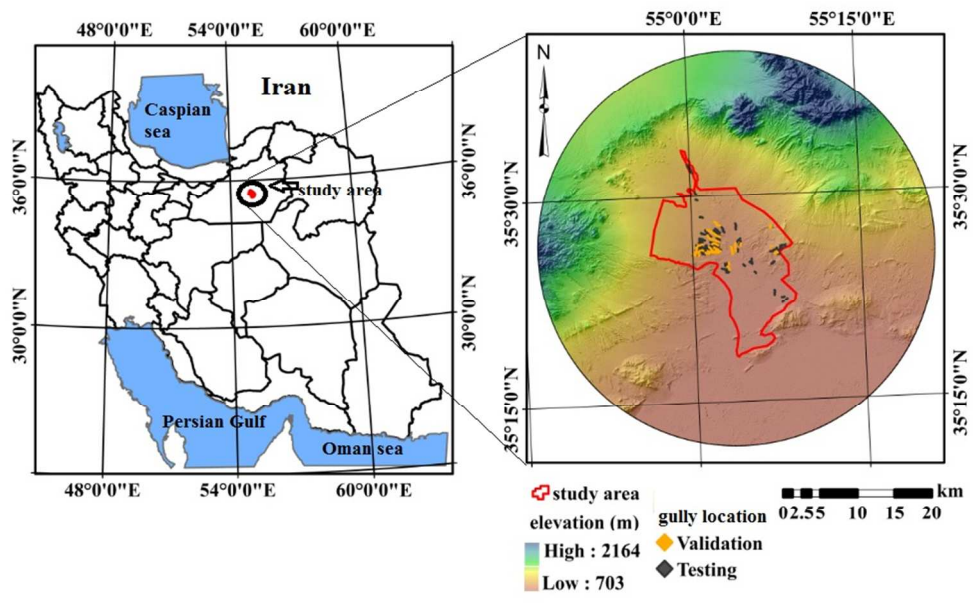


Fig. 1 Study area.

265x166mm (96 x 96 DPI)

Review

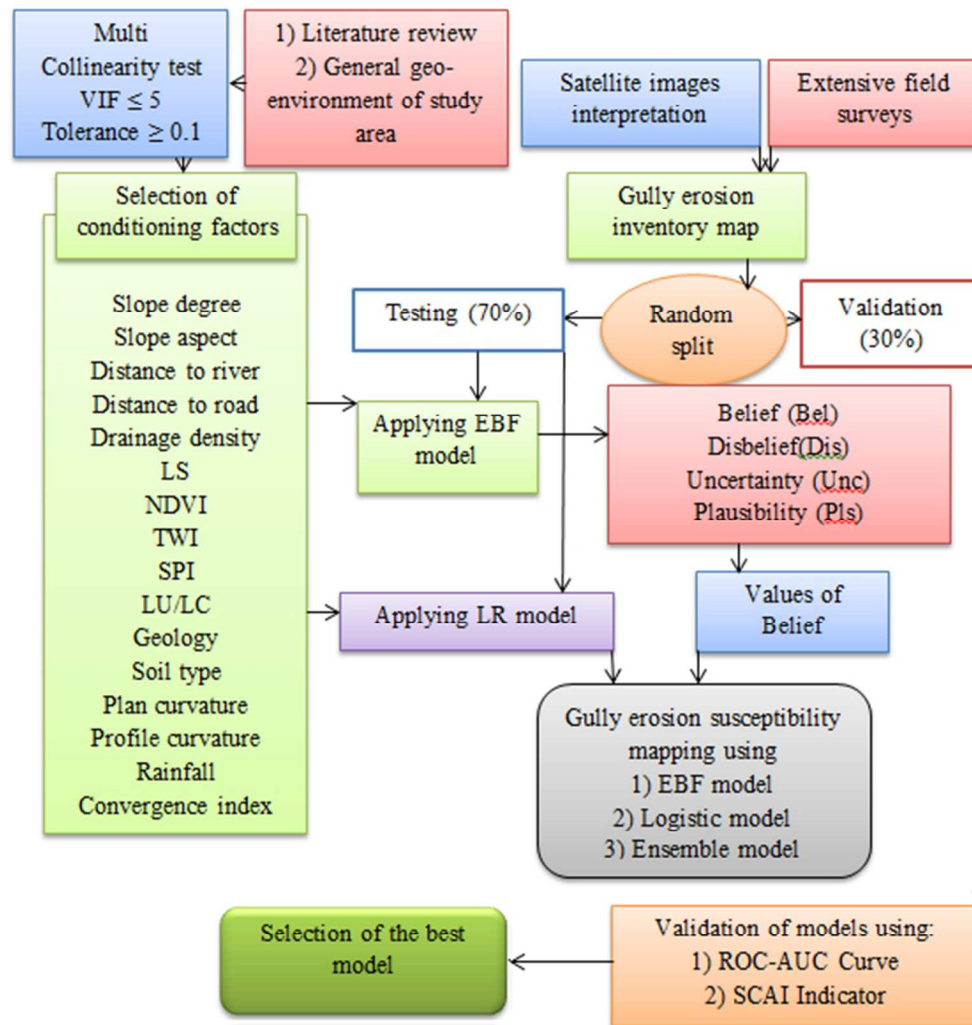


Fig. 2 Flowchart of the methodology used.

