

Assessing the risk of ocean acidification for scleractinian corals on the Great Barrier Reef

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A thesis submitted in fulfillment of the requirements for the degree of Doctor of Philosophy at The University of Technology Sydney in 2013 School of the Environment

Certificate of original authorship

I certify that the work in this thesis has not previously been submitted for a degree nor has it been submitted as part of requirements for a degree except as fully acknowledged within the text.

I also certify that the thesis has been written by me. Any help that I have received in my research work and the preparation of the thesis itself has been acknowledged. In addition, I certify that all information sources and literature used are indicated in the thesis.

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Date:

This thesis is dedicated to my son, Javier Patrick Lloyd. His arrival into this world during my PhD was a timely reminder that we do not inherit this earth from our ancestors; we borrow it from our children. Javier, I wrote this thesis in the hope that it will serve to maintain the beauty of the Great Barrier Reef for your generation and many more to come.

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Abbreviations

α: photosynthetic efficiency at sub-saturating irradiance

ATP: adenosine triphosphate

ADP: adenosine diphosphate

A_T: total alkalinity

CA: carbonic anhydrase

CAP: canonical analysis of principle coordinates

CCA: crustose coralline algae

CCM: carbon concentration mechanism

Chl a: Chlorophyll a

CPCe: Coral Point Count with Excel extensions

DBL: diffusion boundary layer

DGGE: denaturing gradient gel electrophoresis

DbRDA: distance based redundancy analysis

Dd: diadinoxanthin

DIC: dissolved inorganic carbon

DISTLM: distance based linear model

Dt: diatoxanthin

E_k: minimum saturating irradiance

ECM: extracellular calcifying medium

fCO₂: fugacity CO₂

 F_0 : minimum fluorescence

 F_m : maximum fluorescence

 F_m ': light-adapted maximum fluorescence

 F_t : steady-state fluorescence

 ΔF : light-adapted variable fluorescence

 F_{ν} : variable fluorescence

 F_{ν}/F_m : maximum dark-adapted quantum yield of PSII

 $\Delta F/F_m$ ': effective dynamic quantum yield of PSII

FP: fluorescent proteins

G3P: glyceraldehyde 3-phosphate

GFP: green fluorescent protein

HIRS: Heron Island Research Station

HPLC: high performance liquid chromatography

iPAM: imaging pulse amplitude modulation fluorometer

IPCC: Intergovernmental Panel on Climate Change

ITS1/2: internal transcribed spacer regions of rDNA

 K'_{sp} : solubility constant

LEC: light enhanced calcification

LEDR: light enhanced dark respiration

LHC: light harvesting complex

LIRS: Lizard Island Research Station

LR: Loomis Reef

MC: Mermaid Cove

NAD(P)⁺: nicotinamide adenine dinucleotide (phosphate)

NADPH: nicotinamide adenine dinucleotide phosphate

NPQ: non-photochemical quenching

OA: ocean acidification

 $\Omega_{\rm calc}$: saturation state of calcite

 Ω_{arag} : saturation state of aragonite

P₆₈₀: PSII primary donor

PCR: polymerase chain reaction

PG: phsophoglycolate

PGA: phosphoglycerate

PGPase: phosphoglycolate phosphatase

Phe: phaeophytin

P_{net}: net photosynthesis

P_nmax: maximum rate of net photosynthesis

pCO₂: partial pressure of carbon dioxide

P:E curve: photosynthesis-irradiance curve

PERMANOVA: permutational analysis of variance

P_g: gross photosynthesis

P_g:R: gross photosynthesis to respiration ratio

 $pH = -log[H^+]$

pH_T: total pH

PQ: plastoquinone

PRIMER: Plymouth routines in multivariate ecological research

PSII: photosystem II

PSU: practical salinity unit

Q_A: quinone acceptor

Q_m: excitation pressure

RC: reaction centre

R_{dark}: dark respiration

ROS: reactive oxygen species

RuBisCO: ribulose bisphosphate carboxylase/oxygenase

RuBP: ribulose-1,5-bisphosphate

SR: Station Reef

SST: sea surface temperatures

 Φ_{NO} : yield of fluorescence

 Φ_{NPQ} : yield of non-photochemical quenching

Abstract

Ocean acidification will impact the photo-physiology of reef-building corals as it can lead to dysfunction of the symbiosis and loss of productivity. The major objective of this thesis was to provide insight into the mechanism of CO₂-induced bleaching and productivity loss across multiple life-history stages and interpret these findings in an ecological context.

Chapter 1 provides a review of the literature investigating the photo-physiological impact of ocean acidification, emphasizing the experimental conditions in studies that observed *Symbiodinium* dysfunction and productivity loss. Chapter 2 presents a working hypothesis to describe the fundamental physiological aspects of coral bleaching under ocean acidification. This research investigates the response of *Acropora aspera* using pulse amplitude modulation (PAM) fluorometry and oxygen respirometry under increased pCO₂ with concomitant high light conditions. The dinoflagellate density and HPLC pigment analysis are utilised to characterise the CO₂-induced bleaching response. We present a conceptual model linking photorespiration to CO₂-induced bleaching and productivity loss.

The impact of ocean acidification on coral reef ecosystems is likely to deviate from oceanic climate models due to diel modification of carbonate chemistry by community metabolism. Chapter 3 characterises the diurnal variation in carbonate chemistry at sites around Lizard Island and links this to the ocean acidification response of *Acropora millepora* collected from these sites. Furthermore, we utilise permutational multivariate statistical analyses to partition the variation in carbonate chemistry attributable to community composition at these sites. It was hypothesized that greater diurnal variation in carbonate chemistry may improve resilience of scleractinian corals to future ocean acidification conditions. This chapter highlights that site-specific physiological trade-offs may influence the response of reef-building corals to future ocean acidification scenarios.

Chapter 4 reports a visual bleaching response in *A. millepora* juveniles under future ocean acidification conditions. The effect of ocean acidification on coral juveniles is

hypothesised to impact *Symbiodinium* uptake and photochemical efficiency. We utilised the iPAM to align the photochemistry in the juveniles with their visual bleaching response and *Symbiodnium* type, as assessed by denaturing gradient gel electrophoresis (DGGE) of the internal transcribed spacer region 1 (ITS1) of the ribosomal genes. This study links the bleaching response with recruits containing a dominant population of *Symbiodinium* type D1 or D1-4, with potential implications for post-settlement survivorship and population dynamics.

Lastly, in Chapter 5 the key findings of this thesis are discussed in light of the ecological implications for the Great Barrier Reef. The synopsis outlines the effect of ocean acidification on the photo-physiology, productivity, calcification, reproduction and symbiont acquisition of reef-building corals. Future avenues for research are suggested based on new research gaps revealed by this thesis with the aim to continue to provide up-to-date scientific information to policy makers and reef managers.

Chapter 1: General Introduction

Reef-building corals and their threatened ecosystem

The value of coral reefs span biological, social and economic boundaries. They are among the most productive and diverse ecosystems on the planet, providing humanity with services such as coastal protection, building materials and novel biochemical compounds and providing economic returns through the fisheries and tourism industries (Costanza et al. 1997; Moberg and Folke 1999). In addition, coral reefs provide aesthetic and cultural benefits, which remain difficult to quantify. Nevertheless, the economic value of the Great Barrier Reef fisheries and tourism industries were estimated to be AU\$ 912 million in 1996 (Driml 1999) and the global economic value of coral reefs has been estimated up to US\$ 600,000 per km² (UNEP-WCMC 2006). Therefore, it is not surprising that the loss of coral reefs would amount to hundreds of billions of dollars per year worldwide (Bryant et al. 1998; Hoegh-Guldberg 1999). Yet despite their value and importance, coral reefs are declining due a number of direct anthropogenic influences including eutrophication, sedimentation, destructive fishing, overfishing and mining (Bruno and Selig 2007). In addition, anthropogenic CO₂ emissions threaten coral reefs due to i) the radiative scattering effect of CO2 in the atmosphere, which raises the sea surface temperature (SST) and ii) ocean acidification (OA) as CO2 dissolves into oceanic waters and reduces the availability of dissolved inorganic carbon (DIC) species to marine organisms (Caldeira and Wickett 2003; Canadell et al. 2007; Hoegh-Guldberg et al. 2007).

Scleractinian corals are largely responsible for the reef framework through deposition of their limestone skeleton, creating a structurally complex habitat for a diverse range of reef organisms (Smith 1983). Branching corals in the family Acroporidae are known to be fast growing, averaging 10 cm² growth per year and therefore significantly contributing to reef formation (Connell 2012). The taxonomic classification of the Acroporids investigated in the experimental chapters of this thesis is shown below:

Kingdom: Animalia

Phylum: Cnidaria

Class: Anthozoa

Subclass: Zoantharia

Order: Scleractinia

Family: Acroporidae

Genus: Acropora

These reef-building corals are also referred to as the coral holobiont which encompasses all symbionts of the coral animal, including dinoflagellates (genus: Symbiodinium), bacteria, fungi, protozoa and endolithic algae (Rohwer et al. 2002; Rosenberg et al. 2007). The taxonomic classification of the symbiotic dinoflagellates is listed below:

Kingdom: Protozoa

Division: Dinoflagellata

Subdivision: Dinokaryota

Subclass: Gymnodiniphycidae

Order: Suessiales

Family: Symbiodiniaceae

Genus: Symbiodinium

The exchange of nutrients and energy between the symbiotic dinoflagellate and the host coral is a fundamental factor contributing to the success of this symbiosis, and enables the widespread formation of coral reefs around the world's tropical and subtropical seas. The dinoflagellates are protected in the gastrodermal cells of the coral where they receive the CO₂, nitrogen and phosphorus essential for their photosynthesis and growth (Muscatine and D'Elia 1978). In return, the symbionts can translocate up to 95% of their photosynthate, principally as glucose or glycerol, which significantly contributes to the energy budget of the coral host (Muscatine 1967). This mutualistic symbiosis can be disrupted under certain physical and chemical conditions including changes in light, temperature or CO₂. During these suboptimal conditions, the coral may "bleach" due to expulsion of the dinoflagellates or loss of their photosynthetic pigments. This process represents a major threat to coral reef

ecosystems, as coral growth, survivorship and reproduction are diminished during bleaching episodes (Little et al. 2004; Jones and Berkelmans 2011).

Acropora aspera (blue morph) (Dana, 1846)

Acropora aspera is a staghorn coral with digitate, upright branches, which divide infrequently and taper towards the end (Figure 1.1). The axial corallites are domeshaped while the radial corallites, exhibiting a variety of sizes, are appressed and thick-walled with small openings (Veron 2000). A. aspera branch ends may be pale brown or pale blue due to the presence of host pigment, pocilloporin, a type of Green Fluorescent Protein (GFP) (Dove 2004). In general, fast-growing corals such as arborescent acroporids are more susceptible to thermal bleaching than slow-growing corals (Marshall and Baird 2000), while A. aspera (blue morph) may be hypersensitive due to the vulnerability of pigment synthesis during heat stress (Dove 2004).



Figure 1.1: Colony of *Acropora aspera* (blue morph) at Heron Island showing varied size of the radial corallites and blue colouration around the axial corallite due to the host pigment, pocilloporin. Scale bar = 5cm. (Photo: R Middlebrook).

Acropora millepora (red morph) (Ehrenberg, 1834)

Acropora millepora is a corymbose coral with short branches of uniform shape (Figure 1.2). The axial corallites are distinctly tubular while the radial corallites are compact and of the same size with a scale-like appearance (Veron 2000). The branches of A. millepora may be green, yellow or red due to the presence of host fluorescent proteins (FP) with the red morph containing the least amount of FP and being most susceptible to thermal bleaching (Paley and Bay 2012). Overall, it has been shown that A. millepora is susceptible to thermal bleaching, but not to as great an extent as A. aspera (Fisher et al. 2012).



Figure 1.2: Colony of *Acropora millepora* (red morph) at Lizard Island showing corymbose colony shape and tubular axial corallites. Scale bar = 5cm.

Ocean Acidification on Coral Reefs

Increased anthropogenic CO₂ emissions, due to fossil fuel burning, cement production and land-use changes, are causing global climate change and ocean acidification (Caldeira and Wickett 2003; Feely et al. 2004; Orr et al. 2005). Atmospheric CO₂ equilibrates with the sea surface through air-sea gas exchange and this leads to an increase in DIC and a decline in the pH of seawater (Sabine et al. 2004; Doney et al. 2009). The seawater carbonate chemistry is governed by the following reversible chemical reactions:

$$CO_2 + H_2O \rightleftharpoons H_2CO_3 \rightleftharpoons HCO_3^- + H^+ \rightleftharpoons CO_3^{2-} + 2H^+$$
 (1)

The dissolution of CO₂ in seawater forms carbonic acid (H₂CO₃), a weak acid that rapidly dissociates into hydrogen (H⁺), bicarbonate (HCO₃⁻) and carbonate (CO₃²⁻) ions. To balance the equilibrium, some of the additional H^+ reacts with $CO_3^{\,2-}$ forming more HCO₃⁻. Overall, the net effect of adding CO₂ to seawater is an increase in HCO₃⁻ and a decrease in CO_3^{2-} , along with lowered pH (pH = $-\log [H^+]$) (Doney et al. 2009). The Intergovernmental Panel on Climate Change (IPCC) predicts that oceanic pH may drop 0.3 - 0.5 pH units by the end of the century, depending on the CO₂ emission scenario (Zeebe and Wolf-Gladrow 2001; Caldeira and Wickett 2005; IPCC 2007). Prior to the industrial revolution the atmospheric partial pressure of CO₂ (pCO₂) was 280 μatm and Total pH (pH_T) was 8.16; yet today the pCO₂ has risen to 400 μatm, which represents a 43% increase in pCO₂ (Jones 2013). In addition, pH has declined to 8.04 due to the \sim 30% increase in H⁺ (Kleypas et al. 2006). These changes in H⁺ and pCO₂ may have important implications for extracellular and intracellular acid-base regulation, ion-transport mechanisms and the occurrence of hypercapnia in marine organisms (Pörtner 2008). As the pH and CO₃²⁻ concentration decline, so too does the saturation state of the seawater with respect to calcite (Ω_{calc}) and aragonite (Ω_{arag}). Calcium carbonate (CaCO₃) mineral formation, imperative to the shells and skeletons of many marine organisms, is dependent on Ω , which is defined as:

$$\Omega = [Ca^{2+}] [CO_3^{2-}] / K'_{sp}$$
 (2)

where K'_{sp} is the apparent solubility constant for calcite or aragonite, which depends on temperature, salinity and pressure. Aragonite is the CaCO₃ polymorph found in scleractinian coral skeletons and as K'_{sp} (aragonite) is approximately double K'_{sp} (calcite), aragonite is ~50% more soluble than calcite (Doney et al. 2009). Although dissolution of aragonite is expected to occur in under-saturated waters ($\Omega_{arag} < 1$), calcification of particular species and net accretion of coral communities can cease even in super-saturated water (Langdon and Atkinson 2005; Silverman et al. 2009).

Throughout the oceans, marine life has adapted to the ambient CO₂ conditions, which includes the high concentrations at underwater volcanic vents and the highly variable levels in shallow reef waters (Pörtner 2008); yet the current rate of increase in pCO₂ is

approximately 1000x faster than has occurred in the past 420,000 years (Hoegh-Guldberg et al. 2007). Actually, there are no geological records of a decline in Ω_{arag} comparable to the present rate, although the pCO₂ has previously reached 20x preindustrial levels during steady state conditions whereby slow geochemical feedbacks (ie. rock-weathering) allowed Total Alkalinity (A_T) to increase (Veron 2008; Pandolfi et al. 2011). A_T represents the ability of the seawater to buffer changes in pH upon addition of acid and is equal to the sum of all proton acceptors less proton donors (Dickson et al. 2007):

$$A_{T} = [HCO_{3}^{-}] + 2[CO_{3}^{2-}] + [B(OH)_{4}^{-}] + [OH_{-}] + [HPO_{4}^{2-}] + 2[PO_{4}^{3-}] + [SiO(OH)^{3-}] + [NH_{3}] + [HS_{-}] + ... - [H_{-}^{+}]_{F} - [HSO_{4}^{-}] - [HF] - [H_{3}PO_{4}] - ...$$
(3)

Despite this potential buffering, coral reefs have ceased to accrete in the past in alignment with geological records of mass extinction during OA and warming events (Veron 2008).

Large diel variability in carbonate chemistry has been observed on coral reefs (Ohde and vanWoesik 1999; Yates and Halley 2006; Shamberger et al. 2011; Shaw et al. 2012) and with increasing anthropogenic pCO₂, the magnitude of this fluctuation is projected to increase due to changes in the ratio of DIC to A_T (Egleston et al. 2010; Shaw et al. 2013). The carbonate chemistry within the benthic boundary layer will be affected by upstream community metabolism and this may have implications for the pH gradient within the diffusive boundary layer (DBL) (Shashar et al. 1996; Hurd et al. 2011; Jokiel 2011). During the day, net photosynthesis on coral reefs removes CO₂ from the water, which leads to an increase in pH (Figure 1.3). In contrast, the net respiration occurring at night releases CO₂ into the water and reduces the pH. Calcification and dissolution affects both the pH and A_T as calcification releases CO₂ and removes CO₃²⁻ from the water. Overnight accumulation of CO₂ in shallow reef waters can drop the pH to levels as low as those predicted by the end of the century by the IPCC (Kline et al. 2012; Shaw et al. 2012).

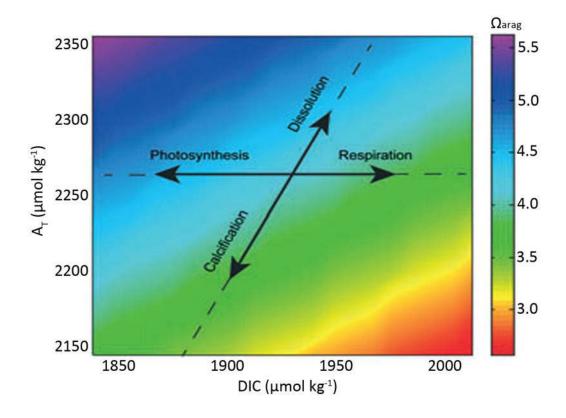


Figure 1.3: Deffeyes diagram showing iso-contours of the saturation state of aragonite (Ω_{arag}) plotted as a function of Dissolved Inorganic Carbon (DIC) and Total Alkalinity (A_T) (Anthony et al. 2011a). Photosynthesis removes CO₂ thereby reducing DIC, whereas calcification removes HCO₃⁻ and CO₃²⁻ which impacts both DIC and A_T.

The large diel variability of carbonate chemistry on coral reefs is at odds with climate models of OA that are based on constant A_T (Kleypas et al. 1999; Silverman et al. 2009; Shaw et al. 2013), hence the risks associated with OA are difficult to predict. Recent models and experiments have begun to investigate the diel variation, with evidence that day-time decreases in pCO₂ may provide some relief (Dufault et al. 2012) and predictions that the diel variation will increase on shallow reef flats (Shaw et al. 2013). Previous models that did not incorporate these diel changes are likely invalid (Silverman et al. 2009), as are models without projected changes in community composition, as this is now known to have strong implications for the diel variation in carbonate chemistry (Anthony et al. 2011b; Kleypas et al. 2011) and will occur under future OA conditions (Hall-Spencer et al. 2008; Inoue et al. 2013).

Impact of ocean acidification on the coral symbiosis and bleaching

OA has the potential to impact the coral symbiosis directly through increased pCO₂, HCO₃⁻ and H⁺, which are important to physiological processes such as photosynthesis, acid-base regulation and ion-transport mechanisms (Badger et al. 1998; Pörtner 2008). Previous research has predominately focused on the radiative forcing of atmospheric CO₂, which increases the SST and is linked to thermal coral bleaching events (Glynn 1991; Hoegh-Guldberg 1999). Yet recent research has indicated that OA may act synergistically with warming to lower the thermal bleaching threshold of corals (Pecheux 2002; Anthony et al. 2008).

For thermally induced bleaching, the initial sites of dysfunction are proposed to be either the D1 protein within the reaction centre (RC) of Photosystem II (PSII) or impairment of CO₂ fixation by Ribulose bisphosphate carboxylase/oxygenase (RuBisCO), both of which lead to damage to PSII within the symbiotic dinoflagellates (Iglesias-Prieto et al. 1992; Warner et al. 1996; Jones et al. 1998; Hill et al. 2004; Lilley et al. 2010). Under these conceptual models of bleaching, there is reduced translocation of photosynthates to the host and production of reactive oxygen species (ROS) due to overwhelmed photo-protective and anti-oxidant mechanisms (Iglesias-Prieto et al. 1992; Shick et al. 1995). Following this damage, a signal transduction is initiated that ends with host cell death or apoptosis (Dunn et al. 2004; Dunn et al. 2007). In general, however, this signal transduction is preceded by reduction in the flow of electrons and rate of photosynthesis, which would lead to the accumulation of excess irradiance within the dinoflagellates (Brown 1997).

In line with these hypotheses, a CO₂-induced bleaching mechanism has been proposed which implicates a decline in the photoprotective mechanism, photorespiration, as the initial site of dysfunction (Anthony et al. 2008; Crawley et al. 2010). *Symbiodinium* possess a Form II RuBisCO, which lacks discrimination between CO₂ and O₂ fixation, the latter of which leads to the photorespiratory cycle rather than the Calvin-Benson cycle. Under elevated CO₂ conditions, genetic expression of the first enzyme in the photorespiratory cycle is down regulated, suggesting that the potential to use photorespiration for photoprotection is reduced (Crawley et al. 2010).

Impacts on productivity

An increase in CO₂ may have beneficial implications for photosynthetic organisms but the situation is complicated for dinoflagellates residing within the endoderm of corals. Corals have the ability to transport HCO₃-, a carbonate species that also increases under OA conditions (Goiran et al. 1996). This ability arises from the presence of a carbon concentrating mechanism (CCM) consisting of isoforms of carbonic anhydrase (CA), which catalyze the inter-conversion of HCO₃- and CO₂ (Bertucci et al. 2013). Furthermore, H⁺-ATPase plays a role in CO₂ protonation (Bertucci et al. 2013). The CCM allows CO₂ to readily diffuse into host ectodermal cells, remain within the cytosol as HCO₃- and be transported across the endoderm and symbiosome as CO₂ to the dinoflagellates for photosynthesis (Al-Moghrabi et al. 1996).

The impact of OA on the productivity of reef-building corals has been shown to be either positive (Langdon and Atkinson 2005; Crawley et al. 2010; Suggett et al. 2012), negative (Langdon and Atkinson 2005; Anthony et al. 2008; Iguchi et al. 2012; Kaniewska et al. 2012; Anthony et al. 2013) or have no effect (Leclercq et al. 2002; Reynaud et al. 2003; Schneider and Erez 2006; Marubini et al. 2008; Rodolfo-Metalpa et al. 2010; Takahashi and Kurihara 2013) (Table 1.1). It is difficult to ascertain how much of this variation in results is due to variable species resilience or different experimental procedures. Experiments that measure productivity on the same coral species are rare, the duration of the experiments range from hours to months and the experimental light fields range from 30 – 3800 μmol quanta m⁻² s⁻¹ (Table 1.1). As a general rule, it is apparent that the decline in productivity usually coincides with high temperature or high light, indicating that increased CO₂ may exacerbate these stressful conditions.

Impacts on calcification

The biomineralisation mechanism is enhanced in hermatypic corals due to the phenomenom of Light-Enhanced Calcification (LEC) (Yonge and Nicholls 1931), yet its link to carbonate chemistry renders the process vulnerable under OA conditions. LEC is proposed to occur due to photosynthetic removal of CO₂ by the symbiotic dinoflagellates in accordance with the following stoichiometric equations:

calcification
$$\rightarrow$$
 photosynthesis \rightarrow

$$Ca^{2+} + 2 HCO_3^{-} \rightleftharpoons CaCO_3 + H_2O + CO_2 \rightleftharpoons CaCO_3^{-} + CH_2O + O_2 \qquad (4)$$

$$\leftarrow \text{dissolution} \qquad \leftarrow \text{oxidation}$$

Although LEC has been observed in many coral species, the mechanism remains elusive and probably includes the contribution of several factors in addition to CO₂ removal, including increasing O₂ (Colombo-Pallotta et al. 2010), buffering H⁺ (Bertucci et al. 2013), removing inhibitors (Ferrier-Pagès et al. 2000), providing energy to the host (Colombo-Pallotta et al. 2010) and supplying precursors for organic matrix synthesis (Muscatine and Cernichiari 1969). While there are numerous studies describing the decline of calcification under increased pCO₂ (Kroeker et al. 2010; Erez et al. 2011) the mechanism is more complicated than the direct effect of the seawater at the site of calcification, due to the abovementioned processes in connection to photosynthesis.

Potential for adaptation or acclimatization

In relation to the ability of scleractinian corals to adapt to climate change and OA, the general concensus among scientists is that the rate of environmental change is too rapid for adaptation, due to their slow generational time (Kleypas et al. 1999; Hoegh-Guldberg et al. 2007; Veron 2008). Nevertheless, field acclimatization or phenotypic plasticity is evident in corals as indicated by reciprocal transplant experiments (Coles and Jokiel 1978; Foster 1979) and studying species across environmental gradients (Falkowski and Dubinsky 1981). A further promising aspect of coral evolution, however, is their association with *Symbiodinium*, which have short generation times (Wilkerson et al. 1988) and have shown the ability to adapt to local conditions (Howells et al. 2012). Furthermore, evidence of acclimatization has been shown in adult coral colonies under thermal stress through shuffling the proportion of genetic types of *Symbiodinium* (Jones et al. 2008). In addition, coral juveniles have an opportunity to establish symbiosis with many types of *Symbiodinium* during early ontogeny and this represents an avenue for adaptation to climate change and OA (Abrego et al. 2009; Cumbo et al. 2013).

Table 1.1: Summary of the current literature investigating the impact of OA on the productivity of scleractinian corals. Change in net photosynthesis (P_{net}) and dark respiration (R_{dark}) relative to the control and normalised to branch surface area or equivalent. Control pCO₂ approximately 380 μatm, mid pCO₂ ranges 500 - 800 μatm and high pCO₂ ranges 800 – 2100 μatm.

Reference	Coral	CO ₂ manipulation	Duration	μmol quanta m ⁻² s ⁻¹	P _{net} at mid CO ₂	P _{net} at high CO ₂	R _{dark} at mid CO ₂	R _{dark} at high CO ₂	Comments
Leclercq et al., 2002	Community	Bubbling CO ₂	4 weeks	220	0%	0%	0%	22%	
Reynaud et al., 2003	Stylophora pistillata	Bubbling CO ₂	5 weeks	380	0%	186	0%	8:28	At +3°C Pnet declined 32% at mid CO ₂
Langdon & Atkinson, 2005	Community	Acid addition	1.5 hours	1300-1700	34%	-8%		92	Based on O ₂ method reported
Schneider & Erez, 2006	Acropora eurystoma	Acid addition	3 hours	6m depth, Red Sea	§ 2	0%	9	0%	100
Marubini et al., 2008	Stylophora pistillata	Acid addition	8 days	300	0%	0%	*		DIC enrichment increased Pnet by 70%
	300								At +3°C productivity increased 52% at mid CO ₂
Anthony et al., 2008	Acropora intermedia	Bubbling CO ₂	8 weeks	700-1200	*0%	*-88%	*P _{net} here	is actually	and declined 85% at high CO2
							Daily Productivity		At +3°C productivity declined 86% at mid CO2
Anthony et al., 2008	Porites lobata	Bubbling CO ₂	8 weeks	700-1200	*-11%	*-42%	which incor	porates R _{dark}	and declined 95% at high CO ₂
Crawley et al., 2010	Acropora formosa	Bubbling CO ₂	4 days	110	38%	0%	0%	0%	
Rodolfo Metalpa et al., 2010	Cladocora caespitosa	Bubbling CO ₂	1 month- 1 year	30-60	0%	17 2 1	0%	tute: ises	Temperate coral
Iguchi et al., 2012	Porites australiensis	Bubbling CO ₂	8 weeks	120-140	-	*-12-33%	2	3	*P _{net} here is actually Fv/Fm
Kaniewska et al., 2012	Acropora millepora	Bubbling CO ₂	4 weeks	860	0%	-60%	-25%	-75%	
Suggett et al., 2012	Acropora horrida	Bubbling CO ₂	5-6 weeks	100-400	*60%	(# 1)	0%	·	*P _{net} here is actually P _{gross}
Suggett et al., 2012	Porities cylindirca	Bubbling CO ₂	5-6 weeks	100-400	*20%	8#86	-10%		which incorporates R _{dark}
Anthony et al., 2013	Acropora aspera	Bubbling CO ₂	1 day	1000	-30%	£73	50%	888	
Takahashi & Kurihara, 2013	Acropora digitifera	Bubbling CO ₂	5 weeks	152-3789	0%	0%	0%	0%	Based on DIC method (no O ₂ measurement)

Understanding CO₂-induced coral bleaching

Coral bleaching under elevated CO₂ conditions has recently been observed in *A. millepora* (Kaniewska et al. 2012), *Acropora intermedia* and *Porites lobata* (Anthony et al. 2008). Differentiation of the bleaching as either loss of pigments or loss of symbiont cells could not be ascertained in Anthony et al. (2008) due to loss of the biological samples in a fire. For *A. millepora* the bleaching was due to loss of symbiont cells (Kaniewska et al. 2012). These studies have highlighted the impact of increased CO₂ on *Symbiodinium*, as well as increased H⁺ on the coral. While it is likely that acidosis plays a role in disruption of host cellular mechanisms, the disruption of the symbiosis and cellular processes within *Symbiodinium* are likely to be linked to inhibition of photosynthesis due to loss of photoprotection, as exemplified by down-regulated expression of the photorespiratory enzyme, phosophoglycolate phosphatase (PGPase) (Takahashi et al. 2007; Crawley et al. 2010). An alternative hypothesis states that OA may cause coral bleaching through dinoflagellate cell division upon release from CO₂-limitation and subsequent changes in carbon translocation leading to disruption of the host CCM (Wooldridge 2013).