150x155mm (96 x 96 DPI)

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Fig. 3 Field photographs showing identified gullies in the study area.

216x197mm (220 x 220 DPI)

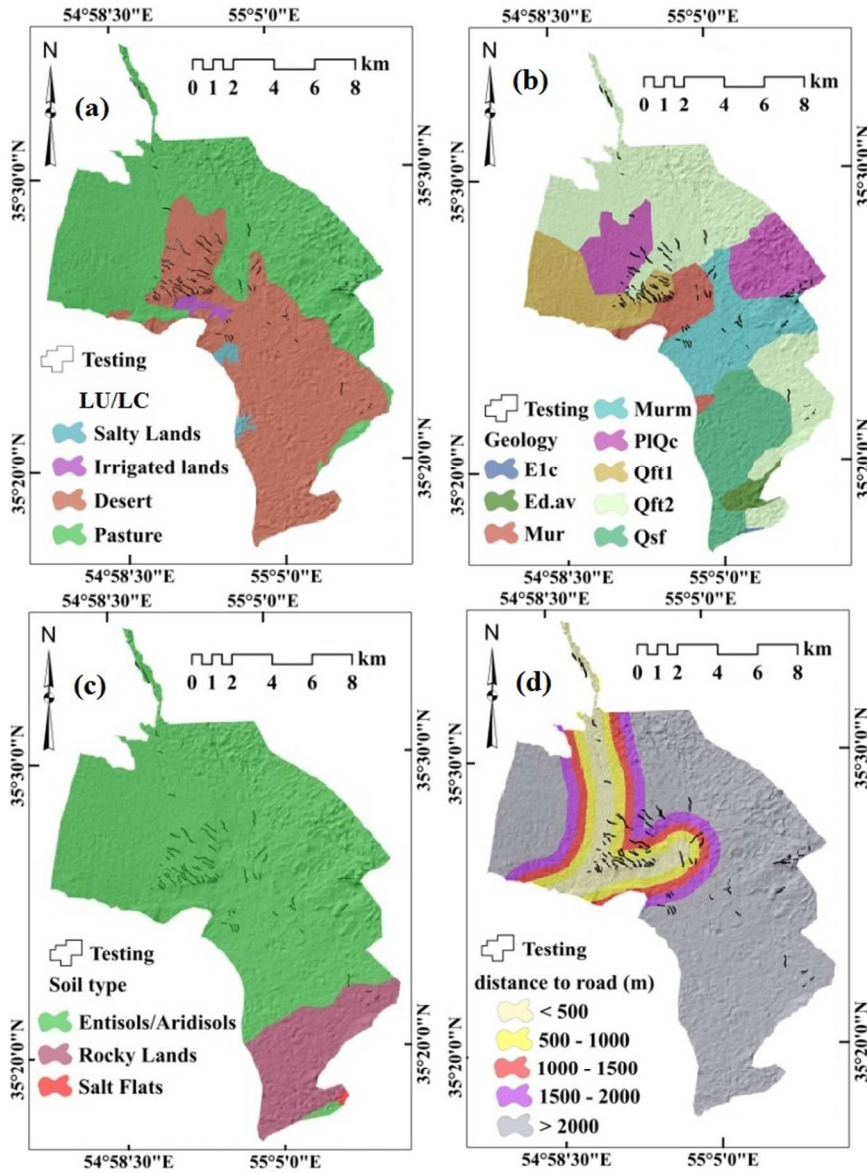


Fig. 4 Some examples of gully erosion conditioning factors, a) Land use/Land cover, b) lithology, c) soil type, and d) distance to road.