Photodamage and photoinhibition

Photodamage occurs in the photosynthetic apparatus when it is exposed to excess light, as the electron transport chain cannot process any further excitons. To prevent photodamage, the first line of defense is dynamic photoinhibition through reversible photoprotective processes such as non-photochemical quenching (NPQ) (Brown et al. 1999). Should excitons overwhelm NPQ, this may lead to the formation of ROS such as singlet oxygen ($^{1}O_{2}$), the superoxide anion radical (O_{2}), hydrogen peroxide ($H_{2}O_{2}$) and the hydroxyl radical (OH·) (Niyogi 1999). There are several photoprotective mechanisms that involve electron transport such as PSI cyclic electron transport, the water-water cycle and photorespiration; all of these result in less photodamage. Yet, under dynamic light conditions, the inevitable production of ROS necessitates antioxidant and repair systems in order to scavenge ROS and replace damaged proteins (Nishiyama et al. 2006). Environmental stress may tip the balance between photodamage and repair mechanisms, thereby leading to chronic photoinhibition (Takahashi and Murata 2008). Given that previous OA experiments have only shown

bleaching or productivity loss with concomitant high light or temperature (Table 1), it is likely that high pCO₂ affects photoprotection and repair mechanisms, while the photodamage itself may be caused by exposure to excess light.

Photoprotective processes

Photoprotective mechanisms can function to quench light prior to photosynthesis (e.g. xanthophyll cycle), to cycle electrons during photosynthesis (e.g. PSI cyclic electron transport and the water-water cycle), or to increase substrate availability downstream of the photosynthetic light reactions (e.g. photorespiration) (Demmig-Adams and Adams 1996; Niyogi 1999; Endo and Asada 2006). In a previous OA experiment (Crawley et al. 2010), xanthophyll cycling increased despite sub-saturating light conditions, which suggests that the quenching was required due to a reduction in downstream electron transport and the resultant exposure to excess light. Xanthophyll pigment de-epoxidation is triggered by the change in pH (ΔpH) across the thylakoid membrane (Demmig-Adams and Adams 1996). In addition, NPQ can occur in functionally detached LHCs of PSII due to the pH-sensing protein, PsbS (Holzwarth et al. 2009). The water-water cycle and PSI cyclic electron transport are linked here as they can maintain the ΔpH in order to trigger thermal dissipation of excess light (Schreiber and Neubauer 1990). These photoprotective processes may change in order to assist photoacclimation of *Symbiodinium* to OA conditions.

The ability of dinoflagellates to utilize the photorespiratory pathway is largely dependant on the proportion of CO₂ and O₂ surrounding the enzyme RuBisCO, which catalyses both carboxylation and oxygenation reactions (Figure 1.4). As dinoflagellates possess a Form II Rubsico with a low affinity for CO₂, a CCM is required to increase the proportion of CO₂ surrounding this enzyme (Badger et al. 1998; Leggat et al. 1999). Carboxylation is preferred due to the subsequent formation of sugars through the Calvin-Benson Cycle, yet the evolutionary persistence of the Form II RuBisCO suggests a role for oxygenation (Badger et al. 2000). Indeed, the fact that high temperature reduces the affinity of RuBisCO for CO₂ implies that the photoprotective role of photorespiration can be harnessed during thermal stress (Badger and Collatz 1977; Jordan and Ogren 1984). Increasing pCO₂ due to OA is at

odds with this fact, however, as the oxygenase reaction will be suppressed thus reducing the photoprotective role of photorespiration (Crawley et al. 2010).

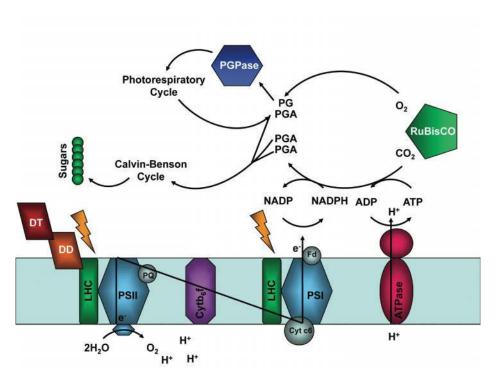


Figure 1.4: Diagram showing the photosynthetic "light reactions" across the thylakoid membrane and the "dark reactions" of carbon fixation by Ribulose bisphosphate carboxylase/oxygenase (RuBisCO). Also depicted are the photoprotective mechanisms photorespiration (beginning with phosphoglycolate phosphatase [PGPase]) and the xanthophyll cycle (de-epoxidation of diadinoxanthin [DD] to diatoxanthin [DT]). Image from Crawley et al. (2010).

Chlorophyll a fluorescence

Chlorophyll a (chl a) fluorescence is a powerful non-invasive tool for investigating photosynthetic performance, which relies on the fact that light energy absorbed by PSII has three possible fates (Bulter 1978). Firstly, photons may drive photochemistry as resonance energy allows electrons to be transferred from the PSII RC chl a, known as P_{680} , through the electron carrier, Phaeophytin (Phe), to the quinone acceptor (Q_A). Secondly, absorbed light energy may be dissipated as heat through non-photochemical quenching (NPQ) or thirdly, as chl a fluorescence (Figure 1.5). Chl a fluorescence measurements are commonly used in coral bleaching research as the relative change in chl a fluorescence can be used to ascertain the proportion of closed

PSII RCs, which may indicate acute or chronic photoinhibition (Krause 1988; Iglesias-Prieto 1995; Warner et al. 1999; Fitt et al. 2001; Ralph et al. 2001).

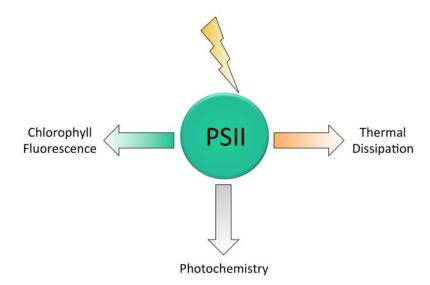


Figure 1.5: The three potential fates of light energy absorbed by Photosystem II (PSII) include: (1) Photochemistry, whereby an electron is transferred from P_{680} to the quinone acceptor (Q_A) ; (2) thermal dissipation as heat through non-photochemical quenching; or (3) chlorophyll fluorescence.

The imaging-Pulse Amplitude Modulated fluorometer (iPAM) (Walz GmbH, Effeltrich, Germany), utilised throughout this thesis, applies saturating pulses of light and captures the resultant fluorescence image. Initially, samples are dark-adapted for one hour to determine the minimum fluorescence (F_0). Following the saturating pulse, the maximum fluorescence (F_m) is measured and the difference, $F_m - F_0$, is the variable fluorescence (F_v) . The ratio F_v/F_m gives an indication of the maximum quantum yield of PSII. In addition, samples can be studied after exposure to the iPAM actinic light, which may close some RCs through NPQ. Here, the steady-state fluorescence (F_t) is higher due to the actinic light. After application of the saturating pulse, the light-adapted maximum fluorescence (F_m) is less than F_m due to NPQ. The light-adapted variable fluorescence (ΔF) is similarly calculated as $F_m' - F_t$ and the ratio, $\Delta F/F_m$ ', is the effective quantum yield of PSII. Acute photoinhibition through a decrease in F_m ' may indicate the use of photoprotective mechanisms, whereas chronic photoinhibition indicates that photodamage may have occurred and results in persistent loss of quantum yield through time (weeks or months). Excitation pressure (Q_m) is a measure of the relative change in the quantum yield of PSII:

$$Q_{\rm m} = 1 - [(\Delta F/F_m') \cdot (F_{\nu}/F_m)^{-1}]$$
 (5)

where $\Delta F/F_m$ ' is measured at midday and F_v/F_m is measured at dusk. An increase in Q_m suggests an increased need for NPQ to reduce over-excitation of PSII, while acclimation to high Q_m can lead to resistance to photoinhibition (Maxwell et al. 1995).

Understanding CO₂-induced productivity loss

Reef-building corals may acquire carbon via autotrophy, utilising the photosynthate of their mutualistic symbiont or via heterotrophic feeding, although the latter is considered to be a compensatory carbon acquisition strategy (Muscatine and Porter 1977; Muscatine et al. 1981). Research suggests that the ability for scleractinian corals to live autotrophically depends on the light regime and therefore the photosynthetic capacity of the symbiont, with heterotrophy being up regulated during episodes of poor water clarity or coral bleaching (Muscatine et al. 1984; Anthony and Fabricius 2000; Grottoli et al. 2006). Despite this ability to compensate, the autotrophy of the coral symbiosis is highly important to the growth, reproduction and survival the holobiont (Hatcher 1988; Hoegh-Guldberg 1999). Respiration is needed for the growth of new biomass and for the maintenance and repair of existing biomass (Amthor 2000). Up to 95% of the photosynthetically fixed carbon can be translocated to the host (Falkowski et al. 1984; Muscatine et al. 1984) while the contribution of dinoflagellate photosynthates may even exceed the host metabolic requirements (Muscatine et al. 1981; Davies 1984). The energy budget is defined by the level that gross photosynthesis (Pg) exceeds respiration (R), with Pg:R > 1 indicating the potential to operate autotrophically and P_g : R < 1 indicating that heterotrophic feeding is required. The P_g:R ratio for the coral symbiosis operates within the range of 0.5 – 5.0 and on average at 2.4 (Hatcher 1988; Battey 1992).

Intuitively, increased pCO₂ should be an advantage for photosynthetic organisms as this could increase the size of the sink for photon energy thereby allowing production of more glucose through the Calvin-Benson cycle. Yet increased pCO₂, with resultant increased HCO₃⁻ and H⁺, will also affect other cellular mechanisms of the coral holobiont and thus complicates the response of their symbiotic dinoflagellates

(Pörtner 2008; Kaniewska et al. 2012). OA has been shown to reduce the host metabolism and change the expression of genes involved in acid-base regulation, ion/molecular transport, apoptosis, calcium homeostasis, cell signaling and cell structure (Kaniewska et al. 2012; Moya et al. 2012). A model was proposed implicating increased ROS as the starting point for the coral bleaching response and this oxidative stress may originate from the *Symbiodinium* cells or host mitochondria, or both (Weis 2008; Kaniewska et al. 2012). Overall, it appears that the paradox of decreased productivity of photosynthetic dinoflagellates under OA conditions (Langdon and Atkinson 2005; Anthony et al. 2008; Iguchi et al. 2012; Kaniewska et al. 2012; Anthony et al. 2013) may linked to the phenomenon of CO₂-induced coral bleaching.

Coral bleaching impacts on productivity

Reducing the population of symbiotic dinoflagellates within the coral host might ultimately lead to a reduction in the photosynthate transferred to the host with implications for the energy budget of the holobiont (Baird and Marshall 2002). Indeed, previous research has indicated that bleached corals consumed their own biomass in order to survive thermal bleaching episodes (Szmant and Gassman 1990). Prolonged coral bleaching conditions have been shown to inhibit gametogenesis (presumably due to the lack of lipids available for eggs) (Szmant and Gassman 1990; Ward et al. 2000), decrease growth rates (Goreau and Macfarlane 1990) and lead to coral mortality (Baird and Marshall 2002). Coral bleaching may be indicated by a decline in coral productivity, which can be investigated with respirometry chambers that measure the rate of photosynthesis oxygen evolution.

Coral respirometry techniques

Coral productivity is estimated from rates of net photosynthesis (P_{net}), dark respiration (R_{dark}) and light enhanced dark respiration (LEDR) measured by the oxygen evolution of a coral sample within respirometry chambers. These chambers sit within a temperature-controlled water bath, contain filtered seawater and a magnetic stirrer for recirculation and are connected to an optical oxygen electrode (Oxy-4, Presens, Germany). The electrodes measure the oxygen flux, which is regressed against time in order to ascertain the steady-state rates of P_{net} , R_{dark} and LEDR. The coral sample is initially dark-acclimated for one hour prior to the respirometry assay, which

commences in darkness to measure R_{dark} . P_{net} is then measured under various light levels to determine both the sub-saturating rate of photosynthesis, α , and the maximum rate of photosynthesis (P_n max). Lastly, the coral samples are returned to darkness to measure LEDR. Data from the respirometry assays can be used to generate Photosynthesis-Irradiance (P:E) curves to estimate productivity or to express the ratio of gross photosynthesis to respiration ($P_g:R$) (Jassby and Platt 1976; Chalker 1981).

Understanding OA impacts on calcification

Calcification occurs in the extracellular calcifying medium (ECM), which exists in close contact with the calicoblastic cells of the coral (Allemand et al. 2004). In accordance with Reaction 4, Ca2+ and HCO3- must be transported to the ECM while CO₂ and H⁺ must be transported away from the ECM, or converted to HCO₃. For calcification to occur, it is imperative that pH remains elevated in the ECM relative to seawater (Al-Horani et al. 2003; Reis 2011; Venn et al. 2011; McCulloch et al. 2012; Venn et al. 2013). In addition, Ca-ATPase actively transports Ca²⁺to the ECM (Al-Horani et al. 2003). As biomineralisation requires both passive and active epithelial transport of those molecules to and from the ECM, any environmental condition that weakens the chemical gradient driving passive transport will lead to an increased requirement for active transport, thereby increasing energy costs of calcification (Allemand et al. 2004). Similarly, any environmental condition that leads to a decrease in productivity will reduce the energy available for active transport and therefore decrease the rate of calcification (Erez et al. 2011). Although the exact mechanism is not known, OA research has indicated that coral calcification rates may decline up to 70% once pCO2 levels double the pre-industrial levels (Renegar and Riegl 2005; Kroeker et al. 2010; Erez et al. 2011). This is a significant issue for coral reefs, as calcification must exceed the rate of bioerosion in order for net accretion to occur (Hoegh-Guldberg et al. 2007).

Metrics for coral calcification

Coral calcification is commonly assessed by measuring the linear growth rate (Lamberts 1978), using the buoyant weight technique (Jokiel et al. 1978) or the

alkalinity anomaly technique (Smith and Key 1975). The linear growth rate is measured by staining the skeleton with alizarin red and requires a long time (> monthly intervals) between measurement points in order to get accurate readings (Allemand et al. 2011). Accordingly, the buoyant weight technique is generally preferred over the linear growth rate as it can be used for short-term experiments (> weekly intervals) and has been shown to have a direct correlation with the skeletal dry weight, although correction may be required for tissue weight (Jokiel et al. 1978; Davies 1989; Langdon et al. 2010). The alkalinity anomaly technique can be used to measure the calcification rate during short-term incubation (< 1 h) due to the fact that A_T decreases two moles for every one mole of CaCO₃ produced, however it is limited to very short-term rates of calcification (Smith and Key 1975; Langdon et al. 2010). In Chapter 3, the skeletal dry weight normalised to the surface area of the branch is used to assess calcification, and this measurement gives a proxy for the skeletal density or branch compactness. This unit of measurement is relevant given that calcification and dissolution processes may change the skeletal density or branch compactness over the long-term and this may not be detectable from short-term measurements of the calcification rate (Kaniewska et al. 2012).

Metrics for community calcification

Community scale measurements of calcification are often employed to project calcification rates under future OA conditions (Silverman et al. 2009; Anthony et al. 2011a; Kleypas et al. 2011; Anthony et al. 2013; Shaw et al. 2013). The change in A_T can be assessed during a slack-water low tide or as the water flows across the reef using Eulerian or Langrangian approaches yet these techniques are restricted to shallow reef flats (Barnes 1983; Langdon et al. 2010). Furthermore, the flow methods require unidirectional flow and do not account for lateral mixing, which limits their application to certain ecosystems, such as reef flats with broad biological zones. In Chapter 3, the community calcification rate could not be assessed as wind direction changed the current seasonally. In addition, lateral mixing would have been significant, as the sites were not characterised by broad biological zones. Accordingly, here the diurnal variation in carbonate chemistry is the best metric to characterise the metabolism of the benthic communities.

Characterising reef acidification

A distinction lies between the process of OA and its effect on coral reef ecosystems as community metabolism in shallow reef waters leads to a diel pattern of reef acidification. Factors that will influence reef acidification include the community assemblage, the water depth, the tidal cycle, the current speed and the current direction (Anthony et al. 2011a; Kleypas et al. 2011; Price et al. 2012; Shaw et al. 2012; Anthony et al. 2013; Shaw et al. 2013).

Carbonate chemistry parameters

All carbonate chemistry parameters can be derived from two of four measurable parameters:

- 1. Total Alkalinity (A_T);
- 2. Dissolved Inorganic Carbon (DIC), also called Total CO₂ (TCO₂);
- 3. pH, and
- 4. Partial pressure of CO₂ (pCO₂) or fugacity of CO₂ (fCO₂).

In addition, the temperature, pressure and salinity must be entered and dissociation constants specified in order for the calculations to run in the program CO2calc (Robbins et al. 2010). This program draws on the dissociation constants for carbonic acid (K₁ and K₂) (Mehrbach et al. 1973; Dickson and Millero 1987), boric acid (K_B) (Dickson 1990a), water (K_W) (Millero 1995), bisulfate ion (K_{SO4}) (Dickson 1990b), hydrogen fluoride (K_F) (Dickson and Riley 1979), phosphoric acid (K_{P1}, K_{P2} and K_{P3}) and silicia acid (K_{Si}) (Millero 1995) as determined by various researchers. The K'_{SD} for calcite and aragonite are from Mucci (1983) and the solubility of CO₂ in seawater (K₀) is from Weiss (1974) and Murray & Riley (1971). The concentrations of boron (Uppstrom 1974), sulfate (Morris and Riley 1966), fluoride (Riley 1965) and calcium (Riley and Tongudai 1967) are assumed to be proportional to salinity. The activity coefficient of the hydrogen ion (f_H) is necessary for converting between pH scales (Takahashi et al. 1982). Values are also needed for the pressure dependence of the dissociation constants (Millero 1995) and solubility products (Ingle 1975; Millero 1979). The virial coefficients of CO₂ and CO₂-air are needed to correct for the nonideal nature of the gas (Weiss 1974) and the vapour pressure of water above seawater is from Weiss & Price (1980). There are 4 pH scales (Total, Free, Seawater and NBS) that have been used in previous research and the NBS scale was used throughout this thesis (Dickson 1984).

Carbonate chemistry techniques

In order to characterise the carbonate chemistry of field sites and experimental treatment aquaria, water samples can be collected and processed for A_T and pH measurements following the standard operating procedures of Dickson et al. (2007). The 500 mL glass sample bottles are rinsed 3 times with sample water before collection and sealed underwater to prevent atmospheric gas exchange. The samples are then fixed with 100 μ L of saturated mercuric chloride, which is 0.02% of the total sample volume, for later processing.

A potentiometric titration procedure can be used whereby the 20 g of sample seawater is placed in open-cell and titrated with hydrochloric acid (HCl). Firstly, the pH is reduced to approximately 3.5 and stirred to allow CO₂ to escape (Dickson et al. 2003). For this small volume titration, the pH data collection is optimal at an incremental addition of 0.015 mL of 0.1 M HCl until pH 3.0 is reached. A_T is calculated from the titrant volume and pH measurements using a Gran approach, which derives the equivalence point from a non-linear least-square fit of the data (Gran 1952). Each daily run on the titrator (T50, Mettler Toledo, Langacher, Switzerland) is preceded by pH sensor (DGi101-SC, Mettler Toledo) calibration using NBS scale buffers (Mettler Toledo) and titrator calibration with Certified Reference Materials (AG Dickson, SIO, Oceanic Carbon Dioxide Quality Control).

Coral Reproduction under OA

Sexual reproduction in reef-building corals is inherently important to coral reef ecosystems by providing scope for evolutionary processes and population growth, yet many aspects of this process are threatened by future OA conditions (Albright 2011; Harrison 2011). The processes of gametogenesis, fertilisation, larval dispersal, settlement, metamorphosis, post-settlement growth, calcification, symbiont uptake and post-settlement survival may all experience potential impacts from increasing pCO₂ (Suwa et al. 2010; Albright 2011; Doropoulos et al. 2012b).

Gametogenesis and fertilisation are critical aspects of sexual reproduction, yet there are limited studies on the impact of OA on these processes. Given that gametogenesis is affected by coral bleaching due to depleted energy reserves (Szmant and Gassman 1990) and OA can increase bleaching susceptibility (Anthony et al. 2008), it is likely that gametogenesis will be indirectly affect by OA. While a previous study did not see an effect of OA on gametogenesis, colonies with partial mortality were excluded from the analysis, which may have confounded the results (Jokiel et al. 2008). Coral that survived as decalcified fleshy polyps were still able to undergo gametogenesis (Fine and Tchernov 2007) whereas in another study, female colonies were impacted more than male colonies, suggesting that the energy budget is a key factor determining the impact of OA on gametogenesis (Holcomb et al. 2010). Acid-base regulation is an another important process that may be affected by OA with implications for sperm motility, egg secretion of a motility-suppressing substance to prevent polyspermy and subsequent embryo development (Morita et al. 2010).

The success of the coral planulae larvae in dispersal, settlement and metamorphosis may be impacted by OA directly through suppression of larvae metabolism and indirectly through decreasing settlement cues (Albright and Langdon 2011; Nakamura et al. 2011). While temporary metabolic suppression may be advantageous during acute stress from hypercapnia, OA is a chronic stressor and subsequent decreases in growth and motility may have implications for dispersal and settlement (Pörtner 2008; Albright 2011). Furthermore, OA has been shown to interfere with the settlement cues from crustose coralline algae (CCA) and microbial biofilms. Firstly, CCA substrate coverage declines in OA conditions and secondly, the normally preferred CCA are avoided under increasing pCO₂ (Kuffner et al. 2008; Doropoulos et al. 2012b). These changes in larval physiology and behavior, in combination with ecological community shifts, will have significant implications for successful coral recruitment.

Many studies have investigated the effect of OA on post-settlement growth, calcification and survival as this early life-history stage may be particularly vulnerable to increasing pCO₂ (Cohen et al. 2009; Anlauf et al. 2011; dePutron et al. 2011; Doropoulos et al. 2012a). Initially, a direct impact on calcification may occur due to the change in carbonate chemistry of the surrounding seawater. Thereafter, for corals that acquire their symbiotic dinoflagellates horizontally (environmental

transmission) rather than vertically (maternal transmission), calcification and growth may be further impacted by symbiont fitness. A previous study has shown that *Symbiodinium* uptake was delayed and polyp growth declined under future OA conditions (Suwa et al. 2010). While symbiont fitness has been shown to increase the thermal tolerance of the holobiont, the differential response of the coral holobiont to the effects of OA in relation to the type of *Symbiodinium* they harbour is largely unknown.

Establishment of symbiosis with Symbiodinium

The eggs and larvae of *A. millepora* are devoid of *Symbiodinium* but uptake occurs from the environment approximately 5-13 days after settlement (Babcock 1985; Babcock and Heyward 1986). *Symbiodinium* are ingested while the coral host is feeding and are subsequently phagocytosed by the endodermal cells (Schwarz et al. 1999). While vertical transmission ensures symbiont infection, horizontal transmission allows for flexible establishment of symbiosis with different types of *Symbiodinium*, which may have important implications to the adaptation of the coral holobiont to changing environmental conditions (Baird et al. 2007; Abrego et al. 2009; Cumbo et al. 2013). In addition, the short life cycle of dinoflagellate cells (1-3 days) allows *Symbiodinium* to locally adapt to environmental conditions, which may explain instances of phylogenetic similarities across ecological zones (Fitt and Trench 1983; Howells et al. 2012; Noonan et al. 2013).

Techniques for assessing Symbiodinium types

There are currently nine broad genetic clades of *Symbiodinium* (A-I) described (Rowan and Powers 1991; Pochon and Gates 2010), which contain genetic and ecologically distinct types, some of which are considered species (LaJeunesse et al. 2004; Sampayo et al. 2009). The clade lineages were originally designated using chloroplast and mitochondrial DNA sequences yet these classifications were not ecologically distinct (Sampayo et al. 2007; Frade et al. 2008). Utilising more rapidly evolving regions of DNA, such as internal transcribed spacer regions of the rDNA (ITS1 and ITS2), has revealed significant phylogenetic *Symbiodinium* types within these clades (Sampayo et al. 2009).

Denaturing gradient gel electrophoresis (DGGE) analysis of the ITS1 rDNA region is a technique found to differentiate high species diversity in comparison to single stranded conformation polymorphism (SSCP) of the ITS1/2 regions or restriction fragment length polymorphism (RFLP) of the large subunit (LSU or 28S) and small subunit (SSU or 18S) of rDNA (Sampayo et al. 2009). Accordingly, DGGE analysis of ITS1 was used in chapter 3 of this thesis. The double stranded DNA is separated by an increasing gradient of denaturing chemicals and migrates across the gel at a different rate with sensitivity for a single-base pair mutation (Muyzer 1999). Bands are subsequently excised from the gels for direct sequencing in order to confirm the *Symbiodinium* type (LaJeunesse et al. 2003). Although quantitative polymerase chain reaction (qPCR) techniques have recently been developed to assess background type populations, the primers developed to date can only differentiate to the clade level (Mieog et al. 2007; Correa et al. 2009; Mieog et al. 2009).

Research Objectives and Thesis Outline

The major objective of this research thesis is to improve our understanding of the physiological response of scleractinian corals to future pCO₂ conditions in an ecological context. In particular, this project investigates the impact of ocean acidification on the productivity and bleaching response of reef-building corals at various life-history stages and across spatial scales. This research will provide insight to policy makers and reef managers in relation to coral reef ecosystem changes that may occur over the coming century.

Chapter 1 provides a review of the literature and techniques relevant to this thesis, including its significance and research objectives.

Chapter 2 presents a working hypothesis towards the mechanism of coral bleaching under ocean acidification. This research investigates the response of *A. aspera* to increasing pCO₂ with concomitant high light conditions and presents a conceptual model linking photorespiration to CO₂-induced bleaching and productivity loss.

Chapter 3 characterises the diurnal variation in carbonate chemistry at sites around Lizard Island and links this to the OA response of *A. millepora* collected from each of these sites. It was hypothesized that greater diurnal variation in carbonate chemistry may improve resilience of scleractinian corals to future OA conditions. This chapter highlights that site-specific physiological trade-offs may influence the response of reef-building corals to future OA scenarios.

Chapter 4 reports a bleaching response in *A. millepora* juveniles under future OA conditions. The effect of OA on coral juveniles is hypothesised to impact *Symbiodinium* uptake and photochemical efficiency. This study links the photophysiology and bleaching response with recruits containing a dominant population of *Symbiodinium* type D1 or D1-4, with potential implications for post-settlement survivorship and population dynamics.

Chapter 5 recapitulates the key findings of this thesis and discusses the results in the light of ecological implications for the Great Barrier Reef. The synopsis outlines the effect of OA on the photo-physiology, productivity, calcification, reproduction and symbiont acquisition of reef-building corals. Future avenues for research are suggested based on new research gaps identified in this thesis with the aim to continue to provide up-to-date scientific information to policy makers and reef managers.

References

- Abrego D, van Oppen MJH, Willis BL (2009) Onset of algal endosymbiont specificity varies among closely related species of *Acropora* corals during early ontogeny. Molecular Ecology 18:3532–3543
- Al-Horani FA, Al-Moghrabi SM, de Beer D (2003) The mechanism of calcification and its relation to phtosynthesis and respiration in the scleractinian coral *Galaxea fascicularis*. Marine Biology 142:419-426
- Al-Moghrabi S, Goiran C, Allemand D, Speziale N, Jaubert J (1996) Inorganic carbon uptake for photosynthesis by the symbiotic coral-dinoflagellate association II. Mechanisms for bicarbonate uptake. Journal of Experimental Marine Biology and Ecology 199:227-248
- Albright R (2011) Reviewing the Effects of Ocean Acidification on Sexual
 Reproduction and Early Life History Stages of Reef-Building Corals. Journal
 of Marine Biology 2011:Article ID 473615
- Albright R, Langdon C (2011) Ocean acidification impacts multiple early life history processes of the Caribbean coral *Porites asteroides*. Global Change Biology 17:2478-2487
- Allemand D, Tambutté ZD, Tambutté S (2011) Coral calcification, cells to reefs. In:

 Dubinsky Z, Stambler N (eds) Coral reefs: an ecosystem in transition.

 Springer Press, New York, pp 119-150
- Allemand D, Ferrier-Pages C, Furla P, Houlbreque F, Puverel S, Reynaud S,

 Tambutte E, Tambutte S, Zoccola D (2004) Biomineralisation in reef-building
 corals: from molecular mechanisms to environmental control. Comptes
 Rendus Palevol 3:453-467
- Amthor JS (2000) The McCree-de Wit-Penning de Vries-Thornley Respiration Paradigms: 30 Years Later. Annals of Botany 86:1-20
- Anlauf H, D'Croz L, O'Dea A (2011) A corrosive concoction: The combined effects of ocean warming and acidification on the early growth of a stony coral are multiplicative. Journal of Experimental Marine Biology and Ecology 397:13-20
- Anthony KRN, Fabricius KE (2000) Shifting roles of heterotrophy and autotrophy in coral energetics under varying turbidity. Journal of Experimental Marine Biology and Ecology 252:221-253

- Anthony KRN, Kleypas JA, Gattuso J-P (2011a) Coral reefs modify their seawater carbon chemistry implications for impacts of ocean acidification. Global Change Biology 17:3655-3666
- Anthony KRN, Kline DI, Diaz-Pulido G, Dove S, Hoegh-Guldberg O (2008) Ocean acidification causes bleaching and productivity loss in coral reef builders.

 Proceedings of the National Academy of Science 105:17442-17446
- Anthony KRN, Diaz-Pulido G, Verlinden N, Tilbrook B, Andersson AJ (2013)

 Benthic buffers and boosters of ocean acidification on coral reefs.

 Biogeosciences 10:4897-4909
- Anthony KRN, Maynard JA, Diaz-Pulido G, Mumby PJ, Marshall PA, Cao L, Hoegh-Guldberg O (2011b) Ocean acidification and warming will lower coral reef resilience. Global Change Biology 17:1798-1808
- Babcock RC (1985) Growth and mortality in juvenile corals (Goniastrea, Platygyra and Acropora): the first year. Proceedings of the Fifth International Coral Reef Congress, Tahiti, French Polynesia, pp 355-360
- Babcock RC, Heyward AJ (1986) Larval development of certain gamete-spawning scleractinian corals. Coral Reefs 5:111-116
- Badger MR, Collatz GJ (1977) Studies on the kinetic mechanism of RuBP carboxylase and oxygenase reactions, with particular reference to the effect of temperature on kinetic papameters. Carnegie Institution of Washington Year Book 76:355-361
- Badger MR, Caemmerer Sv, Ruuska S, Nakano H (2000) Electron flow to oxygen in higher plants and algae: rates and control of direct photoreduction (Mehler reaction) and rubisco oxygenase. Philosophical Transactions of the Royal Society B Biological Sciences 355:1433-1446
- Badger MR, Andrews TJ, Whitney SM, Ludwig M, Yellowlees DC, Leggat W, Price GD (1998) The diversity and coevolution of Rubisco, plastids, pyrenoids, and chloroplast-based CO₂-concentrating mechanisms in algae. Canadian Journal of Botany 76:1052-1071
- Baird AH, Marshall PA (2002) Mortality, growth and reproduction in scleractinian corals following bleaching on the Great Barrier Reef. Marine Ecology Progress Series 237:133-141
- Baird AH, Cumbo VR, Leggat W, Rodriguez-Lanetty M (2007) Fidelity and flexibility in coral symbioses. Marine Ecology Progress Series 347:307-309

- Barnes DJ (1983) Profiling coral reef productivity and calcification using pH and oxygen electrodes. Journal of Experimental Marine Biology and Ecology 66:149-161
- Battey JF (1992) Carbon metabolism in zooxanthellae-coelenterate symbioses. In: Reisser W (ed) Algae and symbioses: plants, animals, fungi, viruses, interactions explored. Biopress Ltd, Bristol, pp174-187
- Bertucci A, Moya A, Tambutte S, Allemand D, Supuran CT, Zoccola D (2013)

 Carbonic anhydrases in anthozoan corals- A review. Bioorganic & Medicinal

 Chemistry 21:1437-1450
- Brown B (1997) Coral bleaching: causes and consequences. Coral Reefs 16:S129-S138
- Brown BE, Ambarsari I, Warner ME, Fitt WK, Dunne RP, Gibb SW, Cummings DG (1999) Diurnal changes in photochemical efficiency and xanthophyll concentrations in shallow water reef corals: evidence for photoinhibition and photoprotection. Coral Reefs 18: 99-105
- Bruno JF, Selig ER (2007) Regional decline of coral cover in the Indo-Pacific: Timing, extent, and subregional comparisons. PLoS ONE 2:e711
- Bryant D, Burke L, McManus J, Spalding M (1998) Reefs at risk: a map-based indicator of threats to the world's coral reefs. World Resources Institute, Washington DC
- Bulter WL (1978) Energy distribution in the photochemical apparatus of photosynthesis. Annual Review of Plant Physiology 29:345-378
- Caldeira K, Wickett ME (2003) Anthropogenic carbon and ocean pH. Nature 425:365
- Caldeira K, Wickett ME (2005) Ocean model predictions of chemistry changes from carbon dioxide emissions to the atmosphere and ocean. Journal of Geophysical Research 110:C09S04
- Canadell JG, Quere CL, Raupacha MR, Fielde CB, Buitenhuisc ET, Ciaisf P,
 Conwayg TJ, Gillettc NP, Houghtonh RA, Marlandi G (2007) Contributions to
 accelerating atmospheric CO₂ growth from economic activity, carbon
 intensity, and efficiency of natural sinks. Proceedings of the National
 Academy of Science 104:18866-18870
- Chalker BE (1981) Simulating light-saturation curves for photosynthesis and calcification by reef-building corals. Marine Biology 63:135-141

- Cohen AL, McCorkle DC, dePutron S, Gaetani GA, Rose KA (2009) Morphological and compositional changes in the skeletons of new coral recruits reared in acidified seawater: Insights into the biomineralization response to ocean acidification. Geochemistry, Geophysics, Geosystems 10:Q07005
- Coles SL, Jokiel PL (1978) Synergistic effects of temperature, salinity and light on the hermatypic coral *Montipora verrucosa*. Marine Biology 49:187-195
- Colombo-Pallotta M, Rodríguez-Román A, Iglesias-Prieto R (2010) Calcification in bleached and unbleached Montastraea faveolata: evaluating the role of oxygen and glycerol. Coral Reefs 29:899-907
- Connell JH (1973) Population ecology of reef building corals. In: Jones OA, Endean RE (eds) Biology and geology of coral reefs. Academic Press, New York, pp 205-245
- Correa AMS, McDonald MD, Baker AC (2009) Development of clade-specific Symbiodinium primers for quantitative PCR (qPCR) and their application to detecting clade D symbionts in Caribbean corals. Marine Biology 156:2403-2411
- Costanza R, d'Arge R, Groot Rd, Farberk S, Grasso M, Hannon B, Limburg K, Naeem S, O'Neill RV, Paruelo J, Raskin RG, Sutton P, Belt Mvd (1997) The value of the world's ecosystem services and natural capital. Nature 387: 253-260
- Crawley A, Kline DI, Dunn S, Anthony K, Dove S (2010) The effect of ocean acidification on symbiont photorespiration and productivity in *Acropora formosa*. Global Change Biology 16:851-863
- Cumbo VR, Baird AH, van Oppen MJH (2013) The promiscuous larvae: flexibility in the establishment of symbiosis in corals. Coral Reefs 32:111-120
- Davies PS (1984) The role of zooxanthellae in the nutritional energy requirements of *Pocillopora eydouxi*. Coral Reefs 2:181-186
- Davies PS (1989) Short-term growth measurements of corals using an accurate buoyant weighing technique. Marine Biology 101:389-395
- Demmig-Adams B, Adams WI (1996) The role of xanthophyll cycle carotenoids in the protection of photosynthesis. Trends in Plant Science 1:21-26
- dePutron SJ, McCorkle DC, Cohen AL, Dillon AB (2011) The impact of seawater saturation state and bicarbonate ion concentration on calcification by new recruits of two Atlantic corals. Coral Reefs 30:321-328

- Dickson AG (1984) pH scales and proton-transfer reactions in saline meadia such as sea water. Geochemica et Cosmochemica Acta 48:2299-2308
- Dickson AG (1990a) Thermodynamics of the dissociation of boric acid in synthetic seawater from 273.15 to 318.15 K. Deep Sea Research Part A. Oceanographic Research Papers 37:755-766
- Dickson AG (1990b) Standard potential of the reaction AgCl(s) + .5H2(g) = Ag(s) + HCl(aq) and the standard acidity constant of the ion HSO4– in synthetic sea water from 273.15 to 318.15 K. The Journal of Chemical Thermodynamics 22:113-127
- Dickson AG, Riley JP (1979) The estimation of acid dissociation constants in seawater media from potentiometric titrations with strong base. I. The ionic product of water- K_w. Marine Chemistry 7:89-99
- Dickson AG, Millero F (1987) A comparison of the equilibrium-constants for the dissociation of carbonic-acid in seawater media. Deep Sea Research Part A 34:1733-1743
- Dickson AG, Afghan JD, Anderson GC (2003) Reference materials for oceanic CO₂ analysis: a method for the certification of total alkalinity. Marine Chemistry 80:185-197
- Dickson AG, Sabine CL, Christian J (eds) (2007) Guide to best practices for ocean CO₂ measurements. PICES Special Publication 3, pp 191
- Doney SC, Fabry VJ, Feely RA, Kleypas JA (2009) Ocean acidification: The other CO₂ problem. Annual Review of Marine Science 1:169-192
- Doropoulos C, Ward S, Marshell A, Diaz-Pulido G, Mumby PJ (2012a) Interactions among chronic and acute impacts on coral recruits: the importance of size-escape thresholds. Ecology 93:2131-2138
- Doropoulos C, Ward S, Diaz-Pulido G, Hoegh-Guldberg O, Mumby PJ (2012b)

 Ocean acidification reduces coral recruitment by disrupting intimate larvalalgal settlement interactions. Ecology Letters 15:338-346
- Dove S (2004) Scleractinian corals with photoprotective host pigments are hypersensitive to thermal bleaching. Marine Ecology-Progress Series 272:99-116
- Driml S (1999) Dollar values and trends of major direct uses of the Great Barrier Reef Marine Park. Great Barrier Reef Marine Park Authority, Townsville

- Dufault AM, Cumbo VR, Fan T-Y, Edmunds PJ (2012) Effects of diurnally oscillating pCO₂ on the calcification and survival of coral recruits. Proc R Soc B-Biol Sci 279:2951-2958
- Dunn SR, Schnitzler CE, Weis VM (2007) Apoptosis and autophagy as mechanisms of dinoflagellate symbiont release during cnidarian bleaching: every which way you lose. Proceedings of the Royal Society of London B 274:3079-3085
- Dunn SR, Thomason JC, Le Tissier MDA, Bythell JC (2004) Heat stress induces different forms of cell death in sea anemones and their endosymbiotic algae depending on temperature and duration. Cell Death and Differentiation 11:1213-1222
- Egleston ES, Sabine CL, Morel FMM (2010) Revelle revisited: Buffer factors that quantify the response of ocean chemistry to changes in DIC and alkalinity. Global Biogeochemical Cycles 24:GB1002
- Endo T, Asada K (2006) Photosystem I and photoprotection: Cyclic electron flow and water-water cycle. In: Demmig-Adams B, Adams W, Mattoo A (eds)

 Photoprotection, Photoinhibition, Gene Regulation and Environment.

 Springer, The Netherlands, pp 205-221
- Erez J, Reynaud S, Silverman J, Schneider K, Allemand D (eds) (2011) Coral calcification under ocean acidification and global change. Springer Press, New York
- Falkowski PG, Dubinsky Z (1981) Light-shade adaptation of *Stylophora pistillata*, a hermatypic coral from the Gulf of Eilat. Nature 289:172-174
- Falkowski PG, Dubinsky Z, Muscatine L, Porter JW (1984) Light and bioenergetics of a symbiotic coral. BioScience 34:705-709
- Feely RA, Sabine CL, Lee K, Berelson W, Kleypas J, Fabry VJ, Millero FJ (2004) Impact of anthropogenic CO₂ on the CaCO₃ system in the oceans. Science 305:362-366
- Ferrier-Pagès C, Gattuso J-G, Dallot S, Jaubert J (2000) Effect of nutrient enrichment on growth and photosynthesis of the zooxanthellate coral. Coral Reefs 19:103-113
- Fine M, Tchernov D (2007) Scleractinian coral species survive and recover from decalcification. Science 315:1811

- Fisher PL, Malme MK, Dove S (2012) The effect of temperature stress on coral— *Symbiodinium* associations containing distinct symbiont types. Coral Reefs 31:473-485
- Fitt W, Brown B, Warner M, Dunne R (2001) Coral bleaching: Interpretation of thermal tolerance limits and thermal thresholds in tropical corals. Coral Reefs 20:51-65
- Fitt WK, Trench RK (1983) The relation of diel patterns of cell division to diel patterns of motility in the symbiotic dinoflagellate *Symbiodinium microadriaticum* Freudenthal in culture. New Phytologist 94:421-432
- Foster AB (1979) Phenotypic plasticity in the reef corals *Montastraea annularia* (Ellis & Solander) and *Siderastrea siderea* (Ellis & Solander). Journal of Experimental Marine Biology and Ecology 39:25-54
- Frade PR, De Jongh F, Vermeulen F, Van Bleijswijk J, Bak RPM (2008) Variation in symbiont distribution between closely related coral species over large depth ranges. Molecular Ecology 17:691-703
- Glynn P (1991) Coral reef bleaching in the 1980s and possible connections with global warming. Trends in Ecology & Evolution 6:175-179
- Goiran C, Al-Moghrabi S, Allemand D, Jaubert J (1996) Inorganic carbon uptake for photosynthesis by the symbiotic coral /dinoflagellate association I.

 Photosynthetic performances of symbionts and dependence on sea water bicarbonate. Journal of Experimental Marine Biology and Ecology 199:207-255
- Goreau TJ, Macfarlane AH (1990) Reduced growth rate of *Montastrea annularis* following the 1987-1988 coral-bleaching event. Coral Reefs 8:211-215
- Gran G (1952) Determination of the equivalence point in potentiometric titrations: Part II. Analyst 77:661-670
- Grottoli AG, Rodrigues LJ, Palardy JE (2006) Heterotrophic plasticity and resilience in bleached corals. Nature 440:1186-1189
- Hall-Spencer JM, Rodolfo-Metalpa R, Martin S, Ransome E, Fine M, Turner SM, Rowley SJ, Tedesco D, Buia M-C (2008) Volcanic carbon dioxide vents show ecosystem effects of ocean acidification. Nature 454:96-99
- Harrison PL (2011) Sexual reproduction of scleractinian corals. In: Dubinsky Z, Stambler N (eds) Coral Reefs: An Ecosystem in Transition. Springer, Dordrecht, pp 59-85

- Hatcher BG (1988) Coral reef primary productivity: A beggar's banquet. Trends In Ecology & Evolution 3:106-111
- Hill R, Larkum AWD, Frankart C, Kuhl M, Ralph PJ (2004) Loss of functional Photosystem II reaction centres in zooxanthellae of corals exposed to bleaching conditions: using fluorescence rise kinetics. Photosynthesis Research 82:59-72
- Hoegh-Guldberg O (1999) Climate change, coral bleaching and the future of the world's coral reefs. Marine and Freshwater Research 50:839-866
- Hoegh-Guldberg O, Mumby PJ, Hooten AJ, Steneck RS, Greenfield P, Gomez E, Harvell CD, Sale PF, Edwards AJ, Caldeira K, Knowlton N, Eakin CM, Iglesias-Prieto R, Muthiga N, Bradbury RH, Dubi A, Hatziolos ME (2007) Coral reefs under rapid climate change and ocean acidification. Science 318:739-742
- Holcomb M, Cohen AL, McCorkle DC (2010) Gender bias in the impact of ocean acidification on corals. Proceedings of the Oceans Sciences Meeting, Portland, Ore, USA
- Holzwarth AR, Miloslavina Y, Nilkens M, Jahns P (2009) Identification of two quenching sites active in the regulation of photosynthetic light-harvesting studied by time-resolved fluorescence. Chemical Physics Letters 483:262-267
- Howells EJ, Beltran VH, Larsen NW, Bay LK, Willis BL, van Oppen MJH (2012)

 Coral thermal tolerance shaped by local adaptation of photosymbionts. Nature

 Climate Change 2:116-120
- Hurd CL, Cornwall CE, Currie K, Hepburn CD, McGraw CM, Hunter KA, Boyd PW (2011) Metabolically induced pH fluctuations by some coastal calcifiers exceed projected 22nd century ocean acidification: a mechanism for differential susceptibility? Global Change Biology 17:3254-3262
- Iglesias-Prieto R (1995) The effects of elevated temperature on the photosynthetic responses of symbiotic dinoflagellates. In: Mathis P (ed) Photosynthesis: from Light to Biosphere. Kluwer Academic Publishers, Dordrecht, The Netherlands, pp 793-796
- Iglesias-Prieto R, Matta JL, Robins WA, Trench RK (1992) Photosynthetic response to elevated temperature in the symbiotic dinoflagellate *Symbiodinium microadriaticum* in culture. Proceedings of the National Academy of Science 89:10302-10305

- Iguchi A, Ozaki S, Nakamura T, Inoue M, Tanaka Y, Suzuki A, Kawahata H, Sakai K (2012) Effects of acidified seawater on coral calcification and symbiotic algae on the massive coral *Porites australiensis*. Marine Environmental Research 73:32-36
- Ingle SE (1975) Solubility of calcite in the ocean. Mar Chem 3:301-319
- Inoue S, Kayanne H, Yamamoto S, Kurihara H (2013) Spatial community shift from hard to soft corals in acidified water. Nature Climate Change DOI: 10.1038/NCLIMATE1855
- IPCC (2007) The fourth assessment report of the Intergovernmental Panel on Climate Change (IPCC). Cambridge University Press, Cambridge, UK
- Jassby AD, Platt T (1976) Mathematical formulation of the relationship between photosynthesis and light for phytoplankton. Limnology & Oceanography 21:540-547
- Jokiel PL (2011) Ocean Acidifcation and control of reef coral calcification by boundary layer limitation of proton flux. Bulletin of Marine Science 87:639–657
- Jokiel PL, Rodgers KS, Kuffner IB, Andersson AJ, Cox EF, Mackenzie FT (2008)

 Ocean acidification and calcifying reef organisms: a mesocosm investigation.

 Coral Reefs 27:473-483
- Jokiel RL, Maragos JE, Franzisket L (1978) Coral growth: buoywant weight technique. In: Stoddart DR, Johannes RE (eds) Monographs on Oceanographic Methodology. UNESCO, Paris, France, pp 529-542
- Jones AM, Berkelmans R (2011) Tradeoffs to thermal acclimation: energetics and reproduction of a reef coral with heat tolerant *Symbiodinium* Type-D. Journal of Marine Biology 2011:185890
- Jones AM, Berkelmans R, van Oppen MJH, Mieog JC, Sinclair W (2008) A community change in the algal endosymbionts of a scleractinian coral following a natural bleaching event: field evidence of acclimatization. Proceedings of the Royal Society of London: B 275:1359-1365
- Jones N (2013) Troubling milestone for CO₂. Nature Geoscience 6:589-589
- Jones RJ, Hoegh-Guldberg O, Larkum AWD, Schreiber U (1998) Temperatureinduced bleaching of corals begins with impairment of the CO₂ fixation mechanism in zooxanthellae. Plant Cell and Environment 21:1219-1230

- Jordan DB, Ogren WL (1984) The CO₂/O₂ specificity of ribulose1,5-bisphosphate carboxylase oxygenase dependence on ribulose bisphosphate concentration, pH and temperature. Planta 161:308-313
- Kaniewska P, Campbell PR, Kline DI, Rodriguez-Lanetty M, Miller DJ, Dove S, Hoegh-Guldberg O (2012) Major cellular and physiological impacts of ocean acidification on a reef building coral. Plos One 7:e34659
- Kleypas J, Buddemeier R, Archer D, Gattuso J, Langdon C, Opdyke B (1999)

 Geochemical consequences of increased atmospheric carbon dioxide on coral reefs. Science 284:118-120
- Kleypas JA, Anthony KRN, Gattuso J-P (2011) Coral reefs modify their seawater carbon chemistry case study from a barrier reef (Moorea, French Polynesia). Global Change Biology 17:3667-3678
- Kleypas JA, Feely RA, Fabry VJ, Langdon C, Sabine CL, Robbins LL (2006) Impacts of Ocean Acidification on Coral Reefs and Other Marine Calcifiers: A Guide for Future Research. Workshop held 18–20 April 2005, sponsored by NSF, NOAA, and the US Geological Survey, St Petersburg, FL, pp 88
- Kline DI, Teneva L, Schneider K, Miard T, Chai A, Marker M, Headley K, Opdyke B, Nash M, Valetich M, Caves JK, Russell BD, Connell SD, Kirkwood BJ, Brewer P, Peltzer E, Silverman J, Caldeira K, Dunbar RB, Koseff JR, Monismith SG, Mitchell BG, Dove S, Hoegh-Guldberg O (2012) A short-term in situ CO₂ enrichment experiment on Heron Island (GBR). Scientific Reports 2:413
- Krause GH (1988) Photoinhibition of photosynthesis. An evaluation of damaging and protective mechanisms. Physiologia Plantarum 74:566-574
- Kroeker KJ, Kordas RL, Crim RN, Singh GG (2010) Meta-analysis reveals negative yet variable effects of ocean acidification on marine organisms. Ecological Letters 13:1419-1434
- Kuffner iB, Andersson AJ, Jokiel PL, Rodgers KS, Mackenzie FT (2008) Decreased abundance of crustose coralline algae due to ocean acidification. Nature Geosciences 1:114-117
- LaJeunesse TC, Loh WKW, Woesik Rv, Hoegh-Guldberg O, Schmidt GW, Fitt WK (2003) Low symbiont diversity in southern Great Barrier Reef corals, relative to those of the Caribbean. Limnology & Oceanography 48:2046-2054

- LaJeunesse TC, Bhagooli R, Hidaka M, DeVantier L, Done T, Schmidt GW, Fitt WK, Hoegh-Guldberg O (2004) Closely related *Symbiodinium* spp. differ in relative dominance in coral reef host communities across environmental, latitudinal and biogeographic gradients. Marine Ecology-Progress Series 284:147-161
- Lamberts AE (1978) Coral growth: alizarin method. In: Stoddart A, Johannes RE (eds) Coral Reefs: research methods. UNESCO, Paris, pp 523 527
- Langdon C, Atkinson M (2005) Effect of elevated *p*CO₂ on photosynthesis and calcification of corals and interactions with seasonal change in temperature/irradiance and nutrient enrichment. Journal of Geophysical Research 110:1-16
- Langdon C, Gattuso J-P, Andersson A (2010) Measurements of calcification and dissolution of benthic organisms and communities. In: U Riebesell, Fabry VJ, Hansson L, Gattuso JP (eds) Guide to best practices for ocean acidification research and data reporting. Publications Offi ce of the European Union, Luxembourg, pp 213-232
- Leclercq N, Gattuso J-P, Jaubert J (2002) Primary production, respiration, and calcification of a coral reef mesocosm under increased CO₂ partial pressure. Limnology and Oceanography 47:558-564
- Leggat W, Badger M, Yellowlees D (1999) Evidence for an inorganic carbonconcentrating mechanism in the symbiotic dinoflagellate *Symbiodinium* sp. Plant Physiology 121:1247-1255
- Lilley RM, Ralph PJ, Larkum AWD (2010) The determination of activity of the enzyme Rubisco in cell extracts of the dinoflagellate alga Symbiodinium sp. by manganese chemiluminescence and its response to short-term thermal stress of the alga. Plant Cell and Environment 33:995-1004
- Little AF, van Oppen MJH, Willis BL (2004) Flexibility in algal endosymbiosies shapes growth in reef corals. Science 304:1492-1494
- Marshall PA, Baird AH (2000) Bleaching of corals on the Great Barrier Reef: differential susceptibilities among taxa. Coral Reefs 19:155-163
- Marubini F, Ferrier-Pages C, Furla P, Allemand D (2008) Coral calcification responds to seawater acidification: a working hypothesis towards a physiological mechanism. Coral Reefs 27:491-499
- Maxwell DP, Falk S, Huner NP (1995) Photosystem II excitation pressure and development of resistance to photoinhibition. Plant Physiology 197:687-694

- McCulloch M, Falter J, Trotter J, Montagna P (2012) Coral resilience to ocean acidification and global warming through pH up-regulation. Nature Climate Change 2:623-627.
- Mehrbach C, Culberso CH, Hawley JE, Pytkowic RM (1973) Measurement of apparent dissociation-constants of carbonic acid in seawater at atmospheric-pressure. Limnology and Oceanography 1973:897-901
- Mieog JC, van Oppen MJH, Cantin NE, Stam WT, Olsen JL (2007) Real-time PCR reveals a high incidence of Symbiodinium clade D at low levels in four scleractinian corals across the Great Barrier Reef: implications for symbiont shuffling. Coral Reefs 26:449-457
- Mieog JC, van Oppen MJH, Berkelmans R, Stam WT, Olsen JL (2009)

 Quantification of algal endosymbionts (*Symbiodinium*) in coral tissue using real-time PCR. Molecular Ecology Resources 9:74-82
- Millero FJ (1979) The thermodynamics of the carbonate system in seawater. Geochemica et Cosmochemica Acta 43:1651-1661
- Millero FJ (1995) Thermodynamics of the carbon dioxide system in the oceans.

 Geochemica et Cosmochemica Acta 59:661-677
- Moberg F, Folke C (1999) Ecological goods and services of coral reef ecosystems. Ecological Economics 29:215-233
- Morita M, Suwa R, Iguchi A, Nakamura M, Shimada K (2010) Ocean acidification reduces sperm flagellar motility in broadcast spawning reef invertebrates.