214x288mm (96 x 96 DPI)

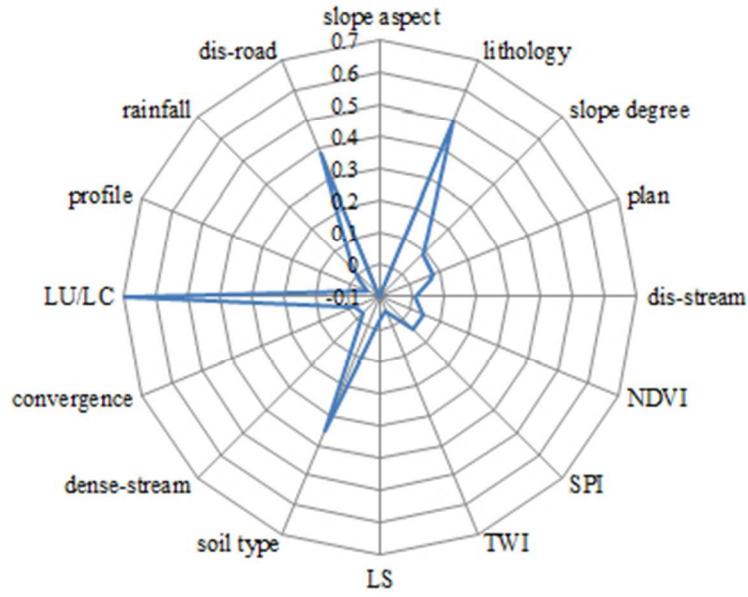


Fig. 5 Coefficients of conditioning factors in the logistic regression model.

101x79mm (96 x 96 DPI)

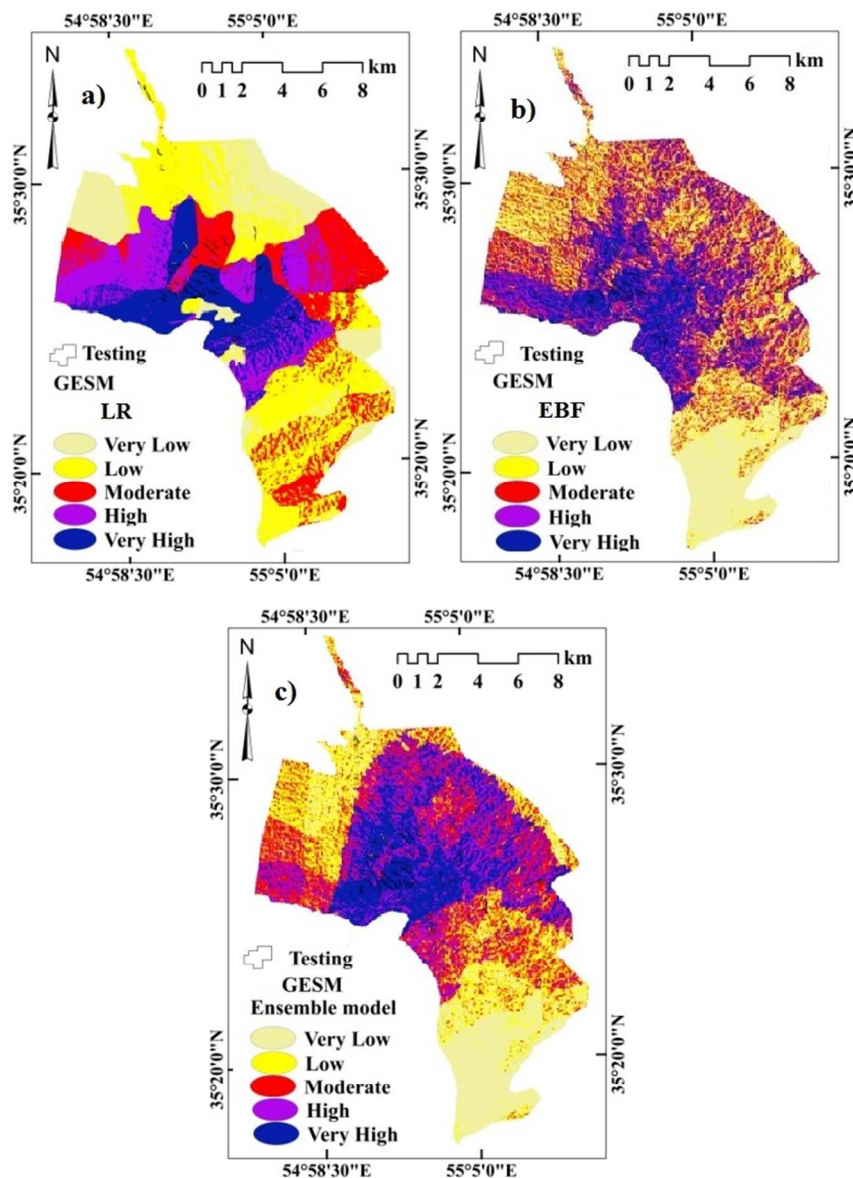


Fig. 6 Gully erosion susceptibility mapping using Logistic Regression (LR), Evidential Belief Function (EBF), and ensemble models.

212x292mm (96 x 96 DPI)