 Zygote 18:103-107
- Morris AW, Riley JP (1966) The bromide/chlorinity and sulphate/chlorinity ratio in sea water. Deep-Sea Research 13:699-705
- Moya A, Huisman L, Ball EE, Hayward DC, Grasso LC, Chua CM, Woo HN, Gattuso J-P, Foret S, Miller DJ (2012) Whole transcriptome analysis of the coral *Acropora millepora* reveals complex responses to CO₂-driven acidification during the initiation of calcification. Molecular Ecology 21:2440-2454
- Mucci A (1983) The solubility of calcite and aragonite in seawater at various salinities, temperatures, and one atmosphere total pressure. American Journal of Science 283:780-799
- Murray CN, Riley JP (1971) The solubility of gases in distilled water and seawater. IV. Carbon dioxide. Deep-Sea Research 18:533-541

- Muscatine L (1967) Glycerol excretion by symbiotic algae from corals and *Tridacna* and its control by the host. Science 156:516-519
- Muscatine L, Cernichiari E (1969) Assimilation of photosynthetic products of zooxanthellae by a reef coral. Biological Bulletin 137:506-523
- Muscatine L, Porter JW (1977) Reef corals: mutualistic symbioses adapted to nutrient-poor environments? BioScience 27:454-459
- Muscatine L, D'Elia CF (1978) The uptake, retention, and release of ammonium by reef corals. Limnology and Oceanography 23:725-734
- Muscatine L, McCloskey L, Marian RE (1981) Estimating the daily contribution of carbon from zooxanthellae to coral animal respiration. Limnology & Oceanography 26:601-611
- Muscatine L, Falkowski PG, Porter JW, Dubinsky Z (1984) Fate of photosynthetic fixed carbon in light and shade-adapted colonies of the symbiotic coral *Stylophora pistillata*. Proceedings of the Royal Society of London B 222:181-202
- Muyzer GM (1999) DGGE/TGGE: a method for identifying genes from natural ecosystems. Current Opinion in Microbiology 2:317-322
- Nakamura M, Ohki S, Suzuki A, Sakai K (2011) Coral larvae under ocean acidification: survival, metabolism, and metamorphosis. PLoS One 6:e14521
- Nishiyama Y, Allakhverdiev SI, Murata N (2006) A new paradigm for the action of reactive oxygen species in the photoinhibition of photosystem II. Biochimica et Biophysica Acta 1757:742-749
- Niyogi K (1999) Photoprotection revisited: genetic and molecular approaches. Annual review of plant physiology and plant molecular biology 50:333-359
- Noonan SHC, Fabricius KE, Humphrey C (2013) *Symbiodinium* community composition in scleractinian corals Is not affected by life-long exposure to elevated carbon dioxide. PLoS One 8:e63985
- Ohde S, vanWoesik R (1999) Carbon dioxide flux and metabolic processes of a coral reef, Okinawa. Bulletin of Marine Science 65:559-576
- Orr JC, Fabry VJ, Aumont O, Bopp L, Doney SC, Feely RA, Gnanadesikan A, Gruber N, Ishida A, Joos F, Key RM, Lindsay K, Maier-Reimer E, Matear R, Monfray P, Mouchet A, Najjar RG, Plattner G-K, Rodgers KB, Sabine CL, Sarmiento JL, Schlitzer R, Slater RD, Totterdell IJ, Weirig M-F, Yamanaka Y,

- Yool A (2005) Anthropogenic ocean acidification over the twenty-first century and its impact on calcifying organisms. Nature 437:681-686
- Paley AS, Bay LK (2012) Bleaching condition varies among *Acropora millepora* color morphs. Proceedings of the 12th International Coral Reef Symposium, Cairns, QLD, Australia pp 1-5
- Pandolfi JM, Connolly SR, Marshall DJ, Cohen AL (2011) Projecting coral reef futures under global warming and ocean acidification. Science 333:418-422
- Pecheux M (2002) CO₂ increase, a direct cause of coral reef mass bleaching? Marine Life 12:63-68
- Pochon X, Gates RD (2010) A new *Symbiodinium* clade (Dinophyceae) from soritid foraminifera in Hawaii. Molecular Phylogenetics and Evolution 56:492-497
- Pörtner H-O (2008) Ecosystem effects of ocean acidification in times of ocean warming: a physiologist's view. Marine Ecology Progress Series 373:203-217
- Price NN, Martz TR, Brainard RE, Smith JE (2012) Diel variability in seawater pH relates to calcification and benthic community structure on coral reefs. PLoS One 7:e43843
- Ralph PJ, Gademann R, Larkum AWD (2001) Zooxanthellae expelled from bleached corals at 33°C are photosynthetically competent. Marine Ecology Progress Series 220:163-168
- Ries JB (2011) A physicochemical framework for interpreting the biological calcification response to CO₂-induced ocean acidification. Geochimica et Cosmochimica Acta 75: 4053-4064
- Renegar DA, Riegl BM (2005) Effect of nutrient enrichment and elevated CO₂ partial pressure on growth rate of Atlantic scleractinian coral *Acropora cervicornis*.

 Marine Ecology Progress Series 293:69-76
- Reynaud S, Leclercq N, Romaine-Lioud S, Ferrier-Pages C, Jaubert J, Gattuso J (2003) Interacting effects of CO₂ partial pressure and temperature on photosynthesis and calcification in a scleractinian coral. Global Change Biology 9:1660-1668
- Riley JP (1965) The occurance of anomalously high fluoride concentrations in the North Atlantic. Deep-Sea Research 12:219-220
- Riley JP, Tongudai M (1967) The major cation/chlorinity ratios in sea water.

 Chemical Geology 2:263-269

- Robbins LL, Hansen ME, Kleypas JA, Meylan SC (2010) CO2calc—A user-friendly seawater carbon calculator for Windows, Max OS X, and iOS (iPhone). U.S. Geological Survey, Florida
- Rodolfo-Metalpa R, Martin S, Ferrier-Pages C, Gattuso JP (2010) Response of the temperate coral *Cladocora caespitosa* to mid- and long-term exposure to pCO₂ and temperature levels projected for the year 2100 AD. Biogeosciences 7:289-300
- Rohwer F, Seguritan V, Azam F, Knowlton N (2002) Diversity and distribution of coral-associated bacteria. Marine Ecology Progress Series 243:1-10
- Rosenberg E, Koren O, Reshef L, Efrony R, Zilber-Rosenberg I (2007) The role of microorganisms in coral health, disease and evolution. Nature Reviews 5: 355-362
- Rowan R, Powers DA (1991) A molecular genetic classification of zooxanthellae and the evolution of animal-algal symbioses. Science 251:1348-1351
- Sabine CL, Feely RA, Gruber N, Key RM, Lee K, Bullister JL, Wanninkhof R, Wong CS, Wallace DWR, Tilbrook B, Millero FJ, Peng T-H, Kozyr A, Ono T, Rios AF (2004) The oceanic sink for anthropogenic CO₂. Science 305:367-371
- Sampayo EM, Dove S, Lajeunesse TC (2009) Cohesive molecular genetic data delineate species diversity in the dinoflagellate genus *Symbiodinium*.

 Molecular Ecology 18:500-519
- Sampayo EM, Francheschinis L, Hoegh-Guldberg O, Dove S (2007) Niche partitioning of closely related symbiotic dinoflagellates. Molecular Ecology 16:3721-3733
- Schneider K, Erez J (2006) The effect of carbonate chemistry on calcification and photosynthesis in the hermatypic coral *Acropora eurystoma*. Limnology and Oceanography 51:1284-1293
- Schreiber U, Neubauer C (1990) O₂—dependent electron flow, membrane energization and the mechanism of nonphotochemical quenching of chlorophyll fluorescence. Photosynthesis Research 25:279-293
- Schwarz JA, Krupp DA, Weis VM (1999) Late larval development and onset of symbiosis in the scleractinian coral *Fungia scutaria*. Biological Bulletin 196:70-79

- Shamberger KEF, Feely RA, Sabine CL, Atkinson MJ, DeCarlo EH, Mackenzie FT, Drupp PS, Butterfield DA (2011) Calcification and organic production on a Hawaiian coral reef. Marine Chemistry 127:64-75
- Shashar N, Kinane S, Jokiel PL, Patterson MR (1996) Hydromechanical boundary layers over a coral reef. Journal of Experimental Marine Biology and Ecology 199:17-28
- Shaw EC, McNeil BI, Tilbrook B (2012) Impacts of ocean acidification in naturally variable coral reef flat ecosystems. Journal of Geophysical Research 117:C03038
- Shaw EC, McNeil BI, Tilbrook B, Matear R, Bates ML (2013) Anthropogenic changes to seawater buffer capacity combined with natural reef metabolism induce extreme future coral reef CO₂ conditions. Global Change Biology 19:1632-1641
- Shick JM, Lesser MP, Dunlap WC, Stochaj WR, Chalker BE, WuWon. J (1995)

 Depth-dependant responses to solar ultraviolet radiation and oxidative stress in the zooxanthellate coral *Acropora microphthalma*. Marine Biology 122:41-51
- Silverman J, Lazar B, Cao L, Caldeira K, Erez J (2009) Coral reefs may start dissolving when atmospheric CO₂ doubles. Geophysical Research Letters 36:L05606, doi:05610.01029/02008GL036282
- Smith SV (1983) Coral reef calcification. In: Barnes DJ (ed) Perspectives on coral reefs. Brian Clouston, ACT, Australia, pp 240-247
- Smith SV, Key GS (1975) Carbon dioxide and metabolism in marine environments. Limnology and Oceanography 20:493-495
- Suggett DJ, Dong LF, Lawson T, Lawrenz E, Torres L, Smith DJ (2012) Light availability determines susceptibility of reef building corals to ocean acidification. Coral Reefs 32:327-337
- Suwa R, Nakamura M, Morita M, Shimada K, Iguchi A, Sakai K, Suzuki A (2010)

 Effects of acidified seawater on early life stages of scleractinian corals (Genus *Acropora*). Fish Sci 76:93-99
- Szmant AM, Gassman NJ (1990) The effects of prolonged "bleaching" on the tissue biomass and reproduction of the reef coral *Montastrea annularis*. Coral Reefs 8:217-224

- Takahashi A, Kurihara H (2013) Ocean acidification does not affect the physiology of the tropical coral *Acropora digitifera* during a 5-week experiment. Coral Reefs 32:305-314
- Takahashi S, Murata N (2008) How do environmental stresses accelerate photoinhibition? Trends in Plant Science 13:178-182
- Takahashi S, Bauwe H, Badger M (2007) Impairment of the photorespiratory pathway accelerates photoinhibition of photosystem II by suppression of repair but not acceleration of damage processes in *Arabidopsis*. Plant Physiology 144:487-494
- Takahashi T, Williams RT, Bos DL (1982) Carbonate chemistry. In: Broecker WS,
 Spencer DW, Craig H (eds) GEOSECS Pacific Expedition, Volume 3:
 Hydrographic Data 1973-1974. National Science Foundation, Washington DC
 pp 77-106
- UNEP-WCMC (2006) In the front line: Shoreline protection and other ecosystem services from mangroves and coral reefs in the front line. UNEP-WCMC, Cambridge
- Uppstrom LR (1974) The boron/chloronity ratio of deep-sea water from the Pacific Ocean. Deep-Sea Research 21:161-162
- Venn AA, Tambutté E, Holcomb M, Allemand D, Tambutté S (2011) Live tissue imaging shows reef corals elevate pH under their calcifying tissue relative to seawater. PLoS ONE 6: e20013
- Venn AA, Tambutte E, Holcomb M, Laurent J, Allemand D, Tambutte S (2013)

 Impact of seawater acidification on pH at the tissue-skeleton interface and calcification in reef corals. Proceedings of the National Academy of Sciences of the United States of America 110: 1634-1639
- Veron JEN (2000) Corals of the World. Australian Institute of Marine Science, Townsville, QLD
- Veron JEN (2008) Mass extinctions and ocean acidification: biological constraints on geological dilemmas. Coral Reefs 27:459-472
- Ward S, Harrison P, Hoegh-Guldberg O (2000) Coral bleaching reduces reproduction of scleractinian corals and increases susceptibility to future stress. Proceedings of the Ninth International Coral Reef Symposium 2:1123-1128
- Warner ME, Fitt WK, Schmidt GW (1996) The effects of elevated temperature on the photosynthetic efficiency of zooxanthellae in hospite from four different

- species of reef coral: A novel approach. Plant Cell and Environment 19:291-299
- Warner ME, Fitt WK, Schmidt GW (1999) Damage to photosystem II in symbiotic dinoflagellates: A determinant of coral bleaching. Proceedings of the National Academy of Science 96:8007–8012
- Weis VM (2008) Cellular mechanisms of Cnidarian bleaching: stress causes the collapse of symbiosis. Journal of Experimental Biology 211:3059-3066
- Weiss RF (1974) Carbon dioxide in water and seawater: The solubility of a non-ideal gas. Marine Biology 2:203-215
- Weiss RF, Price BA (1980) Nitrous oxide solubility in water and seawater. Marine Chemistry 8:347-359
- Wilkerson FP, Kobayashi D, Muscatine L (1988) Mitotic index and size of symbiotic algae in Caribbean Reef corals. Coral Reefs 7:29-36
- Wooldridge SA (2013) Breakdown of the coral-algae symbiosis: towards formalising a linkage between warm-water bleaching thresholds and the growth rate of the intracellular zooxanthellae. Biogeosciences 10:1647-1658
- Yates KK, Halley RB (2006) Diurnal variation in rates of calcification and carbonate sediment dissolution in Florida Bay. Estuaries and Coasts 29:24-39
- Yonge CM, Nicholls AG (1931) Studies on the physiology of corals. V. On the relationship between corals and zooxanthellae. In: Yonge CM (ed) Scientific Reports Great Barrier Reef Expedition. Trustees of the British Museum, London, UK, pp177-211
- Zeebe RE, Wolf-Gladrow D (2001) CO₂ in seawater. Equilibrium, kinetics, isotopes. Oceanography Series 65, Elsevier, Amsterdam

Chapter 2: A working hypothesis to describe the mechanism of coral bleaching under ocean acidification

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Abstract

Ocean Acidification has been shown to lower the thermal bleaching threshold and reduce productivity of reef building corals, yet the mechanism of this process remains unclear. To clarify this issue and to better understand if a CO₂ bleaching threshold exists, we investigated the photoacclimatory response of Acropora aspera to CO₂ enrichment under high light conditions. Pulse Amplitude Modulation (PAM) fluorometry indicated that there was no treatment effect on the excitation pressure of Photosystem II (PSII), suggesting that PSII is not the initial site of dysfunction. We suggest that a CO₂-induced bleaching mechanism may be initiated by dinoflagellate cell proliferation due to relaxation of carbon limitation, as exhibited by A. aspera exposed to the conservative (IPCC A2) CO₂ scenario (pH 7.8 ± 0.1). In contrast, under the business-as-usual (IPCC A1FI) CO₂ scenario (pH 7.6 ± 0.1), dinoflagellate cell density decreased by 72% and the maximum rate of areal net photosynthesis (P_nmax) declined significantly by 25% in comparison to controls. Overall, the photosynthesis: respiration ratio did not significantly change and remained greater than one for all treatments, suggesting that the host retained autotrophic capacity and that down-regulation of the host carbon concentrating mechanism is not the likely cause of CO₂-induced bleaching. Despite high light conditions, increasing CO₂ did not trigger increased thermal dissipation through the xanthophyll cycle nor increased capacity for thermal dissipation through the xanthophyll pool. The loss of productivity under the A1FI scenario may be partially explained by the significantly increased Chlorophyll a (Chl a) per cell yet decreased Chl a per phaeophytin in the remnant dinoflagellate cells, representing a change in the photosynthetic electron transport pathway. We describe a conceptual model linking the potential decline in photorespiration, a key photoprotective mechanism, with CO₂-induced bleaching. Overall, this study describes fundamental physiological aspects of CO₂-induced bleaching and productivity loss and recommends that management strive to cap anthropogenic CO₂ emissions below IPCC A2 CO₂ levels predicted for the end of the century in order to retain the socio-economic benefits of coral reefs.

Introduction

Atmospheric carbon dioxide (CO₂) levels have recently reached 400 ppm (Jones 2013), compared to 280 ppm before the Industrial Revolution, presenting a significant challenge to Scleractinian corals firstly due to warming associated with radiative forcing in the atmosphere and secondly by changing the carbonate chemistry of the oceans (Caldeira and Wickett 2003; Canadell et al. 2007; Hoegh-Guldberg et al. 2007). The former is implicated in rising sea surface temperatures and the phenomenon of coral bleaching, which is the loss of photosynthetic symbionts (dinoflagellates, genus: Symbiodinium) or their pigments from coral tissue (Brown 1997; Hoegh-Guldberg 1999). The second effect leads to the process of ocean acidification (OA), a decrease in pH and carbonate ions (CO₃⁻) that is detrimental to formation of the coral calcium carbonate (CaCO₃) skeleton (Erez et al. 2011). While little research has looked at the effect of OA on the primary productivity of Symbiodinium, some studies have reported a decline (Reynaud et al. 2003; Anthony et al. 2008) while others have reported no change (Leclercq et al. 2002; Schneider and Erez 2006), which is likely a reflection of diversity amongst species and different experimental manipulations. It has recently been suggested that OA can lower the thermal bleaching threshold of Acropora intermedia and Porites lobata, two ubiquitous reef-building corals on the Great Barrier Reef (Anthony et al. 2008). A further study demonstrated that OA caused down-regulation of the first enzyme involved in photorespiration and increased thermal dissipation, which are photoprotective mechanisms of the symbiont; however, this experiment was performed under low light (Crawley et al. 2010). In the present study, we assess the response of Acropora aspera under high light conditions and future IPCC projections of OA in order to determine the photo-physiological pCO₂ threshold. Furthermore, we explore the potential mechanism of CO₂-induced bleaching using a conceptual model.

The mechanism of thermal coral bleaching is well studied, although debate still exists about whether the site of dysfunction lies within Photosystem II (PSII) (Warner et al. 1999) or downstream at the site of carbon fixation by the enzyme RuBisCO (Ribulose-1,5-bisphosphate carboxylase oxygenase) (Jones et al. 1998; Leggat et al. 2004). Light-induced photodamage is also exacerbated by thermal suppression of the protein synthesis involved in PSII repair (Takahashi et al. 2008), which was shown to

occur when Adenosine Triphosphate (ATP) synthesis and/or RuBisCO were chemically inhibited (Takahashi and Murata 2008). Furthermore, RuBisCO has reduced activity under thermal stress (Lilley et al. 2010). Regardless of the starting point, a blockage in the photosynthetic pathway will ultimately lead to similar symptoms, such as an increase in excitation pressure (Q_m) (Iglesias-Prieto et al. 2004). An increase in light beyond the capacity of PSII or RuBisCO, leads to an increase in reactive oxygen species (ROS) and by-products of photosynthesis. There are several photoprotective cycles to prevent build up of these ROS, such as non-photochemical quenching (NPQ), photorespiration, chlororespiration, Photosystem I (PSI) cyclic electron transport and the water-water cycle (Heber et al. 1996; Niyogi 1999; Badger et al. 2000; Nixon 2000). In addition, photoinhibition itself may be an indication of inadequate photoprotection (Matsubara and Chow 2004; Murchie and Niyogi 2011; Hill et al. 2012). Once the capacity of these photoprotective mechanisms along with the antioxidant and scavenging systems are exceeded, ROS damage occurs to the photosynthetic apparatus and surrounding pigments and lipids (Weis 2008). The ROS may also damage the host coral tissue, thereby triggering a cell signalling cascade leading to programmed cell death or necrosis of host cells containing the symbiont (Dunn et al. 2004) or leading to expulsion of the dinoflagellate from the coral host (Weis 2008). While thermal bleaching may occur due to the concomitant high photon flux, which exceeds the rate of photoprotective, scavenging and repair mechanisms, the process of CO₂-induced bleaching has not been elucidated.

Photoprotective mechanisms can function to quench light prior to primary photochemistry (e.g. xanthophyll cycle) (Demmig-Adams and Adams 1996), to cycle electrons during photosynthesis (e.g. PSI cyclic electron transport and the water-water cycle) (Endo and Asada 2006), or to increase substrate availability downstream of the photosynthetic light reactions (e.g. photorespiration) (Niyogi 1999). The xanthophyll cycle contributes to NPQ that thermally dissipates excess light energy within dinoflagellates through the de-epoxidation of diadinoxanthin (Dn) to diatoxanthin (Dt), and therefore diverts energy from downstream assimilatory carbon fixation. Cyclic electron transport is also a non-assimilatory but functions to produce the pH gradient across the thylakoid membrane, which generates ATP and partly triggers the xanthophyll cycle (Miyake et al. 2005). On the other hand, as photorespiration generates 19% more Adenosine diphosphate (ADP) and 50% more Nicotinamide

adenine dinucleotide (phosphate) (NAD(P)⁺) than carbon fixation, it can relieve blockage downstream from the light reactions of photosynthesis (Badger et al. 2000; Ort and Baker 2002) (Figure 1). This is explained by the ADP and NAD(P)+ outputs of the following equations:

 O_2 fixation by RuBisCO: $12 O_2 + 12 RuBP + 57 ATP + 18 NADH_2 + 18 NADP = 18$ $G3P + 57 ADP + 57 P + 36 NAD(P)^+ + 6 CO_2$

 CO_2 fixation by RuBisCO: 12 CO_2 + 12 RuBP + 48 ATP + 24 NADPH = 24 G3P + 48 ADP + 44 P + 24 NAD(P)⁺

Where RuBP is Ribulose-1,5-bisphosphate, the substrate acted upon by the RuBisCO enzyme and G3P is Glyceraldehyde 3-phosphate, a triose from which all hexose sugars are synthesised, and P is inorganic phosphate. The primary role of photorespiration is to remove the toxic compound glycolate (Figure 1). Additional benefits of photorespiration include removal of O₂, which could otherwise potentially form ROS, and regeneration of CO₂, thereby improving ability for carbon fixation. Photorespiration is dependent on the ratio of CO₂:O₂ at the site of RuBisCO and previous research has shown that high CO₂ conditions in the seawater (1100 ppm CO₂) led to a 50% down-regulation of the first enzyme in the *Symbiodinium* photorespiratory pathway in *Acropora formosa* (Crawley et al. 2010). Here we consider the impact of this significant loss of an alternate electron sink under high light, as opposed to low light conditions.

In addition to their photoprotective mechanisms, dinoflagellates have the ability to photoacclimate in order to reduce photoinhibition and photodamage (Niyogi 1999); through differential survival over generations this could lead to photoadaption (Ralph et al. 2002). Photoacclimation can occur on short time scales (seconds to minutes), for example through a change in enzyme activity (Lilley et al. 2010). Over longer time scales (hours to days), photoacclimation may occur through reduced Chlorophyll *a* (Chl *a*) concentration in the Light Harvesting Complexes (LHCs) to decrease light captured (Ralph et al. 2002) or increased xanthophyll to Chl *a* ratio to increase ability for thermal dissipation (Warner and Berry-Lowe 2006; Kramer et al. 2013) or changes in the photorepair process (Takahashi et al. 2008; Takahashi et al. 2013). In

this experiment, we have captured short-term and long-term photoacclimatory changes and although long-term, trans-generational OA experiments would be required to assess the capacity for photoadaptation, an inability to photoacclimate effectively would not provide a selective advantage.

The energetic costs of photoprotection, photoinhibition, photoacclimation and repair of photodamage represents a significant trade-off to the energy obtained under high light conditions (Hoogenboom et al. 2009). Despite the fact that CO₂ is usually a limiting substrate for photosynthesis, decreased coral productivity has been observed under CO₂ enrichment, a result likely confounded by either increased dinoflagellate cell density leading to increased competition for resources (Reynaud et al. 2003; Wooldridge 2009) or bleaching (Anthony et al. 2008). A decrease in areal productivity under high CO₂ may represent a conundrum for the holobiont as the symbionts incur increased energetic costs of photoacclimation, such as pigment synthesis, and additional energy is required to calcify the coral skeleton under OA conditions (Jokiel 2011).

To form a better understanding of coral acclimation to high CO₂ conditions, we examined branches of the reef flat coral, *Acropora aspera*, during exposure to future predicted levels of pCO₂ enrichment over 10 days. We assessed the photo-physiology using Pulse Amplitude Modulation (PAM) fluorometry throughout the experiment and measured O₂ respirometry upon completion. In order to clarify the role of OA in coral bleaching and productivity loss, we propose a conceptual model linking the potential loss of photorespiration in a high CO₂ world. This model provides mechanistic knowledge to assist resource managers in defining the pCO₂ threshold for future coral survival.

Materials and Methods

Experimental Approach

In October 2008, 36 branches of *Acropora aspera* (blue morph) were collected from Heron Island reef flat $(23^{\circ}26.77^{\circ} \text{ S}, 151^{\circ}54.80^{\circ} \text{ E})$, which is exposed at low tide (maximum irradiance ranging $1400 - 2300 \mu \text{mol}$ quanta m² s⁻¹). The branches were hung by nylon line in shaded outdoor aquaria (maximum irradiance ranging 720 -

1040 µmol quanta m² s⁻¹ and average daily light 300 µmol quanta m² s⁻¹) to acclimate for 5 days. Following acclimation, the branches were divided among 4 replicate tanks for the control and treatments (n = 3). The seawater carbonate chemistry was maintained via a computer-controlled solenoid valve system (Aquatronica-AEB Technologies, Cavriago, Italy), which bubbled CO2 to levels representing Intergovernmental Panel on Climate Change (IPCC) CO₂ scenarios for the year 2100 (IPCC 2007). The A1FI treatment was pH 7.6 ± 0.1 , the A2 treatment was pH $7.8 \pm$ 0.1 and the control measured pH 8.1 ± 0.2 . Variation in the controls was naturally higher (McElhany and Shallin Busch 2013) due to diel fluctuation in carbonate chemistry on Heron Island reef flat (Kline et al. 2012). The pH sensor (InPro4501VP, Mettler Toledo, Langacher, Switzerland), calibrated with NBS scale standard buffers (Mettler Toledo), was used to continuously monitor the pH in the 200 L sumps and routinely verify the pH in treatment aquaria. We measured the Total Alkalinity (A_T) throughout the experiment following the standard operating procedures of Dickson et al (2007) performing the Gran titration (Gran 1952; Dickson et al. 2007) as described in Kline et al (2012) using a titrator (T50, Mettler Toledo, Langacher, Switzerland). In order to characterise the carbonate chemistry of the treatments (Table 1), this data along with the ambient temperature (25°C), pressure (10.16 dbars) and salinity (35 PSU), was input into the CO2calc program (Robbins et al. 2010). For the calculations we specified the total hydrogen ion scale, the carbonic acid dissociation constants of Mehrbach et al (1973) as refit by Dickson & Millero (1987) and the sulfonic acid dissociation constant of Dickson (1990).

Chlorophyll Fluorescence

An imaging-PAM (iPAM, Walz GmbH, Effeltrich, Germany) was used to measure Chl fluorescence. We measured the maximum dark-adapted quantum yield of PSII photochemistry (F_v/F_m) at dusk and the effective quantum yield of PSII photochemistry $(\Delta F/F_m)$ at noon with the actinic light set at 925 µmol quanta m⁻² s⁻¹. Excitation pressure (Q_m) was calculated as $Q_m = 1 - [(\Delta F/F_m) \cdot (F_v/F_m)^{-1}]$ (Iglesias-Prieto et al. 2004), This was conducted on days 1, 5 and 7 of the experiment with the same branch region selected across all time points (Ralph et al. 2005).

Respirometry

After 10 days, we measured the O₂ evolution and consumption within custom-made acrylic respirometry chambers following the method of Crawley et al (2010). The water bath was set to 25°C, the ambient seawater temperature at Heron Island in October 2008. Corals were dark-acclimated for one hour and the chambers were filled with 0.22 µm filtered treatment water prior to the start of the respirometry. To obtain steady-state dark respiration (R_{dark}), respirometry began with 10 minutes in the dark. Then the actinic light from the iPAM followed 10 minutes steps at 20, 55, 110 µmol quanta m^{-2} s⁻¹ to obtain an estimate of alpha (α), the photosynthetic efficiency at subsaturating irradiance. Next, the maximum rate of net photosynthesis (P_nmax) was determined as the maximum rate of O₂ evolution following 10 minutes at 925 µmol quanta m⁻² s⁻¹, and 5 minute steps at 1075 and 1250 µmol quanta m⁻² s⁻¹. Shorter steps at the higher irradiances were sufficient to obtain steady-state rates of O2 evolution. Finally, Light-Enhanced Dark Respiration (LEDR) was determined as the rate of O₂ consumption during 10 minutes in the dark, which began after approximately 1 minute of darkness. Respirometry rates were normalized to the respective surface area of the coral branches.

Data from the respirometry assays were used to generate Photosynthesis-Irradiance (P: E) Curves using the hyperbolic tangent function (Jassby and Platt 1976)

$$Pg = (P_n max + LEDR) \tanh (E/E_k)$$
 (1)

Where Pg is Gross Photosynthesis in μ mol O_2 cm⁻² h⁻¹, E is Irradiance in μ mol quanta m⁻² s⁻¹ and E_k is the minimum saturating irradiance at which P_n max and α intercept (Jassby and Platt 1976; Chalker 1981). We calculated daily Pg by integrating Eq. 1 over the average hourly light curve recorded by the light loggers throughout the experiment.

Daily
$$Pg = \int_{t=0}^{24} (P_n max + LEDR) \tanh [E(t)/E_k] dt$$
 (2)

Furthermore, we determined the ratio of daily Pg to daily respiration (daily Pg:R), based on LEDR occurring during the 14 hours of daylight and R_{dark} occurring for the 10 hours at night.

Cell Counts and Pigment Profile

The coral branches were frozen in liquid nitrogen on the last day of the experiment at midday and were later water-piked using a high pressure airgun with injection of 0.22 µm filtered seawater. The samples were centrifuged (4,500 x g, 5 min) to separate the dinoflagellate pellet from host tissue. One aliquot was made for dinoflagellate cell counts using a 0.100 mm Tiefe Depth Profondeur hemocytometer (0.0025 mm²). A second aliquot was made to quantify the photosynthetic pigments, including Chl a, Dd, Dt and phaeophytin (phe), using High Performance Liquid Chromatography (HPLC) following the methods of Zapata et al (2000) and Dove et al. (2006). Cells and pigments were normalized to the surface area of the branch, which was measured using a double-dipping wax method appropriate for skeletons with deep corallites (Stimson and Kinzie 1991; Veal et al. 2010).

Statistical Analysis

The CO₂ treatment effect was assessed by the PERMANOVA procedure in PRIMER using a mixed model with treatment fixed and tank nested in treatment (Anderson 2001; McArdle and Anderson 2001). The variables were log transformed and Euclidean distance resemblance matrices were created. For the respirometry, pigments and cell count data the Pseudo-F distribution was created by unrestricted permutation of the raw data. For the chlorophyll fluorescence data, we used a repeated measures design with the highest-order interaction term excluded to overcome the issue of temporal replication within cells (Gurevitch and Chester 1986) and the Pseudo-F distribution was created by permutation of residuals under a reduced model. Monte-Carlo P-values were used for all pair-wise comparisons (Anderson and Robinson 2003).

Results

Chlorophyll Fluorescence

Overall Q_m increased with time and there was a significant treatment x time interaction (Figure 2) (See Supplementary Materials S1 for full statistic results). Pairwise comparisons for the interaction at the level of treatment indicated that on day five the A1FI scenario had reduced Q_m compared to the A2 scenario (Pseudo- $F_{2,9} = 3.32$, $p_{perm} < 0.02$, pairwise A1FI < A2, $p_{MC} < 0.03$). However, there were no significant differences in Q_m due to treatment after seven days. Pairwise comparisons for the interaction at the level of time revealed that excitation pressure was highest on day seven for all treatments. For the control and A2 scenario, Q_m was higher on day five compared to day one (Pairwise, Control and A2 day 1 < day 5 < day 7, $p_{MC} < 0.02$, A1FI day 5 < day 7, $p_{MC} < 0.03$).

Respirometry

 P_n max significantly declined by 29% under the A1FI scenario in comparison to the A2 scenario (Pseudo- $F_{2,9} = 4.73$, $p_{perm} < 0.04$, pairwise A1FI < A2, $p_{MC} < 0.04$) (Figure 3A). Although LEDR declined under the A1FI scenario, high variation precluded statistical significance (Pseudo- $F_{2,9} = 2.83$, $p_{perm} > 0.07$) (Figure 3B). Furthermore, there was no significant effect of the CO_2 treatment on R_{dark} (Pseudo- $F_{2,9} = 0.63$, $p_{perm} > 0.5$). E_k averaged 682 ± 138 µmol quanta m^{-2} s⁻¹ and was not significantly different across CO_2 treatments (Pseudo- $F_{2,9} = 1.02$, $p_{perm} > 0.4$). Similarly, CO_2 treatment did not significantly affect α , the efficiency of photosynthesis at sub-saturating light (Pseudo- $F_{2,9} = 0.50$, $p_{perm} > 0.6$) nor the daily P_g :R ratio, which was 1.2 ± 0.1 (Pseudo- $F_{2,9} = 2.99$, $p_{perm} > 0.08$) (Table S1).

Cell Counts and Pigment Profile

Areal dinoflagellate cell density significantly decreased by 72% under the A1FI scenario compared to the control (Pseudo- $F_{2,9} = 14.60$, $p_{perm} < 0.01$, pairwise A1FI < A2 = Control, $p_{MC} < 0.03$). The estimated component of variation analysis attributed 80% of the variation in dinoflagellate cell density to CO_2 treatment (Figure 4). In contrast, Chl a per cell significantly increased by 50% in the remnant dinoflagellate cells under the A1FI scenario compared to the control (Pseudo- $F_{2,9} = 14.98$, $p_{perm} < 0.01$, pairwise A1FI < A2 = Control, $p_{MC} < 0.03$). There were no changes in the

xanthophyll pool, at 0.20 ± 0.01 (Dt + Dd) Chl a⁻¹, nor the xanthophyll cycle, at 0.19 ± 0.01 Dt / (Dt + Dd) (Pseudo-F_{2,9} = 0.60, p_{perm} > 0.55; Pseudo-F_{2,9} = 0.61, p_{perm} > 0.64) (Figure 5). Yet, there was a significant, 23% decline in the xanthophyll pool normalised to *phe* under the A1FI scenario (Pseudo-F_{2,9} = 7.40, p_{perm} < 0.02, pairwise A1FI < A2 = Control, p_{MC} < 0.01). Similarly, Chl *a*: *phe* significantly declined by 22% under the A1FI scenario (Pseudo-F_{2,9} = 7.69, p_{perm} < 0.02, pairwise A1FI < A2 = Control, p_{MC} < 0.01).

Discussion

This study has attempted to elucidate the mechanisms behind coral bleaching and productivity loss due to OA. It appears that CO₂-induced bleaching may initially begin with dinoflagellate proliferation upon relaxation of CO₂-limitation as observed in A. aspera branches exposed to IPCC A2 scenario CO₂ conditions (pH 7.8) (Figure 4). Under the high light conditions of this experiment, symbiont density decreased in A. aspera branches under the IPCC A1FI CO₂ conditions (pH 7.6) and this was most likely due to photoinhibition, as exemplified by high Q_m values (Iglesias-Prieto at al. 2004). It is plausible that the increased symbiont cell density at pH 7.8 increased the susceptibility of the coral to bleaching, damage or stress, which was observed at pH 7.6. Increased symbiont cell density has similarly been shown as a precursor to thermal bleaching (Cunning and Baker 2013). Verification of this mechanism could be observed by following symbiont density and productivity through time during increased pCO₂ conditions. As described in Crawley et al (2010), the loss of photorespiration, a key photoprotective mechanism, may have led to increased photoinhibition and photodamage, thereby triggering the expulsion of Symbiodinium cells. In addition, the photoacclimation process undertaken by the A. aspera branches may explain the loss of productivity under the A1FI scenario. A decrease in Ch1 a: phe, where phe is the first electron acceptor in the reaction centre of PSII, indicates a change in electron transport capacity (Klimov and Krasnovskii 1981; Vredenberg 2011) (Figure 5). Alternatively, as *phe* is a degradation product of Chl a, the increased phe: Chl a may represent loss of functional Chl a (Matile et al. 1999). This study confirms the detrimental effect of CO₂ enrichment under high light conditions, which could impact future ecological services provided by structurally complex, fastgrowing, branching corals such as Acropora aspera.

Under conceptual models of thermal bleaching, ROS production occurs due to damaged PSII or impairment of CO₂ fixation by Rubisco, both of which lead to overreducation in the photosystems of dinoflagellates (Iglesias-Prieto et al. 1992; Jones et al. 1998; Warner et al. 1999; Hill et al. 2004). In support of this, CO₂-induced bleaching has only been observed with concomitant high irradiance (Anthony et al. 2008; Kaniewska et al. 2012) and not under low light conditions (Crawley et al. 2010; Putnam et al. 2013; Wall et al. 2014). Just as environmental stressors can tip the balance between photodamage and repair mechanisms (Takahashi and Murata 2008), a CO₂-induced decline in photorespiration represents a loss of photoprotection and may also suppress repair mechanisms (Takahashi et al. 2007; Crawley et al. 2010). In recent studies, a decrease in PGPase gene expression was only observed under the business-as-usual (IPCC A1FI) CO₂ scenario (Crawley et al. 2010) rather than conservative CO₂ scenarios (Crawley et al. 2010; Ogawa et al. 2013; Putnam et al. 2013) suggesting a CO₂ threshold at approximately 1000 µatm, but only in the absence of other stressors. For example, when considering that excess dinoflagellate cell density may increase the susceptibility of corals to thermal bleaching (Cunning and Baker 2013), the IPCC A2 CO₂ conditions are also detrimental to coral considering the predicted rise in global SSTs. Therefore; we suggest that the ecological impact of increasing pCO₂ will be intricately dependent on concomitant environmental factors.

An alternative hypothesis for CO₂-induced bleaching implicates the carbon concentrating mechanism (CCM) of the coral host as the initial site of dysfunction (Wooldridge 2013), yet the results of this study do not align with this perspective. A generally stable symbiosis is characterized by the uncoupling of dinoflagellate photosynthesis and growth due to nutrient-limitation, which hinders amino acid synthesis needed for cell division (Dubinsky and Berman-Frank 2001). Despite high variability, dinoflagellate cell density did increase under the A2 scenario (pH 7.8) (Figure 4), which suggests that cell division was promoted due to release from CO₂-limitation in nutrient-replete conditions. The assumption of nutrient-replete conditions at Heron Island is valid considering the high nutrient inputs associated with the seabird population and human effluent treatment on this coral cay (Staunton-Smith and Johnson 1995) in addition to the nutrients provided from fish defecation (Meyer

and Shultz 1985). The host-CCM hypothesis then postulates that this cell division leads to less carbon translocated to the host and therefore less energy to maintain the CCM, paradoxically resulting in CO₂-limitation (Wooldridge 2013). At odds with this hypothesis, carbon translocation per symbiont cell has been shown to increase under increased pCO₂ (pH 7.2) (Tremblay et al. 2013). In the present study, we did not observe a decrease in the daily Pg:R ratio, which remained above one for all treatments suggesting that the coral retained the capacity to operate using autotrophically derived carbon, although the values in this study were slightly lower than the average of 2.4 reported among 70 species of corals (Battey 1992). Nevertheless, a Pg:R ratio greater than one is necessary but not sufficient to conclude autotrophy, as the actual carbon translocated must meet the energy requirements of the host (Muscatine et al. 1981), and a significant amount of carbon may be excreted as mucus and not utilised by host metabolism (Wild et al. 2004). Yet, if carbon translocation to the host declined without heterotrophic supplementary feeding, respiration rates would also decline, and this was not observed in our experiment. In addition, Kaniewska et al. (2012) did not observe a change in expression of the carbonic anhydrase (CA) genes of the host CCM in a similar OA experiment conducted under high irradiance. Ogawa et al. (2013) reported decreased CA expression, yet this occurred subsequent to the decrease in dinoflagellate cell density and was therefore not the trigger of coral bleaching, although CA enzyme activity is needed for verification. Overall, it is highly unlikely that CO₂-limitation due to host CCM dysfunction played a role in CO₂-induced bleaching.

A recent study has implicated the role of nutrient imbalance, specifically phosphate starvation, in coral bleaching due to substitution of phospholipids with presumably less stable sulpholipids (Wiedenmann et al. 2012), yet this is unlikely to have played a role in the CO₂-induced bleaching observed here. These authors suggested that, just as uncoupling of photosynthesis from growth occurs due to departure from the C:N aspect of the "Redfield ratio" (Redfield et al. 1963; Dubinsky and Berman-Frank 2001), departure from the N:P aspect of the "Redfield ratio" may alter the composition of lipid membranes thereby disrupting the functionality of the photosynthetic apparatus (Wiedenmann et al. 2012). A comprehensive seasonal water sampling regime at Heron Island reported benthic N:P ratios ranging from approximately 7:1 in summer to 2:1 in winter, therefore indicating adequate

phosphate levels for balanced growth (Staunton-Smith and Johnson 1995). Accordingly, despite proliferation of the dinoflagellate cells under the IPCC A2 scenario, the nutrient levels were adequate to support the building of stable thylakoid membranes and it is unlikely that this phenomenon influenced in the mechanism of CO₂-induced bleaching.

The increase in Q_m throughout the experiment confirms that the treatments were conducted in high light and indicates that the proportion of closed PSII reaction centres was gradually increasing with time (Maxwell et al. 1995; Iglesias-Prieto et al. 2004) (Figure 2). A previous thermal stress experiment has shown that Symbiodinium cell density in $\emph{A. aspera}$ decreased by 50% while \emph{Q}_{m} increased by 50% to approximately 0.8 (Fisher et al. 2012), which is in agreement with the Q_m levels in the present study and signifies potential bleaching conditions. After 5 days, the A1FI CO₂ treatment temporarily alleviated Q_m, but this effect was no longer apparent after 7 days suggesting that high CO₂ levels do not permanently increase the proportion of open PSII reaction centres. Exposure to high excitation pressure can improve resistance to photoinhibition provided that there is capacity to reduce the light harvested or increase thermal dissipation (Maxwell et al. 1995). In the A1FI CO₂ treatment, however, the dinoflagellate cells increased Chl a per cell, presumably due to the significantly reduced cell density and increased capacity for carbon fixation. Despite high light levels, the xanthophyll cycle may not have been running at full capacity, which has been reported as high as 0.8 Dt / (Dt + Dd) in A. aspera under thermal stress (Middlebrook et al. 2008). In the present study, however, increasing CO₂ did not lead to up-regulation in the xanthophyll pool or de-epoxidation, which is in contrast to the increased thermal dissipation observed due to CO₂ enrichment under low light conditions (Crawley et al. 2010). While these studies estimated the daytime xanthophyll cycling conditions, a unique feature of the dinoflagellate Dd-Dt xanthophyll cycle is that it may also function at night. Accordingly, future studies should quantify the nocturnal xanthophyll cycle, which may increase under OA conditions. In the present study, it appears that exposure to high excitation pressure did not allow improved resistance to photoinhibiton.

Photoacclimation can occur via changes in the LHC pigment concentration or alternatively through changes in the number of PSII reaction centres (Falkowski et al.

1981). In the A1FI CO₂ treatment, we observed a decrease in Chl *a: phe*, which implies a decrease in light-harvesting relative to the subsequent transport of excitation energy from P₆₈₀ to Plastoquinone (PQ), assuming the *phe* is located within the PSII reaction centre (Klimov and Krasnovskii 1981; Vredenberg 2011) (Figure 5). In addition, the increased chlorophyll concentration observed in the high CO₂ treatment most likely decreased the efficiency of light capture due to self-shading and increasing the path of excitation energy flow to the PSII reaction centre (Walters 2005) (Figure 4). On the other hand, increased light may have been experienced by the dinoflagellates due to multiple scattering in the coral tissue (Wangpraseurt et al. 2014) and especially by the coral skeleton (Enriquez et al. 2005; Wangpraseurt et al. 2012). Overall, It appears that CO₂ enrichment leads to a cascade of events resulting in suboptimal photoacclimation of the dinoflagellates, which may in turn explain the decline in areal productivity (Figure 3).

This study has described the photo-physiological response of a reef-building coral to enriched CO₂ conditions. In doing so, we have taken significant steps towards unfolding a new mechanism of CO₂-induced bleaching and productivity loss as previously proposed by Anthony et al (2008). It is now clear that increased anthropogenic CO₂ emissions also represent a significant challenge to the photosynthesis of symbiotic dinoflagellates in reef-building corals. While we only observed bleaching in the A1FI CO₂ scenario, even emissions under the A2 CO₂ scenario may increase bleaching susceptibility due to increased dinoflagellate density, as recently shown by Cunning and Baker (2013). The activity and regulation of RuBisCO may also contribute to the response of Symbiodinium to CO2 enrichment and this warrants future investigation. As the present study did not incorporate the thermal effect of increased anthropogenic CO₂ emissions, the results should be interpreted with caution as the additional impact of temperature-induced bleaching may further lower coral resilience. Overall, our findings support the view that CO₂ emissions should be capped to ensure we do not reach the IPCC A2 CO₂ level in order to preserve reef-building corals and their socio-economic benefit for future generations.

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Table 1: Carbonate chemistry parameters calculated from the pH, total alkalinity (A_T) , temperature $(25^{\circ}C)$, salinity (35 PSU) and pressure (10.16 dbars) using the program CO2calc (Robbins et al. 2010).

	\mathbf{A}_{T}	pCO_2	HCO ₃	CO_3^{2-}	CO_2
pH	$(\mu mol kg^{-1})$	(µatm)	(µmol kg ⁻¹)	(µmol kg ⁻¹)	(µmol kg ⁻¹)
8.1 ± 0.2	2296 ± 39	372 ± 190	1704 ± 234	241 ± 80	11 ± 5
7.8 ± 0.1	2295 ± 39	789 ± 208	1962 ± 100	136 ± 28	22 ± 6
7.6 ± 0.1	2295 ± 41	1319 ± 335	2073 ± 84	91 ± 19	37 ± 9

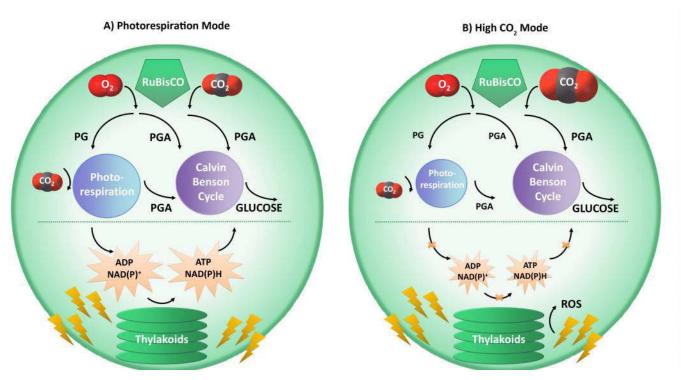


Figure 1: Conceptual model illustrating the photoprotective role of photorespiration. The *Symbiodinium* Form II (Ribulose-1,5-bisphosphate carboxylase oxygenase) does not discriminate between fixation of O₂ or CO₂. O₂ fixation leads to 19% more ADP (Adenosine diphosphate) and 50% more NAD(P)⁺ (Nicotinamide adenine dinucleotide (phosphate)) than through CO₂ fixation and this can relieve the photosynthetic apparatus from additional photon excitation energy. Under high CO₂ conditions, photorespiration is reduced which leads to less availability of ADP and NAD(P)+ to accept the excitation energy flow from the photosystems in the thylakoid membrane. Therefore, excitation energy reacts with oxygen species leading to formation of Reactive Oxygen Species (ROS), which starts the signalling cascade that may lead to coral bleaching. PGA = Phosphoglycerate; PG = Phosphoglycolate.

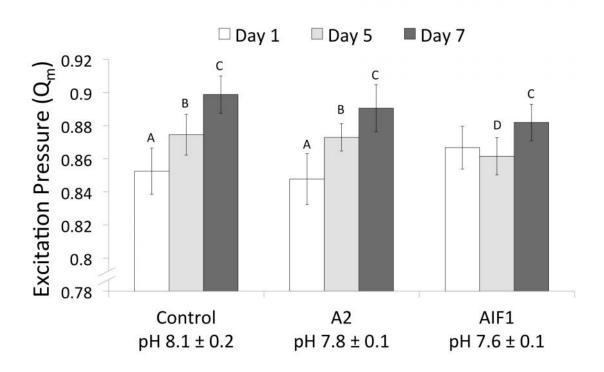


Figure 2: Excitation pressure (Q_m) of photosystem II (PSII) throughout the experiment. On Day 5, A1FI < A2 (p < 0.03). For the control and A2 treatment Day 1 < Day 5 < Day 7 (p < 0.02). For the A1FI treatment Day 5 < Day 7 (p < 0.03). Error bars are standard error (n = 3).

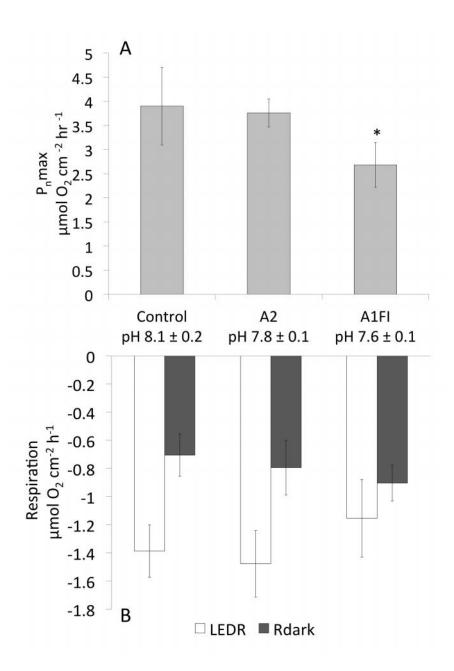


Figure 3: Changes in respirometry under the CO_2 treatments. A) P_n max significantly declines under the IPCC A1FI CO_2 scenario (p < 0.04). B) There were no significant changes to LEDR or Rdark. Error bars are standard error (n = 3).

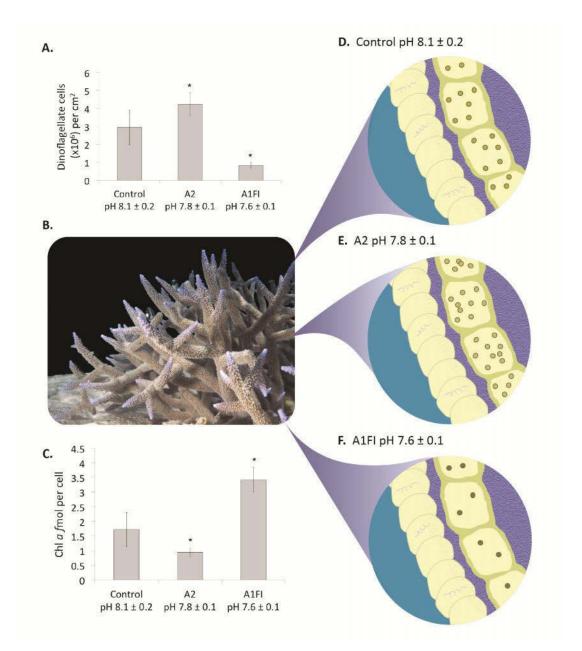


Figure 4: The physiological response of (B) *Acropora aspera* to the CO_2 treatments. (A) Dinoflagellate cell density is significantly decreased under the A1FI IPCC scenario (p < 0.03). (C) Chlorophyll a (chl a) per cell is significantly increased under the A1FI IPCC scenario (p < 0.03). Error bars are standard errors (n = 3). Insets show conceptual coral tissue cross sections of the ectoderm and mesoglea with dinoflagellate cell density under (D) ambient conditions, (E) with increased dinoflagellate cell density under the A2 scenario and (F) after expulsion of dinoflagellates yet increased chl a per cell under the A1FI scenario (Photo: Mark Priest; Drawings not to scale).

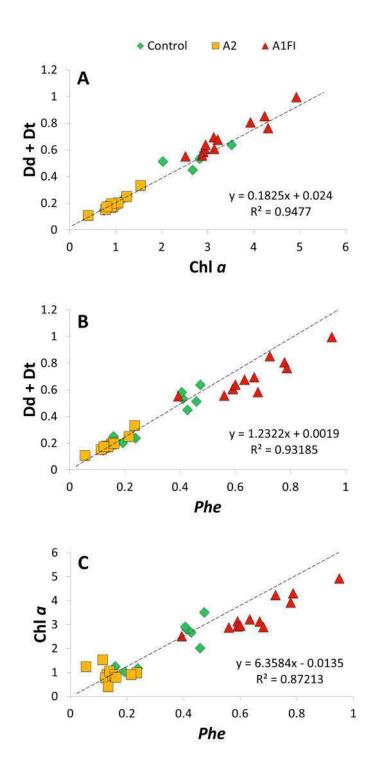


Figure 5: Ratios of photosynthetic pigments in *f*mol per cell. A) The xanthophyll pool, Diadinoxanthin + Diatoxanthin (Dd + Dt), remained in proportion to the quantity of Chlorophyll a (Chl *a*) across all treatments. Yet in the A1FI scenario, both B) the ratio of (Dd +Dt) to pheophytin (*phe*) and C) the ratio of Chl *a* to *phe* significantly declined. Linear regression of control data points overlaid on all graphs.

Table S1: Table of statistical results

Variable	Source of Variation	df	Pseudo F	P(perm)	Pairwise (MC)
R _{dark} (μmol O ₂ cm ⁻² h ⁻¹)	Treatment	2	0.62585	0.5231	
	Tank (Treatment)	9	5.1246	0.0011	
P _n max (μmol O ₂ cm ⁻² h ⁻¹)	Treatment	2	4.7285	0.0352	A2 > A1FI p=0.0303
	Tank (Treatment)	9	1.987	0.0805	III III
LEDR (μmol O ₂ cm ⁻² h ⁻¹)	Treatment	2	2.8303	0.0711	
	Tank (Treatment)	9	1.598	0.1669	
Ek (µmol quanta m-2 s-1)	Treatment	2	1.0199	0.4047	
	Tank (Treatment)	9	3.7384	0.0048	
Alpha (μmol O ₂ cm ⁻² h ⁻¹)	Treatment	2	0.50201	0.6672	
	Tank (Treatment)	9	3.7585	0.0078	
Daily Pg:R	Treatment	2	2.9935	0.0895	
(μmol O ₂ cm ⁻² day ⁻¹)	Tank (Treatment)	9	1.0795	0.408	
Cells per cm²	Treatment	2	14.595	0.0079	Control > A1FI p=0.0266
	Tank (Treatment)	9	5.7983	0.0007	A2 > A1FI p=0.0002
Chl <i>a f</i> mol per cell	Treatment	2	14.979	0.002	Control < A1FI p=0.0243
	Tank (Treatment)	9	4.8168	0.0013	A2 < A1FI p=0.0002
Chl a per cm²	Treatment	2	1.7478	0.2249	
	Tank (Treatment)	9	3.5688	0.0041	
Xanthophyll Cycle	Treatment	2	0.60768	0.6451	
Dt <mark>(Dd + Dt)⁻¹</mark>	Tank (Treatment)	9	1.6808	0.0799	
Xanthophyll Pool per Chl a	Treatment	2	0.60224	0.5566	
(Dd + Dt) Chl a ⁻¹	Tank (Treatment)	9	0.73264	0.6934	
Xanthophyll Pool per Pheophytin	Treatment	2	7.3984	0.0153	A2 > A1FI p=0.0022
(Dd + Dt) Pheophytin ⁻¹	Tank (Treatment)	9	1.7646	0.1219	
Chl a per Pheophytin	Treatment	2	7.6892	0.0155	A2 > A1FI p=0.0011
	Tank (Treatment)	9	2.2337	0.0572	Control > A1FI p=0.0501
Fv/Fm	Treatment	2	0.016125	0.9672	Control Time 1 < 5 < 7 p < 0.004
	Time	2	25.721	0.0001	A2 Time 1< 5< 7 p<0.03
	Tank (Treatment)	9	5.8978	0.0001	A1FI Time no sig diff
	Treatment x Time	4	4.0468	0.0044	
$Q_m = 1 - [(\Delta F/Fm') \cdot (Fv/Fm)^{-1}]$	Treatment	2	0.052335	0.966	At time 5, A1FI < A2 p=0.0248
	Time	2	30.939	0.0001	Control Time 1 < 5 < 7 p<0.008
	Tank (Treatment)	9	6.3031	0.0001	A2 Time 1 < 5 < 7 p<0.02
	Treatment x Time	4	3.3183	0.0133	A1FI Time 5 < 7 p<0.03

References

- Anderson MJ (2001) A new method for non-parametric multivariate analysis of variance. Austral Ecology 26: 32-46
- Anderson MJ, Robinson J (2003) Generalised discriminant analysis based on distances. Australian and New Zealand Journal of Statistics 45: 301-318
- Anthony KRN, Kline DI, Diaz-Pulido G, Dove S, Hoegh-Guldberg O (2008) Ocean acidification causes bleaching and productivity loss in coral reef builders. Proceedings of the National Academy of Science 105: 17442-17446
- Badger MR, Caemmerer Sv, Ruuska S, Nakano H (2000) Electron flow to oxygen in higher plants and algae: rates and control of direct photoreduction (Mehler reaction) and rubisco oxygenase. Philosophical Transactions of the Royal Society B Biological Sciences 355: 1433-1446
- Battey JF (1992) Carbon metabolism in zooxanthellae-coelenterate symbioses. In: Reisser W (ed) Algae and symbioses: plants, animals, fungi, viruses, interactions explored. Biopress Ltd, Bristol, pp 174-187
- Brown B (1997) Coral bleaching: causes and consequences. Coral Reefs 16: S129-S138
- Caldeira K, Wickett ME (2003) Anthropogenic carbon and ocean pH. Nature 425: 365
- Canadell JG, Quere CL, Raupacha MR, Fielde CB, Buitenhuisc ET, Ciaisf P, Conwayg TJ, Gillettc NP, Houghtonh RA, Marlandi G (2007) Contributions to accelerating atmospheric CO₂ growth from economic activity, carbon intensity, and efficiency of natural sinks. Proceedings of the National Academy of Science 104: 18866-18870
- Chalker BE (1981) Simulating light-saturation curves for photosynthesis and calcification by reef-building corals. Marine Biology 63: 135-141
- Crawley A, Kline DI, Dunn S, Anthony K, Dove S (2010) The effect of ocean acidification on symbiont photorespiration and productivity in *Acropora formosa*. Global Change Biology 16: 851-863
- Cunning R, Baker AC (2013) Excess algal symbionts increase the susceptibility of reef corals to bleaching. Nature Climate Change 3: 259-262
- Demmig-Adams B, Adams WI (1996) The role of xanthophyll cycle carotenoids in the protection of photosynthesis. Trends in Plant Science 1: 21-26

- Dickson AG (1990) Standard potential of the reaction AgCl(s) + .5H2(g) = Ag(s) + HCl(aq) and the standard acidity constant of the ion HSO4– in synthetic sea water from 273.15 to 318.15 K. The Journal of Chemical Thermodynamics 22: 113-127
- Dickson AG, Millero F (1987) A comparison of the equilibrium-constants for the dissociation of carbonic-acid in seawater media Deep-Sea Research Part A Oceanographic Research Papers 34: 1733-1743
- Dickson AG, Sabine CL, Christian J (2007) Guide to best practices for ocean CO₂ measurements. PICES Special Publication 3, pp 191
- Dove S, Ortiz J, Fine M, Fisher P, Iglesias-Prieto R, Thornhill D, Hoegh-Guldberg O (2006) Response of holosymbiont pigments from the scleractinian coral *Montipora monasteriata* to short-term heat stress. Limnology & Oceanography 51: 1149-1158
- Dubinsky Z, Berman-Frank I (2001) Uncoupling primary production fromo population growth in photosynthesizing organisms in aquatic ecosystems.

 Aquatic Sciences 63: 4-17
- Dunn SR, Thomason JC, Le Tissier MDA, Bythell JC (2004) Heat stress induces different forms of cell death in sea anemones and their endosymbiotic algae depending on temperature and duration. Cell Death and Differentiation 11: 1213-1222
- Endo T, Asada K (2006) Photosystem I and photoprotection: Cyclic electron flow and water-water cycle. In: Demmig-Adams B, Adams W, Mattoo A (eds) Photoprotection, Photoinhibition, Gene Regulation and Environment. Springer, The Netherlands, pp 205-221
- Enriquez S, Mendez ER, Iglesias-Prieto R (2005) Multiple scattering on coral skeletons enhances light absorption by symbiotic algae. Limnology and Oceanography 50: 1025-1032
- Erez J, Reynaud S, Silverman J, Schneider K, Allemand D (2011) Coral calcification under ocean acidification and global change. In: Dubinsky Z, Stambler N (eds) Coral reefs: an ecosystem in transition. Springer Press, New York, pp 151-176
- Falkowski PG, Owens TG, Ley AC, Mauzerall DC (1981) Effects of growth irradiance levels on the ratio of reaction centres in two species of marine phytoplankton. Plant Physiology 68: 969-973

- Fisher PL, Malme MK, Dove S (2012) The effect of temperature stress on coral— *Symbiodinium* associations containing distinct symbiont types. Coral Reefs 31: 473-485
- Gran G (1952) Determination of the equivalence point in potentiometric titrations: Part II. Analyst 77: 661-670
- Gurevitch J, Chester ST (1986) Analysis of repeated measures experiments. Ecology 67: 251-255
- Heber U, Bligny R, Streb P, Douce R (1996) Photorespiration is essential for the protection of the photosynthetic apparatus of C3 plants against photoinactivation under sunlight. Botanica acta 109: 307-315
- Hill R, Larkum AWD, Frankart C, Kuhl M, Ralph PJ (2004) Loss of functional Photosystem II reaction centres in zooxanthellae of corals exposed to bleaching conditions: using fluorescence rise kinetics. Photosynthesis Research 82: 59-72
- Hill R, Larkum AWD, Prasil O, Kramer DM, Szabo M, Kumar V, Ralph PJ (2012) Light-induced dissociation of antenna complexes in the symbionts of scleractinian corals correlates with sensitivity to coral bleaching. Coral Reefs 31: 963-975
- Hoegh-Guldberg O (1999) Climate change, coral bleaching and the future of the world's coral reefs. Marine and Freshwater Research 50: 839-866
- Hoegh-Guldberg O, Mumby PJ, Hooten AJ, Steneck RS, Greenfield P, Gomez E, Harvell CD, Sale PF, Edwards AJ, Caldeira K, Knowlton N, Eakin CM, Iglesias-Prieto R, Muthiga N, Bradbury RH, Dubi A, Hatziolos ME (2007) Coral reefs under rapid climate change and ocean acidification. Science 318: 739-742
- Hoogenboom MO, Connolly SR, Anthony KRN (2009) Effects of photoacclimation on the light niche of corals: a process-based approach. Marine Biology 156: 2493-2503
- Iglesias-Prieto R, Beltran V, LaJeunesse T, Reyes-Bonilla H, Thome P (2004)

 Different algal symbionts explain the vertical distribution of dominant reef corals in the eastern Pacific. Proceedings of the Royal Society of London. B 271: 1757-1763
- Iglesias-Prieto R, Matta JL, Robins WA, Trench RK (1992) Photosynthetic response to elevated temperature in the symbiotic dinoflagellate *Symbiodinium*

- *microadriaticum* in culture. Proceedings of the National Academy of Science 89: 10302-10305
- IPCC (2007) The fourth assessment report of the Intergovernmental Panel on Climate Change (IPCC). Cambridge University Press, Cambridge, UK
- Jassby AD, Platt T (1976) Mathematical formulation of the relationship between photosynthesis and light for phytoplankton. Limnology & Oceanography 21: 540-547
- Jokiel PL (2011) Ocean Acidifcation and control of reef coral calcification by boundary layer limitation of proton flux. Bulletin of Marine Science 87: 639–657
- Jones N (2013) Troubling milestone for CO₂. Nature Geoscience 6: 589-589
- Jones RJ, Hoegh-Guldberg O, Larkum AWD, Schreiber U (1998) Temperature-induced bleaching of corals begins with impairment of the CO₂ fixation mechanism in zooxanthellae. Plant Cell and Environment 21: 1219-1230
- Kaniewska P, Campbell PR, Kline DI, Rodriguez-Lanetty M, Miller DJ, Dove S, Hoegh-Guldberg O (2012) Major cellular and physiological impacts of ocean acidification on a reef building coral. Plos One 7: e34659
- Klimov VV, Krasnovskii AA (1981) Participation of pheophytin in the primary processes of electron transfer at the reaction centers of photosystem II. Biophysics 27: 186-198
- Kline DI, Teneva L, Schneider K, Miard T, Chai A, Marker M, Headley K, Opdyke B, Nash M, Valetich M, Caves JK, Russell BD, Connell SD, Kirkwood BJ, Brewer P, Peltzer E, Silverman J, Caldeira K, Dunbar RB, Koseff JR, Monismith SG, Mitchell BG, Dove S, Hoegh-Guldberg O (2012) A short-term in situ CO₂ enrichment experiment on Heron Island (GBR). Scientific Reports 2: 413
- Kramer WE, Schrameyer V, Hill R, Ralph PJ, Bischof K (2013) PSII activity and pigment dynamics of *Symbiodinium* in two Indo-Pacific corals exposed to short-term high-light stress. Marine Biology 160: 563-577
- Leclercq N, Gattuso J-P, Jaubert J (2002) Primary production, respiration, and calcification of a coral reef mesocosm under increased CO₂ partial pressure. Limnology and Oceanography 47: 558-564
- Leggat W, Whitney S, Yellowlees D (2004) Is coral bleaching due to the instability of the zooxanthellae dark reactions? Symbiosis 37: 137-153

- Lilley RM, Ralph PJ, Larkum AWD (2010) The determination of activity of the enzyme Rubisco in cell extracts of the dinoflagellate alga Symbiodinium sp. by manganese chemiluminescence and its response to short-term thermal stress of the alga. Plant Cell and Environment 33: 995-1004
- Matile P, Hortensteiner S, Howard T (1999) Chlorophyll degradation. Annual Review of Plant Physiology and Plant Molecular Biology 50: 67-95
- Matsubara S, Chow WS (2004) Populations of photoinactivated photosystem II reaction centers characterized by chlorophyll a fluorescence lifetime in vivo. Proceedings of the National Academy of Science 101: 18234-18239
- Maxwell DP, Falk S, Huner NP (1995) Photosystem II excitation pressure and development of resistance to photoinhibition. Plant Physiology 197: 687-694
- McArdle BH, Anderson MJ (2001) Fitting multivariate models to community data: a comment on distance-based redundancy analysis. Ecology 82: 290-297
- McElhany P, Shallin Busch D (2013) Appropriate pCO₂ treatments in ocean acidification experiments. Marine Biology 160: 1807-1812
- Mehrbach C, Culberso CH, Hawley JE, Pytkowic RM (1973) Measurement of apparent dissociation-constants of carbonic acid in seawater at atmospheric-pressure. Limnology and Oceanography 1973: 897-901
- Meyer JL, Shultz ET (1985) Migrating haemulid fishes as a source of nutrients and organic matter on coral reefs. Limnology & Oceanography 30: 146-156
- Middlebrook R, Hoegh-Guldberg O, Leggat W (2008) The effect of thermal history on the susceptibility of reef-building corals to thermal stress. The Journal of Experimental Biology 211: 1050-1056
- Miyake C, Miyata M, Shinzaki Y, Tomizawa K (2005) CO₂ response of cyclic electron flow around PSI (CEF-PSI) in Tobacco leaves- Relative electron fluxes through PSI and PSII determine the magnitude of non-photochemical quenching (NPQ) of Chl fluorescence. Plant & Cell Physiology 46: 629-637
- Murchie EH, Niyogi KK (2011) Manipulation of photoprotection to improve plant photosynthesis. Plant Physiology 155: 86-92
- Muscatine L, McCloskey L, Marian RE (1981) Estimating the daily contribution of carbon from zooxanthellae to coral animal respiration. Limnology & Oceanography 26: 601-611
- Nixon PJ (2000) Chlororespiration. Philosophical Transactions of the Royal Society London B 355: 1541-1547

- Niyogi K (1999) Photoprotection revisited: genetic and molecular approaches. Annual review of plant physiology and plant molecular biology 50: 333-359
- Ogawa D, Bobeszko T, Ainsworth T, Leggat W (2013) The combined effects of temperature and CO₂ lead to altered gene expression in *Acropora aspera*. Coral Reefs: 1-13
- Ort DR, Baker NR (2002) A photoprotective role for O₂ as an alternative electron sink in photosynthesis? Current Opinion in Plant Biology 5: 193-198
- Putnam HM, Mayfield AB, Fan TY, Chen CS, Gates RD (2013) The physiological and molecular responses of larvae from the reef-building coral *Pocillopora* damicornis exposed to near-future increases in temperature and pCO₂. Marine Biology 160: 2157-2173
- Ralph PJ, Gademann R, Larkum AWD, Kuhl M (2002) Spatial heterogeneity in active chlorophyll fluorescence and PSII activity of coral tissues. Marine Biology 141: 639-646
- Ralph PJ, Schreiber U, Gademann R, Kuhl M, Larkum AWD (2005) Coral photobiology studied with a new imaging pulse amplitude modulated fluorometer. Journal of Phycology 41: 335-342
- Redfield AC, Ketchum BH, Richards FA (1963) The influence of organisms on the composition of sea-water. In: Hill MN (ed) The Sea. Interscience Publishers, John Wiley and Sons, New York, pp 26-77
- Reynaud S, Leclercq N, Romaine-Lioud S, Ferrier-Pages C, Jaubert J, Gattuso J (2003) Interacting effects of CO₂ partial pressure and temperature on photosynthesis and calcification in a scleractinian coral. Global Change Biology 9: 1660-1668
- Robbins LL, Hansen ME, Kleypas JA, Meylan SC (2010) CO2calc—A user-friendly seawater carbon calculator for Windows, Max OS X, and iOS (iPhone). U.S. Geological Survey, Florida
- Schneider K, Erez J (2006) The effect of carbonate chemistry on calcification and photosynthesis in the hermatypic coral *Acropora eurystoma*. Limnology and Oceanography 51: 1284-1293
- Staunton-Smith J, Johnson CR (1995) Nutrient inputs from seabirds and humans on a populated coral cay. Marine Ecology Progress Series 124: 189-200
- Stimson J, Kinzie RA (1991) The temporal pattern and rate of release of zooxanthellae from the reef coral *Pocillopora damicornis* (Linnaeus) under

- Nitrogen-enrichment and control conditions. Journal of Experimental Marine Biology and Ecology 153: 63-74
- Takahashi S, Bauwe H, Badger M (2007) Impairment of the photorespiratory pathway accelerates photoinhibition of photosystem II by suppression of repair but not acceleration of damage processes in *Arabidopsis*. Plant Physiology 144: 487-494
- Takahashi S, Murata N (2008) How do environmental stresses accelerate photoinhibition? Trends in Plant Science 13: 178-182
- Takahashi S, Whitney S, Itoh S, Maruyama T, Badger M (2008) Heat stress causes inhibition of the de novo synthesis of antenna proteins and photobleaching in cultured *Symbiodinium*. Proceedings of the National Academy of Sciences of the United States of America 105: 4203-4208
- Takahashi S, Yoshioka-Nishimura M, Nanba D, Badger MR (2013) Thermal ccclimation of the symbiotic alga *Symbiodinium* spp. alleviates photobleaching under heat stress. Plant Physiology 161: 477-485
- Tremblay P, Fine M, Maguer JF, Grover R, Ferrier-Pages C (2013) Photosynthate translocation increases in response to low seawater pH in a coral-dinoflagellate symbiosis. Biogeosciences 10: 3997-4007
- Veal CJ, Holmes G, Nunez M, Hoegh-Guldberg O, Osborn J (2010) A comparative study of methods for surface area and three dimensional shape measurement of coral skeletons. Limnology and Oceanography: Methods 8: 241-253
- Vredenberg W (2011) Kinetic analyses and mathematical modeling of primary photochemical and photoelectrochemical processes in plant photosystems. Biosystems 103: 138-151
- Wall CB, Fan TY, Edmunds PJ (2014) Ocean acidification has no effect on thermal bleaching in the coral *Seriatopora caliendrum*. Coral Reefs 33: 119-130
- Walters RG (2005) Towards an understanding of photosynthetic acclimation. Journal of Experimental Botany 56: 435-447
- Wangpraseurt D, Larkum AWD, Franklin J, Szabo M, Ralph PJ, Kuhl M (2014)

 Lateral light transfer ensures efficient resource distribution in symbiontbearing corals. Journal of Experimental Biology 217: 489-498
- Wangpraseurt D, Larkum AWD, Ralph PJ, Kühl M (2012) Light gradients and optical microniches in coral tissues. Frontiers in Microbiology 3: 316

- Warner ME, Berry-Lowe S (2006) Differential xanthophyll cycling and photochemical activity in symbiotic dinoflagellates in multiple locations of three species of Caribbean coral. Journal of Experimental Marine Biology and Ecology 339: 86-95
- Warner ME, Fitt WK, Schmidt GW (1999) Damage to photosystem II in symbiotic dinoflagellates: A determinant of coral bleaching. Proceedings of the National Academy of Science 96: 8007–8012
- Weis VM (2008) Cellular mechanisms of Cnidarian bleaching: stress causes the collapse of symbiosis. Journal of Experimental Biology 211: 3059-3066
- Wiedenmann J, D'Angelo C, Smith EG, Hunt AN, Legiret F-E, Postle AD, Achterberg EP (2012) Nutrient enrichment can increase the susceptibility of reef corals to bleaching. Nature Climate Change 3: 160-164
- Wild C, Huettel M, Klueter A, Kremb SG, Rasheed MYM, Jorgensen BB (2004)

 Coral mucus functions as an energy carrier and particle trap in the reef ecosystem. Nature 428: 66-70
- Wooldridge SA (2009) A new conceptual model for the warm-water breakdown of the coral–algae endosymbiosis. Marine and Freshwater Research 60: 483-496
- Wooldridge SA (2013) Breakdown of the coral-algae symbiosis: towards formalising a linkage between warm-water bleaching thresholds and the growth rate of the intracellular zooxanthellae. Biogeosciences 10: 1647-1658
- Zapata M, Rodriguez F, Garrido JL (2000) Separation of chlorophylls and carotenoids from marine phytoplankton: a new HPLC method using a reversed phase C8 column and pyridine containing mobile phases. Marine Ecology Progress Series 195: 29-45

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Chapter 3: Diurnal carbonate chemistry variability impacts the sensitivity of *Acropora millepora* to ocean acidification

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Ocean acidification, community assemblage, carbonate chemistry, photosynthesis, respiration, calcification

Abstract

The predicted decline in oceanic pH facilitated by increasing anthropogenic CO₂ emissions is expected to challenge the calcification and primary productivity of reefbuilding corals. Previous research has either focused on the physiological or ecosystem responses with little understanding of the combined impacts. This study aimed to characterise intra-specific variation in the response of *Acropora millepora* to constant ocean acidification (OA) treatments in the context of their exposure to diurnal variation in carbonate chemistry at distinct sites around Lizard Island. Corals collected from the lagoon sites were exposed to more extreme daily changes in carbonate chemistry than corals collected from outside the lagoon. After subsequent exposure to constant CO₂ levels that correspond to IPCC scenario A1FI (pH 7.7 ± 0.05; pCO₂ 1020 ± 140 μatm) in an aquaria-based OA experimental system, corals from outside the lagoon increased net photosynthesis per cm² (P_{net}), yet there were no significant improvements to the gross photosynthesis to respiration ratio (Pg:R). Furthermore, a site-specific physiological trade-off was exhibited in the form of differential investment in skeletal morphology and tissue thickness with corals from a lagoon-based site having a lower skeletal dry-weight to surface-area ratio, but more protein per cm². Overall, these results indicate that corals collected from reefs with varying diurnal fluxes in carbonate chemistry may respond differently to future OA scenarios, suggesting that the biogeochemical cycles of the collection site must be considered in future experimental studies with implications for the predictive power of climate models when applied to coral reefs.

Introduction

The calcification and primary productivity of marine organisms will be affected by changing marine carbonate chemistry, which is occurring due to the increased oceanic absorption of anthropogenic carbon dioxide (CO₂) emissions, leading to the decline in oceanic pH known as ocean acidification (OA) (Caldeira and Wickett 2003; Sabine et al. 2004; Canadell et al. 2007). Since the industrial revolution there has already been a ~ 30% increase in the oceanic concentration of hydrogen, [H⁺], with a further 55% rise expected once pCO₂ doubles from 280 μatm to 560 μatm (Jokiel 2011). Under these conditions, research has shown that calcification of scleractinian coral may decline by 10-40%, whereas productivity impacts are variable and the overall response is species-specific, with higher sensitivity to OA occurring during early life stages (Kroeker et al. 2010). The variable response of reef-building coral species to OA has made it difficult to model accurate future projections. This research responds to the critical need to characterise intra-specific variation in the response of the common coral species, *Acropora millepora*, to future OA conditions and investigates key carbonate chemistry variables defining their resilience.

Climate models assume a constant oceanic total alkalinity (A_T), as the predicted rate of pCO₂ increase is too rapid for the slower geochemical feedbacks to achieve carbonate chemistry equilibrium (Raven and Falkowski 1999; Kump et al. 2009; Pandolfi et al. 2011). However, as coral reefs exist in the surface layer of the ocean with community metabolism changing the pCO₂ over diel and seasonal timescales, the impact of increased anthropogenic pCO₂ on the carbonate chemistry of coral reefs is likely to deviate from oceanic models with static A_T (Hofmann et al. 2011; Shaw et al. 2013). The process of calcification reduces the A_T by removing carbonate (${\rm CO_3}^{2-}$) or bicarbonate (HCO3-) ions, but A_T is not affected by the CO₂ byproduct of calcification, which is considered to be a source of CO2 to the atmosphere around coral reefs (Bates et al. 2001; Fagan and Mackenzie 2007; Kleypas et al. 2011). In addition, this byproduct is also coupled to photosynthesis whereby CO₂ is removed from the water during the day, leading to a rise in the seawater pH, which contributes to the phenomenon known as light enhanced calcification (Yonge and Nicholls 1931). The diel cycle of carbonate chemistry has been well documented on coral reefs and consists of daytime higher pH due to photosynthesis with lowered A_T due to

calcification compared to nighttime lowered pH due to respiration with higher A_T due to calcium carbonate dissolution (Ohde and vanWoesik 1999; Yates and Halley 2006; Silverman et al. 2007; Santos et al. 2011). As this diel oscillation is driven by residence time of surface waters, shallow reef flats exhibit extreme variability in carbonate chemistry, with seasonal tide and metabolic changes also playing a role (Shamberger et al. 2011; Shaw et al. 2012).

Early research utilized the relationship between A_T and O₂ to derive calcification and primary production measurements of individual corals or whole reef communities (Smith and Key 1975; Barnes 1983). Coupled with future predicted pCO₂ levels, models have described an overall decline in coral community calcification (Silverman et al. 2009; Shaw et al. 2013), yet spatial and temporal variability must be considered to accurately determine the vulnerability of each reef community (Hofmann et al. 2011). A recent study has related the temporal variability in the carbonate chemistry to recruitment success and calcification, with high-Mg calcite coralline algae and bryozoans being particularly sensitive (Price et al. 2012). While *a priori* expectations suggest increased resilience of organisms nocturnally exposed to pH levels predicted by 2100, the potential for carbonate chemistry variability to be used as a predictor of the physiological response of reef-building corals to OA remains largely unexplored.

Despite an abundance of published work, the key drivers of coral calcification remain equivocal (Allemand et al. 2011; Jokiel 2011) and few studies have investigated the potential changes in primary productivity under future OA conditions (Schneider and Erez 2006; Anthony et al. 2008; Crawley et al. 2010). Coral skeletal morphological plasticity is evidence of the response of biomineralisation to environmental heterogeneity, traditionally constrained by flow and light (Lesser et al. 1994; Bruno and Edmunds 1997; Hoogenboom et al. 2008), while OA research has focused on the skeletal mineralogy and rate of calcification (Ries 2011). Furthermore, physiological plasticity is evident by the corals ability to photoacclimate through mechanisms such as manipulating the symbiotic dinoflagellate (genus: *Symbiodinium*) population size or distribution in the tissue (Iglesias-Prieto and Trench 1994; Enriquez et al. 2005) or by the symbionts optimising their photosynthetic pigments (Hoogenboom et al. 2009). Investigating photoacclimatory processes are relevant to this study, as increasing the CO₂ substrate for photosynthesis has been shown to drive photo-physiological

changes (Anthony et al. 2008; Crawley et al. 2010). As photosynthesis and calcification are tightly coupled, it is imperative to consider these together when investigating the response reef-building corals to OA.

Studies investigating the effects of increased pCO₂ on coral physiology have been carried out in aquaria (Ohde and Hossain 2004; Jury et al. 2009), in mesocosms (Langdon et al. 2000; Leclercq et al. 2002; Reynaud et al. 2003) and on the reef using in situ carbon enrichment facilities (Kline et al. 2012). Alternatively, community metabolism calculations and modeling have been used to address this problem (Silverman et al. 2009; Shaw et al. 2012). Although aquaria-based experiments provide an ability to investigate the detailed mechanism behind changes in calcification and productivity in individual species, there may be difficulties in extrapolating this relationship to the ecosystem level due to inter-species variation and interactions. On the other hand, the use of mesocosms and community metabolism models offer insight into community responses, but are often limited to a particular community assemblage. For example, the prediction that reefs will begin to dissolve at 560 ppm CO₂ (Silverman et al. 2009) assumes that the community composition will remain the same, which is very unlikely due to competition, mortality and successional processes, and therefore compromises the future relevance of these types of predictions. Recent studies near naturally occurring CO₂ vents have confirmed that community shifts will occur with increasing pCO₂ (Kroeker et al. 2012; Inoue et al. 2013). In this study, the response of individual corals to future OA conditions has been placed in the context of ecosystem variability. In order to improve future projections of the spatial variability of the response of reef-building corals to OA, the intra-species variation of Acropora millepora was investigated from a range of sites with inherent differences in carbonate chemistry and then a manipulative experiment was performed using samples from these distinct habitats.

In addition to the temporal variation caused by reef metabolism, the carbonate chemistry may exhibit spatial variation due to factors such as the particular community assemblage, water depth, tidal fluctuations and current dynamics (Hofmann et al. 2011; Price et al. 2012). The impact of this temporal and spatial variation in carbonate chemistry on the response of reef-building corals to OA is largely unknown. In this study, the diurnal and seasonal variability in carbonate

chemistry and the community assemblage was assessed at three sites around Lizard Island on the Great Barrier Reef. From these sites, the physiology of *A. millepora* was assessed under future (constant) pCO₂ scenarios in a manipulative, aquaria-based experiment. This study design enabled determination of the coral physiological response to OA in relation to their site of origin, an important consideration to improve the predictive power of climate models as applied to coral reefs.

Materials and Methods

Study Area

The study was conducted at three sites around Lizard Island, Great Barrier Reef, Australia (Figure 1). Station Reef (SR) (14.6798° S, 145.4451° E) is located in front of Lizard Island Research Station (LIRS) in the open lagoon. Loomis Reef (LR) (14.6830° S, 145.4494° E) is also located within the lagoon, while Mermaid Cove (MC) (14.6457° S, 145.4537° E) is located outside the lagoon on the northern side of the Island. While the crest of each reef is exposed at low tide, the reef slopes extend to different depths; approximately 1 m at SR, 3 m at LR and 6 m at MC. The tidal fluctuation of this open lagoon only plays a minor role in the water current dynamics except during spring tides, which occurred during the September 2009 sample time point (see below). The current direction is strongly influenced by the south easterly trade winds which are prevalent from March to November and the intermittent northerly wind from December to February (Crossland and Barnes 1983). Wind speed and direction data was provided via the Australian Institute of Marine Science (AIMS) Integrated Marine Observing System (IMOS).

Carbonate Chemistry

The carbonate chemistry at the sites was determined during three separate trips in September 2009, February 2010 and January 2011. At 3 locations within each site, water samples were collected in 500 mL glass bottles for A_T and pH measurements. The water was collected equidistant from the reef crest at each site to avoid depth bias. Sample frequency was every two hours during the day between the hours of 6:00 to 18:00 except due to logistical constraints. Collection, storage and analysis followed the standard operating procedures as described by Dickson et al. (2007). The bottles were rinsed with sample water three times before collection, sealed underwater to

prevent atmospheric gas exchange and then poisoned with 100 µL of saturated mercuric chloride solution if analysis was to take place more than 12 hours later. A_T was determined by Gran titration (Gran 1952; Dickson et al. 2007) with 0.1M HCl using a titrator (T50, Mettler Toledo, Langacher, Switzerland) modified with a 1 mL burette, pH sensor (DGi101-SC) and small volume stirrer. The pH electrode was calibrated with NBS scale standard buffers (Mettler Toledo) and the titrator was calibrated with Certified Reference Materials (AG Dickson, SIO, Oceanic Carbon Dioxide Quality Control) prior to daily sample analysis. The pH was measured using a pH sensor (DGi101-SC, Mettler Toledo). Measurements were repeated until precision of \pm 3 µmol kg⁻¹ was achieved for A_T and \pm 0.005 for pH. The data was input into the CO2calc program (Robbins et al. 2010) along with ambient temperature, pressure and salinity in order to obtain total dissolved inorganic carbon (DIC), and saturation states for aragonite (Ω_{arag}) and calcite (Ω_{calc}). The calculations were performed using the total hydrogen ion scale, the carbonic acid dissociation constants of Mehrbach et al (1973) as refit by Dickson & Millero (1987) and the sulfonic acid dissociation constant of Dickson (1990).

Community Assemblage

Twenty x 1 m² photo quadrats were randomly taken at each water sample location within the sites and analysed using 100-point intercept method with the Coral Point Count with Microsoft Excel extensions (CPCe) software (Kohler and Gill 2006). The coral, algae and substrate coverage was classified in terms of functional groups and also to genus level, or species level where possible, with the assistance of the Kraft (2007) and the Indo Pacific Coral Finder (Kelley 2009).

OA Experiment

At LIRS in November 2009, 36 branches of *A. millepora* collected near the reef crest at each site (1 - 3 m depth) were hung with nylon line in aquaria with ambient seawater to acclimate for five days before beginning CO_2 treatment. The CO_2 dosing was controlled through a computer interface (Aquatronica-AEB Technologies, Cavriago, Italy) operating solenoid valves, which injected CO_2 gas based on a pH threshold in line with IPCC scenarios B2 (pH 7.9 \pm 0.05; pCO₂ 600 \pm 80 μ atm) and A1FI (pH 7.7 \pm 0.05; pCO₂ 1020 \pm 140 μ atm) (IPCC 2007). The control aquaria had ambient seawater from Lizard Island lagoon (pH 8.1 \pm 0.05; pCO₂ 350 \pm 50 μ atm).

The pH was continuously measured using a pH sensor precise to 0.01 pH units (InPro4501VP, Mettler Toledo) in 200 L sumps and pumped into 4 replicate aquaria for each treatment. The aquaria were maintained outdoors with neutral density shade cloth, which resulted in maximum light levels of 400 µmol quanta m⁻² s⁻¹ as measured by light loggers (Odyssey, Dataflow Systems, Christchurch, New Zealand). The experiment ran for 3 weeks with physiological measurements taken upon completion.

Coral Physiology after OA experiment

Respirometry was performed on the coral samples using the procedure of Crawley et al (2010) with the following modifications. The water bath was set to 27°C, the ambient seawater temperature at Lizard Island in November 2009. Corals were darkacclimated for one hour prior to the start of respirometry analysis. The light program was 10 minutes in darkness to obtain steady-state dark respiration (R_{dark}) followed by 30 minutes at 600 µmol quanta m⁻² s⁻¹ to obtain steady-state net photosynthesis (P_{net}), then 10 minutes in darkness to obtain light-enhanced dark respiration (LEDR), which began approximately 1 minute post illumination. Linear regression of the bulk oxygen (O₂) measurements, obtained with O₂ optodes (Oxy4 v2, PreSens, Regensburg, Germany), provided the rate of O₂ evolution or consumption. The Pg:R ratio was calculated as the ratio of gross photosynthesis ($Pg = P_{net} + LEDR$) to respiration (R =R_{dark} + LEDR), under a 12 hour light: 12 hour dark cycle. A Pg:R ratio greater than 1 denotes that the coral has the potential to grow phototrophically, assuming equal translocation of photosynthates from the dinoflagellates to the coral and equal use of the photosynthates for growth (Muscatine et al. 1981; Amthor 2000). Alternatively, a Pg:R ratio < 1 indicates that the coral must resort to heterotrophy in order to grow.

After completion of the respirometry assays, the branches were frozen in liquid nitrogen the following day at midday, at which time light loggers recorded 300 μ mol quanta m⁻² s⁻¹, which was 25% lower than the average maximum daily light level due to intermittent cloud cover. The branches were later water-piked using 0.22 μ m filtered seawater and centrifuged (4,500 x g, 5 min) to separate the dinoflagellate pellet from host tissue. Two aliquots were made; one for dinoflagellate cell counts using a 0.100 mm Tiefe Depth Profondeur hemocytometer (0.0025 mm²), and the other for pigment analysis using High Performance Liquid Chromatography (HPLC) following the methods of Zapata et al (2000) and Dove et al. (2006). The supernatant

from centrifuging was diluted with 0.22 µm filtered seawater and analysed with a spectrophotometer (SpectraMax M2, Molecular Devices, Sunnyvale, California, United States) to determine host protein using the equations of Whitaker & Granum (1980). In a previous OA experiment, the buoyant weight technique may have been too insensitive to detect a change in the calcification rate of *A. millepora* (Kaniewska et al. 2012). Skeletal dry weight has previously been shown to have perfect correlation with buoyant weight, which has 1-2% error due to the negative buoyancy of the organic matrix and tissue (Jokiel et al. 1978). We therefore measured the skeletal dry-weight to surface-area ratio, which represents a change in the aragonite density and/or branch compactness. The coral surface area was determined using a double-dipping wax method appropriate for skeletons with deep corallites (Stimson and Kinzie 1991; Veal et al. 2010).

Statistical analysis

Site differences in the diurnal carbonate chemistry range were assessed using a oneway PERMANOVA in PRIMER with Site as a fixed factor (Anderson 2001; McArdle and Anderson 2001). We corrected for seasonal baseline differences using the diurnal range in A_T, DIC and pH, which were normalized and projected in Euclidean space for analysis. The diurnal range in Ω_{arag} and Ω_{calc} were similarly analysed. For all PERMANOVA procedures, unrestricted permutation of the raw data created the Pseudo-F distribution for the main test and Monte Carlo P-values were used for pairwise comparisons. We used distance-based linear models (DISTLM) to assess the relationship between the community assemblage and carbonate chemistry predictor variables (Legendre and Anderson 1999; McArdle and Anderson 2001). For this procedure, the community assemblage data was log transformed and a zeroadjusted Bray-Curtis resemblance matrix was generated. Canonical analysis of principle coordinates (CAP) (Anderson and Robinson 2003; Anderson and Willis 2003) was used to determine the best axes through the multivariate data cloud to depict a priori sites. The carbonate chemistry predictor variables found to explain a significant proportion of the data through DISTLM marginal and sequential tests were then overlaid on the CAP model as Pearson correlation coefficients with the CAP axes. We assessed the coral physiological parameters with the PERMANOVA procedure using Euclidean resemblance matrices (Anderson 2001; McArdle and

Anderson 2001). Treatment and Site were specified as fixed factors, while Tank was random and nested within Site. The source of variation was pooled when p > 0.25.

Results

Carbonate Chemistry

The diurnal range in the carbonate chemistry variables was significantly higher at SR compared to MC (Pseudo- $F_{2,24} = 2.79$; $p_{perm} < 0.03$, pairwise MC < SR, $p_{MC} < 0.03$) (Figure 2). Across all seasons sampled the diurnal A_T range was $46 \pm 4 \mu mol \ kg^{-1}$ at SR, $33 \pm 4 \mu mol \ kg^{-1}$ at LR and $28 \pm 4 \mu mol \ kg^{-1}$ at MC. Similarly, the pH range was higher in the lagoon sites than outside the lagoon with SR and LR ranging 0.11 ± 0.01 pH units and 0.12 ± 0.01 pH units, respectively, while MC only 0.08 ± 0.01 pH units. Overnight accumulation of respiratory CO_2 meant that the pH was usually lowest at the morning sample times (See Supplementary Material Figure S1), corresponding with the lowest Ω_{arag} and Ω_{calc} values (Table 1).

Community Assemblage

At the functional group level, there were clear differences in the community assemblage at each site (Figure 3). SR and LR were differentiated from MC by an abundance of soft corals, sea fans and sponges with approximately 55% coverage at SR, 25% coverage at LR and less that 2% coverage at MC. This group was largely composed of the soft corals *Sarcophyton* sp, *Sinularia* sp and *Lobophyton* sp (data not shown). MC was also distinguished by a high abundance of crustose coralline algae (CCA) with up to 20% coverage, whereas SR and LM had less than 3% coverage. Furthermore, branching corals were at approximately 17% cover at MC compared to only approximately 10% at SR and LR.

The DISTLM analysis showed that the community assemblage was distinct at each site and that the seasonal and diurnal carbonate chemistry variables were able to explain a significant proportion of the variation (Figure 4). In September 2009, the maximum daily pH (pH max) significantly explained 25% of the total variation in the community assemblage (Pseudo-F = 2.33, $p_{perm} < 0.02$). In February 2010, the daily minimum A_T (A_T min) and daily maximum A_T (A_T max) significantly explained 26% and 21% of the community assemblage variation, respectively (Pseudo-F = 2.5, p_{perm}

< 0.03). In January 2011, marginal tests showed that the daily DIC minimum (DIC min) and pH max significantly explained 26% and 28% of the variation (Pseudo-F = 2.57, $p_{perm} < 0.02$), although sequential tests indicated overlap in the variation explained as indicated by the plane of the vector on the CAP analysis.

Coral Physiology after OA Experiment

Overall, P_{net} was not significantly affected by the treatment (Pseudo- $F_{2,18} = 3.16$; $p_{perm} = 0.07$), but a one-way PERMANOVA on samples from each site showed that P_{net} increased significantly in MC corals from the B2 to the A1FI scenario (Pseudo- $F_{2,9} = 5.66$; $p_{perm} < 0.03$, pairwise B2 < A1FI, $p_{MC} < 0.02$) (Figure 5). Under the B2 scenario, LEDR significantly increased in corals from SR and MC, but declined in corals from LR (Pseudo- $F_{2,22} = 3.59$; $p_{perm} < 0.04$, pairwise Present \neq B2, $p_{MC} < 0.03$). R_{dark} was significantly higher in corals from SR compared to MC (Pseudo- $F_{2,22} = 3.95$; $p_{perm} < 0.03$, pairwise MC < SR, $p_{MC} < 0.02$). Overall, the Pg:R ratio remained > 1 and there were no significant differences due to treatment or site factors.

Dinoflagellate cell density was significantly higher in SR corals compared to MC across all treatments (Pseudo- $F_{2,9} = 4.94$; $p_{perm} < 0.05$, pairwise SR > MC, $p_{MC} < 0.03$) although there were also some tank effects that may have been related to the diurnal sun angle for which the variability could not be partitioned (Figure 6A). There were no changes in photosynthetic pigments; however, the photoprotective pigment, Diatoxanthin (DT), was upregulated significantly in corals from LR under the B2 scenario (Pseudo- $F_{4,90} = 2.55$; $p_{perm} < 0.04$, pairwise LR: B2 > A1FI, $p_{MC} < 0.02$) (Figure 6B).

Corals from SR had significantly higher protein per surface area than MC corals, although there were some tank effects (Pseudo- $F_{2,18} = 5.80$; $p_{perm} < 0.03$, pairwise MC < SR, $p_{MC} < 0.02$) (Figure 6C). Irrespective of treatment, the skeletal dry-weight to surface-area ratio was higher in corals from MC compared to SR and LM (Pseudo- $F_{2,18} = 10.65$; $p_{perm} < 0.01$, pairwise LR < MC, $p_{MC} < 0.01$, MC > SR, $p_{MC} < 0.01$) (Figure 6D).

Discussion

This study has illustrated the link between benthic community composition and diurnal fluctuations in carbonate chemistry, indicating that the latter may be useful as predictor variables for future models of reef susceptibility to OA. At MC, a site outside Lizard Island lagoon, the diurnal range in carbonate chemistry was significantly smaller than SR, a site inside the lagoon, and this factor could be used to partition the variation in site community composition. A further key finding of this study was the site-specific physiological trade-off demonstrated by variable skeletal morphology and tissue thickness in A. millepora. The skeletal dry-weight to surfacearea ratio, a proxy for aragonite density or branch compactness, was higher in MC coral samples in comparison to SR and LR whereas protein per cm, a proxy for tissue thickness, was higher in SR coral samples compared to MC and LR. In addition, this study has demonstrated that the resilience of A. millepora to future OA may be differentially impacted by the life-history strategies employed by the coral. After three weeks under the pCO₂ conditions of the A1FI scenario, the corals from MC increased P_{net}, yet there were no significant improvements to the Pg:R ratio. This may have been a necessary shift in energy harvest for MC corals in order to increase protein production and tissue thickness in an attempt to buffer the skeleton from acidifying seawater; a response which has been shown in long-term experiments (Fine and Tchernov 2007; Krief et al. 2010). Despite these strategies for survival, a future with less calcification on coral reefs, either through individual decalcification or a shift in community assemblage, seems likely in an elevated-pCO₂ world.

Diurnal variation in carbonate chemistry

The diurnal range in carbonate chemistry reported in this study is comparable to other studies on reef metabolism (Frankignoulle et al. 1996; Gattuso et al. 1996; Bates et al. 2001; Fagan and Mackenzie 2007) although not as extreme as some shallow reef flat studies (Ohde and vanWoesik 1999; Shaw et al. 2012). To our knowledge, the most extreme recording of diel oscillations in Ω_{arag} is 1.83 to 6.36 at Lady Elliot Island reef flat (Shaw et al. 2012) and assuming half the diel range occurred during the daytime, the diurnal range reported here at Lizard Island is still considerably less. It is well understood that coral reef communities can alter the carbonate chemistry of their seawater through the processes of photosynthesis, respiration, calcification, and

dissolution (Smith and Key 1975). As this composite signal is amplified in shallow waters, reef flat studies have subsequently dominated this research. However, it is imperative to understand the full range of diurnal variation in carbonate chemistry and the associated implications for coral reef communities. This study has shown that daily maximum and minimum A_T values can explain up to 40% of the variation in the community assemblage. Although direct causality cannot be attributed, it is highly likely that the carbonate chemistry at MC played a role in the enhanced settlement, growth or survivorship of branching scleractinian corals and CCA in comparison to SR and LR leading to the observed higher percent coverage at MC. This result is not surprising considering the work of Chisholm (2000) suggesting that night-time dissolution lowered the actual accretion rates of CCA on the windward reefs of Lizard Island despite high calcification rates. Previous studies have also predicted enhanced sensitivity of branching scleractinian corals and especially CCA to future elevated pCO₂ scenarios (Anthony et al. 2008; Hall-Spencer et al. 2008; Jokiel et al. 2008; Hurd et al. 2011). On the other hand, the carbonate chemistry at SR and LR were likely more amenable to the life-history characteristics of soft corals, which appear to be physiologically unchanged by elevated pCO₂ conditions (Gabay et al. 2013). Future models of coral reef ecology must consider that reef community composition will be strongly influenced by the diurnal range in their carbonate chemistry.

Coral physiological response to carbonate chemistry

In addition to defining the community, the diurnal variation in carbonate chemistry may also play a role in the underlying morphology of ubiquitous species as exemplified by the site differences in skeletal morphology and protein per cm² in *A. millepora*. Given that calcification responds to changes in atmospheric CO₂ forcing (Kleypas et al. 2011), it is not surprising that skeletal morphology is also potentially shaped by diurnal fluctuations in carbonate chemistry. In this study, the aragonite density or branch compactness of coral samples from outside Lizard Island lagoon was significantly higher than those samples from inside the lagoon at SR and LR. Calcification is enhanced during the day when CO₂ is removed by the process of photosynthesis and hindered each night by the build up of CO₂ from respiration leading to the process of CaCO₃ dissolution. Accordingly, coral skeletal structure will be impacted differentially depending on the extremities of the range in the diel carbonate chemistry and the ability of the coral to maintain optimal conditions in the

extracellular calcifying medium (ECM) (Cohen et al. 2009; Allemand et al. 2011). In order for calcification to occur, the pH must remain elevated in the ECM relative to seawater while Ca²⁺ and HCO₃⁻ must be transported to the ECM. It is likely that the smaller diurnal range in carbonate chemistry was a contributing factor impacting the skeletal morphological plasticity leading to denser aragonite or more compactness with coral samples from outside the lagoon.

One possible avenue to improve the maintenance of optimal calcifying conditions in the ECM is to increase tissue thickness as a means to increase the potential buffering capacity. Increased tissue thickness may restrict entry of seawater and/or result in greater capacity to actively transport ions to and from the ECM. However, as energy allocation to the growth and maintenance of tissue thickness represents a trade-off to skeletal deposition (Anthony et al. 2002), this strategy therefore comes at an energetic cost. In this study, A. millepora at MC had less protein per cm2, but a different skeletal morphology in comparison to A. millepora from SR, which suggests that MC corals were able to build and maintain their skeleton with comparatively less energetic cost. On the other hand, SR corals may have utilized their increased tissue thickness to protect their skeleton from the nocturnal corrosive pH conditions (Chisholm, 2000). While the SR corals aragonite density or branch compactness was still significantly lower than MC, it may have been worse without this additional effort. The higher R_{dark} at SR compared to MC is a testimony to the increased respiratory cost of maintaining thicker tissue (Figure 6C). Given that this trade-off is a life-history strategy that may already be utilized by present day diurnal variation in carbonate chemistry, its future benefit to reef-building corals under OA conditions should be reassessed.

The ability of corals to utilize increased tissue thickness in an attempt to protect their skeleton in the face of future elevated pCO₂ levels also depends on the availability of energy to be allocated to this function. It seemed fortunate for MC corals that P_{net} significantly increased under the A1FI scenario, yet there were no significant differences in the Pg:R ratio, suggesting that the increased energy was being fully utilized. Although we did not detect an increase in protein per cm² in our 3 week study, other long-term experiments have shown that tissue thickness increases under future elevated pCO₂, yet the rate of calcification by those scleractinian corals still

declined (Fine and Tchernov 2007; Krief et al. 2010). So, despite potential increases in photosynthetic output and potential improvements to tissue thickness, the future of coral reefs may still be threatened by decreased skeletal calcification.

Implications for future models

Current research has started to investigate the ecological consequences of OA on coral reefs, with some studies modeling the future carbonate chemistry considering different community composition (Anthony et al. 2011; Shaw et al. 2013). While previous studies have focused on a shift to macroalgal dominance, the observed differences in community composition in the present study indicate that soft coral is also a likely candidate for future reef dominance, which has also been shown near volcanically-acidified waters in Japan (Inoue et al. 2013). This will have cascading effects on the ecology of reefs with different invertebrate, fish and algal communities associated with the habitat of soft corals, which provide overall less structural and topographical complexity.

The overlaid vectors on the CAP model depict the quintessential carbonate chemistry variables that distinguish the community composition of the sites, although these were not common between or within seasons. We believe the driver of this difference was the particular alignment of tidal height and photoperiod. As previous discussed, longer seawater residence times lead to a greater signal of community metabolism, therefore the coincidence of low tide around dawn and dusk will give different values to a low tide regime at midnight and midday. During the September 2009 and February 2010 sample regimes, low tide occurred around midday/midnight whereas January 2011 sampling occurred with low tide on dawn/dusk. In order to partition the variation due to alignment of tidal height and photoperiod, future studies should be designed with comprehensive sampling within seasonal and tidal ranges (Shaw et al. 2012).

Additional factors contributing to the site variability and coral physiology could be included in future studies to further partition the variation. These include current speed and direction, tidal fluctuations, water depth, light attenuation, temperature, nutrients, community metabolism and heterotrophic capacity. This study aimed to investigate the relationship between diurnal variation in carbonate chemistry and

community composition and assess the response of a common reef-building coral from these sites under future OA conditions. Nevertheless, we do not discount the role of the abovementioned factors at these sites. For example, increased heterotrophic feeding at SR could also lead to increased protein and thicker tissue. Furthermore, reflected light from the lagoon sandy bottom could play a greater role in the coral photophysiolgy at SR and LR, in comparison to MC. Similarly, temperature oscillations are likely to be greater at the lagoon sites due to shallow pooled water at low tide with decreased proximity to deeper waters upon inundation at high tide. As our study is purely correlative, we do not exclude these factors as potential drivers and encourage future studies to further investigate.

The natural diel variation in carbonate chemistry on coral reefs offers a unique opportunity to investigate mechanisms of survival under future pCO₂ conditions, akin to studying communities near natural CO₂ vents (Fabricius et al. 2011; Rodolfo-Metalpa et al. 2011). The present study is unique as it investigated the intra-species variation in *A. millepora* in the context of sites with significantly different carbonate chemistry and community assemblage. Site-specific factors have previously been shown to influence the life-history of the dinoflagellates (genus: *Symbiodinium*) growing within *A. millepora* from different sites on the Great Barrier Reef and this led to variability in their response to thermal stress conditions (Howells et al. 2012). Here, the site-specific diurnal variation in carbonate chemistry correlated with differential investment in skeletal morphology and tissue thickness, but this trade-off may be of limited future benefit especially considering the suite of additional factors such as temperature, eutrophication, overfishing, that will impact reef resilience.

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Table 1: Diurnal range of the saturation state of aragonite (Ω_{arag}) and calcite (Ω_{calc}) at the sampling sites around Lizard Island as calculated using the CO2calc program (Robbins et al. 2010).

Site	Sample Timepoint	Range of Ω_{arag}	Range of Ω_{calc}
	September 2009	3.38 - 4.00	5.12 - 6.05
Station Reef	February 2010	4.23 - 4.40	6.33 - 6.58
	January 2011	4.04 - 4.40	6.06 - 6.59
	September 2009	3.94 - 4.00	5.96 - 6.05
Loomis Reef	February 2010	3.72 - 3.88	5.56 - 5.81
	January 2011	4.30 - 4.36	6.45 - 6.52
	September 2009	3.85 - 4.20	5.82 - 6.35
Mermaid Cove	February 2010	3.88 - 4.23	5.81 - 6.34
	January 2011	4.41 - 4.50	6.61 - 6.73

Table 2: PERMANOVA results for *Acropora millepora* physiology after 3 weeks in an OA experiment and the diurnal range in carbonate chemistry at their site of collection.

Variables	Source of Variation	P	Pairwise P
R_{dark} (µmol O_2 cm ⁻² hr ⁻¹)	Treatment	0.25	
	Site	0.02	MC > SR p = 0.0119
	Pooled	0.66	
P_{net} (µmol O_2 cm ⁻² hr ⁻¹)	Treatment	0.07	
	Site	0.17	
	Tank (Site)	0.03	
	Treatment x Site	0.23	
	Treatment x Tank	0.73	
P _{net} MC (μmol O ₂ cm ⁻² hr ⁻¹)	(Site)		A1EI > D2 0.010/
P_{net} WIC (μ IIIOI O_2 CIII III)	Treatment Tank (Treatment)	0.0214 0.3975	A1FI > B2 p = 0.0196
LEDD (um al O am-2 hu-1)			D
LEDR (μ mol O ₂ cm ⁻² hr ⁻¹)	Treatment Site	0.03 0.67	Present \neq B2 p = 0.021
	Tank (Site)	0.67	
	Pooled	0.13	
Cells per cm ²	Treatment	0.4307	
Cens per em	Site	0.0323	MC < SR p = 0.0205
	Tank (Site)	0.004	WIC \ 5K p \ 0.0203
	Treat x Site	0.1868	
	Treat x Tank (Site)	0.2349	
Dt pg per cell	Treatment	0.1234	
	Site	0.1438	
	Tank (Site)	0.3879	
	Treat x Site	0.039	LR: $B2 > A1FI p=0.0135$
Protein mg cm ⁻²	Treatment	0.2865	
	Site	0.028	MC < SR p = 0.0147
	Tank (Site)	0.0141	
	Treatment x Site	0.01911	
Skeleton g cm ⁻²	Treatment	0.4598	
	G!A.	0.0071	LR < MC p=0.0099;
	Site Tonk (Site)	0.0061	SR < LR p=0.0021
A DIC III'	Tank (Site)	0.6494	CD - MC - 0.0440
A _T , DIC, pH diurnal range	Site	0.0263	SR > MC p = 0.0248
			LR > MC P = 0.0794
$\Omega_{ m arag} \Omega_{ m calc}$ diurnal range	Site	0.12	



Figure 1: Map of Lizard Island and surrounding reefs showing the location of sampling sites at Mermaid Cove (MC) 14.6457° S, 145.4537° E, Station Reef (SR) 14.6798° S, 145.4451° E and Loomis Reef (LR) 14.6830° S, 145.4494° E. Image Copyright GeoEye 2005.

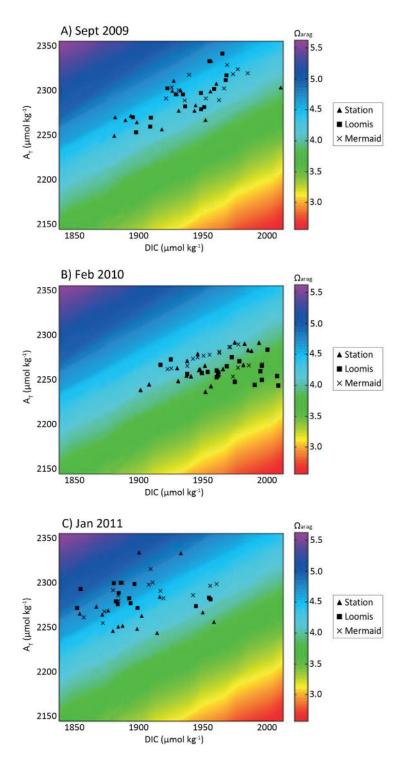


Figure 2: Deffeyes plots depicting Total Alkalinity (A_T) and Dissolved Inorganic Carbon (DIC) at the lagoon sites, Station Reef (SR) and Loomis Reef (LR) and outside the lagoon at Mermaid Cove (MC) during the sampling in A) September 2009, B) February 2010 C) January 2011. Data are overlaid on the contours of aragonite saturation (Ω_{arag}).

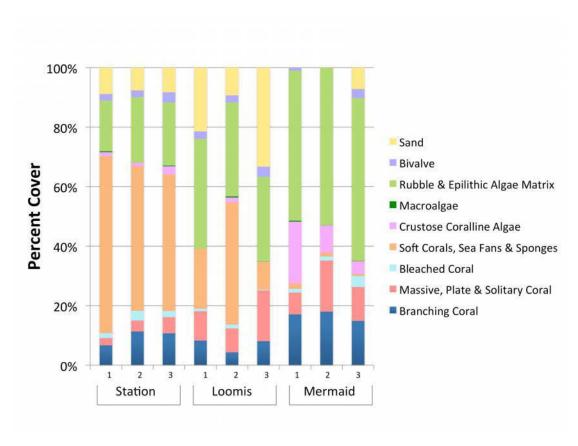


Figure 3: Functional group community assemblage determined from the percent cover at three locations within each site at Lizard Island.

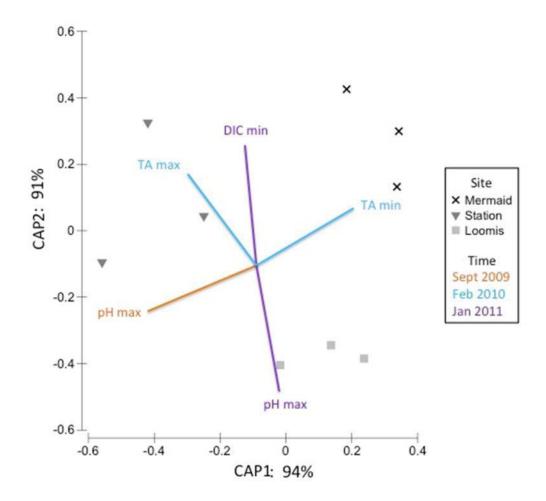


Figure 4: Canonical analysis of principal coordinates (CAP) discriminating the *a priori* sites based on community assemblage. Overlaid vectors are Pearson correlations depicting carbonate chemistry parameters that represent a significant proportion of the variation as partitioned by distance-based linear modelling (DISTLM).

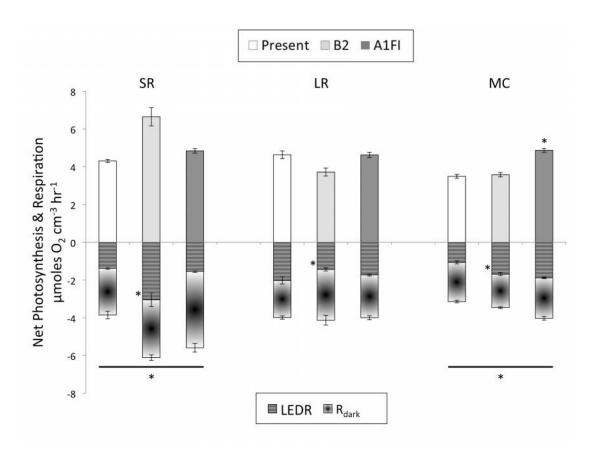


Figure 5: Net Photosynthesis (O_2 production) and Respiration (O_2 consumption) of *Acropora millepora* branches from 3 sites around Lizard Island after 3 weeks in the OA experiment. SR = Station Reef and LR = Loomis Reef are sites inside the lagoon. MC = Mermaid Cove is situated outside the lagoon. LEDR = Light Enhanced Dark Respiration. R_{dark} = Dark Respiration. * = Significant differences ($p_{MC} < 0.05$). Error bars are standard error (n=3).

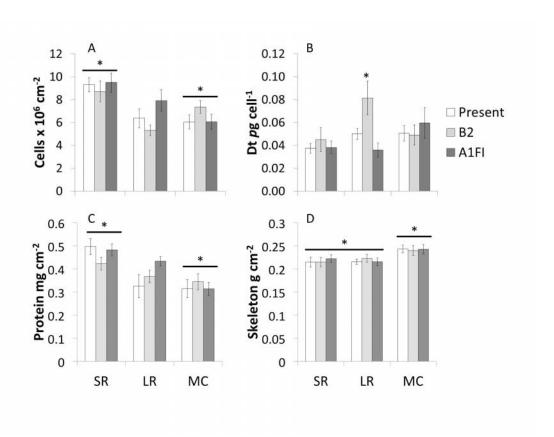


Figure 6: Physiology of *Acropora millepora* collected from 3 sites around Lizard Island after 6 weeks in OA experiment. SR = Station Reef and LR = Loomis Reef are sites inside the lagoon. MC = Mermaid Cove is situated outside the lagoon. Dt = Diatoxanthin * = Significant differences ($p_{MC} < 0.05$). Error bars are standard error (n=3).

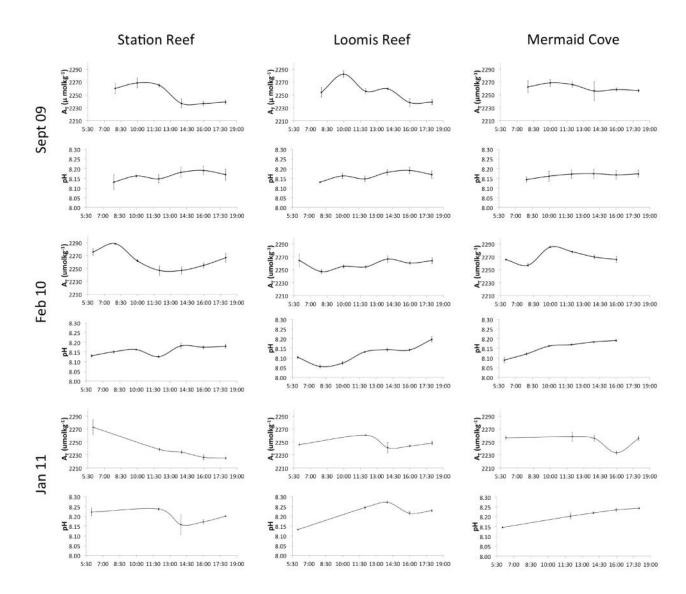


Figure S1: Total Alkalinity (A_T) and pH data from three sites at Lizard Island (Station Reef, Loomis Reef and Mermaid Cove) sampled at three timepoints (September 2009, February 2010 and January 2011). Error bars are standard error (n=3).

References

- Allemand D, Tambutté ZD, Tambutté S (2011) Coral calcification, cells to reefs. In:

 Dubinsky Z, Stambler N (eds) Coral reefs: an ecosystem in transition.

 Springer Press, New York, pp 119-150
- Amthor JS (2000) The McCree-de Wit-Penning de Vries-Thornley Respiration Paradigms: 30 Years Later. Annals of Botany 86:1-20
- Anderson MJ (2001) A new method for non-parametric multivariate analysis of variance. Austral Ecology 26:32-46
- Anderson MJ, Robinson J (2003) Generalised discriminant analysis based on distances. Australian and New Zealand Journal of Statistics 45:301-318
- Anderson MJ, Willis TJ (2003) Canonical analysis of principle coordinates: a useful methods of constrained ordination for ecology. Ecology 84:511-525
- Anthony KRN, Connolly SR, Willis BL (2002) Comparative analysis of energy allocation to tissue and skeletal growth in corals. Limnology & Oceanography 47:1417-1429
- Anthony KRN, Kleypas JA, Gattuso J-P (2011) Coral reefs modify their seawater carbon chemistry implications for impacts of ocean acidification. Global Chang Biology 17:3655-3666
- Anthony KRN, Kline DI, Diaz-Pulido G, Dove S, Hoegh-Guldberg O (2008) Ocean acidification causes bleaching and productivity loss in coral reef builders. Proceedings of the National Academy of Science USA 105:17442-17446
- Barnes DJ (1983) Profiling coral reef productivity and calcification using pH and oxygen electrodes. Journal of Experimental Marine Biology and Ecology 66:149-161
- Bates NR, Samuels L, Merlivat L (2001) Biogeochemical and physical factors influencing seawater fCO₂ and air-sea CO₂ exchange on the Bermuda coral reef. Limnology & Oceanography 46:833-846
- Bruno JF, Edmunds PJ (1997) Clonal variation for phenotypic plasticity in the coral *Madracis mirabilis*. Ecology 78:2177-2190
- Caldeira K, Wickett ME (2003) Anthropogenic carbon and ocean pH. Nature 425:365
- Canadell JG, Quere CL, Raupacha MR, Fielde CB, Buitenhuisc ET, Ciaisf P, Conwayg TJ, Gillettc NP, Houghtonh RA, Marlandi G (2007) Contributions to accelerating atmospheric CO₂ growth from economic activity, carbon

- intensity, and efficiency of natural sinks. Proceedings of the National Academy of Science USA 104:18866-18870
- Chisholm JRM (2000) Calcification by crustose coralline algae on the northern Great Barrier Reef, Australia. Limnology & Oceanography 45:1476-1484
- Cohen AL, McCorkle DC, dePutron S, Gaetani GA, Rose KA (2009) Morphological and compositional changes in the skeletons of new coral recruits reared in acidified seawater: Insights into the biomineralization response to ocean acidification. Geochemistry, Geophysics, Geosystems 10:Q07005
- Crawley A, Kline DI, Dunn S, Anthony K, Dove S (2010) The effect of ocean acidification on symbiont photorespiration and productivity in *Acropora formosa*. Global Change Biology 16:851-863
- Crossland CJ, Barnes D (1983) Dissolved nutrients and organic particulates in water flowing over coral reefs at Lizard Island. Australian Journal of Marine Freshwater Research 34:835-844
- Dickson AG (1990) Standard potential of the reaction AgCl(s) + .5H2(g) = Ag(s) + HCl(aq) and the standard acidity constant of the ion HSO4– in synthetic sea water from 273.15 to 318.15 K. Journal of Chemical Thermodynamics 22:113-127
- Dickson AG, Millero F (1987) A comparison of the equilibrium-constants for the dissociation of carbonic-acid in seawater media. Deep Sea Research Part A 34:1733-1743
- Dickson AG, Sabine CL, Christian J (eds) (2007) Guide to best practices for ocean CO₂ measurements. PICES Special Publication 3
- Dove S, Ortiz J, Fine M, Fisher P, Iglesias-Prieto R, Thornhill D, Hoegh-Guldberg O (2006) Response of holosymbiont pigments from the scleractinian coral *Montipora monasteriata* to short-term heat stress. Limnology & Oceanography 51:1149-1158
- Enriquez S, Mendez ER, Iglesias-Prieto R (2005) Multiple scattering on coral skeletons enhances light absorption by symbiotic algae. Limnology & Oceanography 50:1025-1032
- Fabricius KE, Langdon C, Uthicke S, Humphrey C, Noonan S, De'ath G, Okazaki R, Muehllehner N, Glas MS, Lough JM (2011) Losers and winners in coral reefs acclimatized to elevated carbon dioxide concentrations. Nature Climate Change 1:165-169

- Fagan KE, Mackenzie FT (2007) Air–sea CO₂ exchange in a subtropical estuarinecoral reef system, Kaneohe Bay, Oahu, Hawaii. Marine Chemistry 106:174-191
- Fine M, Tchernov D (2007) Scleractinian coral species survive and recover from decalcification. Science 315:1811
- Frankignoulle M, Gattuso J-P, Biondo R, Bourge I, Copin-Montegut G, Pichon P (1996) Carbon fluxes in coral reefs II Eulerian study of inorganic carbon dynamics and measurement of air-sea CO₂ exchanges. Marine Ecology Progress Series 145:123-132
- Gabay Y, Benayahu Y, Fine M (2013) Does elevated pCO₂ affect reef octocorals? Ecology and Evolution 3:465-473
- Gattuso JP, Pichon M, Delesalle B, Canon C, Frankignoulle M (1996) Carbon fluxes in coral reefs. I. Lagrangian measurement of community metabolism and resulting air-sea CO₂ disequilibrium. Marine Ecology Progress Series 145:109-121
- Gran G (1952) Determination of the equivalence point in potentiometric titrations:

 Part II. Analyst 77:661-670
- Hall-Spencer JM, Rodolfo-Metalpa R, Martin S, Ransome E, Fine M, Turner SM, Rowley SJ, Tedesco D, Buia M-C (2008) Volcanic carbon dioxide vents show ecosystem effects of ocean acidification. Nature 454:96-99
- Hofmann GE, Smith JE, Johnson KS, Send U, Levin LA, Micheli F, Paytan A, Price NN, Peterson B, Takeshita Y, Matson PG, Crook ED, Kroeker KJ, Gambi MC, Rivest EB, Frieder CA, Yu PC, Martz TR (2011) High-Frequency Dynamics of Ocean pH: A Multi-Ecosystem Comparison. PLoS One 6:e28983
- Hoogenboom MO, Connolly SR, Anthony KRN (2008) Interactions between morphological and physiological plasticity optimize energy acquisition in corals. Ecology 89:1144-1154
- Hoogenboom MO, Connolly SR, Anthony KRN (2009) Effects of photoacclimation on the light niche of corals: a process-based approach. Marine Biology 156:2493-2503
- Howells EJ, Beltran VH, Larsen NW, Bay LK, Willis BL, van Oppen MJH (2012) Coral thermal tolerance shaped by local adaptation of photosymbionts. Nature Climate Change 2:116-120

- Hurd CL, Cornwall CE, Currie K, Hepburn CD, McGraw CM, Hunter KA, Boyd PW (2011) Metabolically induced pH fluctuations by some coastal calcifiers exceed projected 22nd century ocean acidification: a mechanism for differential susceptibility? Global Change Biology 17:3254-3262
- Iglesias-Prieto R, Trench RK (1994) Acclimation and adaptation to irradiance in symbiotic dinoflagellates. I. Responses of the photosynthetic unit to changes in photon flux density. Marine Ecology Progress Series 113:163-175
- Inoue S, Kayanne H, Yamamoto S, Kurihara H (2013) Spatial community shift from hard to soft corals in acidified water. Nature Climate Change 3:683-687
- IPCC (2007) The fourth assessment report of the Intergovernmental Panel on Climate Change (IPCC). Cambridge University Press, Cambridge, UK
- Jokiel PL (2011) Ocean Acidifcation and control of reef coral calcification by boundary layer limitation of proton flux. Bulletin of Marine Science 87:639–657
- Jokiel PL, Rodgers KS, Kuffner IB, Andersson AJ, Cox EF, Mackenzie FT (2008)

 Ocean acidification and calcifying reef organisms: a mesocosm investigation.

 Coral Reefs 27:473-483
- Jokiel RL, Maragos JE, Franzisket L (1978) Coral growth: buoywant weight technique. In: Stoddart DR, Johannes RE (eds) Monographs on Oceanographic Methodology. UNESCO, Paris, France, pp529-542
- Jury CP, Whitehead RF, Szmant AM (2009) Effects of variations in carbonate chemistry on the calcification rates of *Madracis auretenra* (= *Madracis mirabilis sensu* Wells, 1973): bicarbonate concentrations best predict calcification rates. Global Change Biology 16:1632-1644
- Kaniewska P, Campbell PR, Kline DI, Rodriguez-Lanetty M, Miller DJ, Dove S, Hoegh-Guldberg O (2012) Major cellular and physiological impacts of ocean acidification on a reef building coral. Plos One 7:e34659
- Kelley R (2009) Indo Pacific Coral Finder. BYO Guides, Townsville
- Kleypas JA, Anthony KRN, Gattuso J-P (2011) Coral reefs modify their seawater carbon chemistry case study from a barrier reef (Moorea, French Polynesia). Global Change Biology 17:3667-3678
- Kline DI, Teneva L, Schneider K, Miard T, Chai A, Marker M, Headley K, Opdyke B, Nash M, Valetich M, Caves JK, Russell BD, Connell SD, Kirkwood BJ, Brewer P, Peltzer E, Silverman J, Caldeira K, Dunbar RB, Koseff JR,

- Monismith SG, Mitchell BG, Dove S, Hoegh-Guldberg O (2012) A short-term in situ CO₂ enrichment experiment on Heron Island (GBR). Scientific Reports 2:413
- Kohler KE, Gill SM (2006) Coral Point Count with Excel extensions (CPCe): A Visual Basic program for the determination of coral and substrate coverage using random point count methodology. Computers and Geosciences 32:1259-1269
- Kraft GT (2007) Algae of Australia: marine benthic algae of Lord Howe Island and the Southern Great Barrier Reef. Australian Biological Resources Study and CSIRO Publishing, Canberra
- Krief S, Hendy EJ, Fine M, Yamd R, Meibom A, Foster GL, Shemesh A (2010)

 Physiological and isotopic responses of scleractinian corals to ocean acidification. Geochimica et Cosmochimica Acta 74:4988–5001
- Kroeker KJ, Micheli F, Gambi MC (2012) Ocean acidification causes ecosystem shifts via altered competitive interactions. Nature Climate Change 3:156-159
- Kroeker KJ, Kordas RL, Crim RN, Singh GG (2010) Meta-analysis reveals negative yet variable effects of ocean acidification on marine organisms. Ecology Letters 13:1419-1434
- Kump LR, Bralower TJ, Ridgwell A (2009) Ocean acidification in deep time.

 Oceanography 22:94-107
- Langdon C, Takahashi T, Sweeney C, Chipman D, Goddard J, Marubini F, Aceves H, Barnett H, Atkinson M (2000) Effect of calcium carbonate saturation state on the calcification rate of an experimental coral reef. Global Biogeochemical Cycles 14:639-654
- Leclercq N, Gattuso J-P, Jaubert J (2002) Primary production, respiration, and calcification of a coral reef mesocosm under increased CO₂ partial pressure. Limnology & Oceanography 47:558-564
- Legendre P, Anderson MJ (1999) Distance-based redundancy analysis: testing multispecies responses in multifactorial ecological experiments. Ecological Monographs 75:435-450
- Lesser MP, Weisb VM, Patterson MR, Jokiel PL (1994) Effects of morphology and water motion on carbon delivery and productivity in the reef coral, *Pocillopora damicornis* (Linnaeus): Diffusion barriers, inorganic carbon

- limitation and biochemical plasticity. Journal of Experimental Marine Biology and Ecology 178:153-179
- McArdle BH, Anderson MJ (2001) Fitting multivariate models to community data: a comment on distance-based redundancy analysis. Ecology 82:290-297
- Mehrbach C, Culberso CH, Hawley JE, Pytkowic RM (1973) Measurement of apparent dissociation-constants of carbonic acid in seawater at atmospheric-pressure. Limnology & Oceanography 1973:897-901
- Muscatine L, McCloskey L, Marian RE (1981) Estimating the daily contribution of carbon from zooxanthellae to coral animal respiration. Limnology & Oceanography 26:601-611
- Ohde S, vanWoesik R (1999) Carbon dioxide flux and metabolic processes of a coral reef, Okinawa. Bulletin of Marine Science 65:559-576
- Ohde S, Hossain MMM (2004) Effect of CaCO₃ (aragonite) saturation state of seawater on calcification of *Porites* coral. Geochemical Journal 38:613-621
- Pandolfi JM, Connolly SR, Marshall DJ, Cohen AL (2011) Projecting coral reef futures under global warming and ocean acidification. Science 333:418-422
- Price NN, Martz TR, Brainard RE, Smith JE (2012) Diel variability in seawater pH relates to calcification and benthic community structure on coral reefs. PLoS One 7:e43843
- Raven JA, Falkowski PG (1999) Oceanic sinks for atmospheric CO₂. Plant Cell nad Environment 22:741-755
- Reynaud S, Leclercq N, Romaine-Lioud S, Ferrier-Pages C, Jaubert J, Gattuso J (2003) Interacting effects of CO₂ partial pressure and temperature on photosynthesis and calcification in a scleractinian coral. Global Change Biology 9:1660-1668
- Ries JB (2011) Skeletal mineralogy in a high-CO₂ world. Journal of Experimental Marine Biology and Ecology 403:54-64
- Robbins LL, Hansen ME, Kleypas JA, Meylan SC (2010) CO2calc—A user-friendly seawater carbon calculator for Windows, Max OS X, and iOS (iPhone). U.S. Geological Survey, Florida
- Rodolfo-Metalpa R, Houlbreque F, Tambutte E, Boisson F, Baggini C, Patti FP, Jeffree R, Fine M, Foggo A, Gattuso J-P, Hall-Spencer JM (2011) Coral and mollusc resistance to ocean acidification adversely affected by warming. Nature Climate Change 1:308-312

- Sabine CL, Feely RA, Gruber N, Key RM, Lee K, Bullister JL, Wanninkhof R, Wong CS, Wallace DWR, Tilbrook B, Millero FJ, Peng T-H, Kozyr A, Ono T, Rios AF (2004) The oceanic sink for anthropogenic CO₂. Science 305:367-371
- Santos IR, Glud RN, Maher D, Erler D, Eyre BD (2011) Diel coral reef acidification driven by porewater advection in permeable carbonate sands, Heron Island, Great Barrier Reef. Geophysical Research Letters 38:L03604
- Schneider K, Erez J (2006) The effect of carbonate chemistry on calcification and photosynthesis in the hermatypic coral *Acropora eurystoma*. Limnology & Oceanography 51:1284-1293
- Shamberger KEF, Feely RA, Sabine CL, Atkinson MJ, DeCarlo EH, Mackenzie FT, Drupp PS, Butterfield DA (2011) Calcification and organic production on a Hawaiian coral reef. Marine Chemisty 127:64-75
- Shaw EC, McNeil BI, Tilbrook B (2012) Impacts of ocean acidification in naturally variable coral reef flat ecosystems. Journal of Geophysical Research 117:C03038
- Shaw EC, McNeil BI, Tilbrook B, Matear R, Bates ML (2013) Anthropogenic changes to seawater buffer capacity combined with natural reef metabolism induce extreme future coral reef CO₂ conditions. Global Change Biology 19:1632-1641
- Silverman J, Lazar B, Erez J (2007) Community metabolism of a coral reef exposed to naturally varying dissolved inorganic nutrient loads. Biogeochemistry 84:67-82
- Silverman J, Lazar B, Cao L, Caldeira K, Erez J (2009) Coral reefs may start dissolving when atmospheric CO₂ doubles. Journal of Geophysical Ressearch 36:L05606, doi:05610.01029/02008GL036282
- Smith SV, Key GS (1975) Carbon dioxide and metabolism in marine environments. Limnology & Oceanography 20:493-495
- Stimson J, Kinzie RA (1991) The temporal pattern and rate of release of zooxanthellae from the reef coral *Pocillopora damicornis* (Linnaeus) under Nitrogen-enrichment and control conditions. Journal of Experimenal Marine Biology and Ecology 153:63-74
- Veal CJ, Holmes G, Nunez M, Hoegh-Guldberg O, Osborn J (2010) A comparative study of methods for surface area and three dimensional shape measurement of coral skeletons. Limnology & Oceanography: Methods 8:241-253

- Whitaker JR, Granum PE (1980) An absolute method for protein determination based on differences in absorbance at 235 and 280 nm. Anal Biochem 109:156-159
- Yates KK, Halley RB (2006) CO₃²⁻ concentration and pCO₂ thresholds for calcification and dissolution on the Molokai reef flat, Hawaii. Biogeosciences 3:357-369
- Yonge CM, Nicholls AG (1931) Studies on the physiology of corals. V. On the relationship between corals and zooxanthellae. In: Yonge CM (ed) Scientific Reports Great Barrier Reef Expedition. Trustees of the British Museum, London, UK, pp177-211
- Zapata M, Rodriguez F, Garrido JL (2000) Separation of chlorophylls and carotenoids from marine phytoplankton: a new HPLC method using a reversed phase C8 column and pyridine containing mobile phases. Marine Ecology Progress Series 195:29-45

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Chapter 4: Coral bleaching in Acropora millepora juveniles under

ocean acidification

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Abstract

Ocean acidification (OA) presents a challenge to the productivity of reef-building corals, with recent studies indicating that OA can cause coral bleaching (loss of endosymbionts or photosynthetic pigments). dinoflagellate While dinoflagellate types (genus Symbiodinium) can infer increased thermal tolerance to the holobiont, other types are more sensitive to heat stress. However, the effects of OA on coral bleaching in relation to symbiont type are largely unknown. In this study, we investigated the impact of OA on the photo-physiology of Symbiodinium within newly settled Acropora millepora recruits during the establishment of endosymbiosis. After two months under increased CO₂ conditions (pH 7.81 ± 0.10 and pH $7.60 \pm$ 0.10), coral recruits were significantly bleached > 2.5 fold relative to controls, using a visual bleaching index at the centre of the polyp. This bleaching response correlated with recruits containing a dominant population of Symbiodinium type D1 or D1-4, indicating that OA may exhibit a phenotypic selective pressure on A. millepora recruits, with implications for post-settlement survivorship and population dynamics. The benthic functional groups surrounding the recruits significantly explained the coral photo-physiological response with crustose coralline algae corresponding with ambient CO₂ levels, while turf and endolithic algae correlated with high CO₂ levels. Overall, this indicates that the additional impact of CO₂ on benthic functional groups may indirectly affect post-settlement survivorship as the diel metabolic cycles of turf and endolithic algae could potentially exacerbate OA conditions.

Introduction

Over the past century, human activities have become more carbon intensive, leading to an increase in atmospheric carbon dioxide (CO₂) emissions and subsequent reduction in oceanic pH (Sabine et al. 2004; Canadell et al. 2007). This shift in carbonate chemistry, termed ocean acidification (OA), is a result of oceanic uptake of CO_2 , having already led to a $\sim 30\%$ increase in the concentration of hydrogen ions $[H^{+}]$ and $\sim 15\%$ decline in carbonate ions (CO₃²⁻) in oceans worldwide (Kleypas et al. 2006). The ocean pH has not dropped this low for the last two million years and while OA events have coincided with coral mass extinctions in the fossil record, the current rate of pH decline is unprecedented (Veron 2008; Kump et al. 2009; Pandolfi et al. 2011). Should the remaining fossil fuel resources be utilised, oceanic pH may drop a further 0.7 units with critical implications for marine calcification and primary productivity (Caldeira and Wickett 2003; Kroeker et al. 2010; Jokiel 2011). Recent studies have indicated that OA will present a challenge for the calcification of reefbuilding corals (Schneider and Erez 2006; Marubini et al. 2008), especially in their early life history stages (Albright and Langdon 2011; Doropoulos et al. 2012b). Furthermore, some adult corals have exhibited a bleaching response (loss of endosymbionts and/or their pigments) to elevated CO₂ (Anthony et al. 2008) potentially due to changes in photoprotection (Crawley et al. 2010), yet there are very few studies investigating the photo-physiology of juvenile corals under OA conditions. Here we describe the response of Acropora millepora recruits after postsettlement exposure to CO₂ levels predicted under Intergovernmental Panel on Climate Change (IPCC) scenarios.

Disruption of the mutualistic symbiosis between corals and their algal symbionts results in coral bleaching, which occurs in response to a variety of stressors including temperature (Jokiel and Coles 1990), light (Falkowski et al. 1984), CO₂ (Anthony et al. 2008), salinity (Hoegh-Guldberg and Smith 1989), nutrients (Wiedenmann et al. 2012) and sedimentation (Anthony et al. 2007). One of the first symptoms of coral bleaching is impairment of the photosynthetic apparatus, which can be measured through changes in the photochemical efficiency of Photosystem II (PSII) (Warner et al. 1999). A decrease in the variable fluorescence yield in relation to the maximum fluorescence yield (F_v/F_m) indicates closure of the PSII reaction centres, which may

reduce the rate of photosynthesis and lead to production of damaging reactive oxygen species (ROS) (Hill and Ralph 2008). Dinoflagellate density and pigments may be adjusted in order to photoacclimate to changing conditions (Falkowski and Dubinsky 1981; Walters 2005). The differential ability to photoacclimate, between (Iglesias-Prieto et al. 2004) and within *Symbiodinium* types (Howells et al. 2012), may have important implications for the future survival of the coral holobiont. In this study, we consider the photochemical efficiency of *Symbiodinium* types at the onset of symbiosis under OA conditions.

In shallow reef flat areas characterised by a smaller volume of water, large diel oscillations in carbonate chemistry are detectable, driven by community metabolism and calcification (Shaw et al. 2012; Anthony et al. 2013). The proportion of photosynthetic and calcifying organisms within the coral reef community can significantly influence the carbonate chemistry (Anthony et al. 2011; Kleypas et al. 2011). Diel variation in carbonate chemistry occurs as photosynthesis and calcification are coupled during the day, with photosynthesis removing CO₂ from the seawater, thereby increasing the pH and providing optimal conditions for the precipitation of calcium carbonate. On the other hand, as night respiration increases CO₂, this creates conditions optimal for dissolution of calcium carbonate. Due to the small size of coral recruits, the impact of OA may vary depending on the surrounding community composition that can alter carbonate chemistry within the diffusive boundary layer (DBL) (Shashar et al. 1996; Hurd et al. 2011; Jokiel 2011). In addition, OA has previously been shown to indirectly affect coral settlement due to changes in the community composition of the substrate (Albright and Langdon 2011; Doropoulos et al. 2012a) and interactions among coral larvae and crustose coralline algae (CCA) (Doropoulos et al. 2012a; Doropoulos and Diaz-Pulido 2013). Accordingly, in this study we characterise the community composition surrounding the coral recruits in order to assess whether this impacts the photo-physiological response of the symbiont to OA conditions.

Small variations in factors that influence recruit success, such as competition or physiological defects, can have profound impacts on coral community structure and their evolutionary development (Gosselin and Qian 1997). Juvenile mortality is naturally high among benthic marine invertebrates, with a less than 20% survival rate

past the first four months, while post-settlement survivorship of corals can be as low as 0.2-6.0% (Gosselin and Qian 1997; Fairfull and Harriott 1999; Wilson and Harrison 2005). Stochastic events and chronic stressors dictate survivorship during these early life history stages (Vermeij and Sandin 2008). The situation is complicated for corals as selective pressures can act phenotypically or genetically on the coral host (Marfenin 1997) or their symbiotic dinoflagellates (genus: Symbiodinium) (Howells et al. 2012) and potential combinations of both (Little et al. 2004; LaJeunesse et al. 2010). Nine major genetic clades (A-I) exist within the genus Symbiodinium (Rowan and Powers 1991; Pochon and Gates 2010) and these clades contain genetically and ecologically distinct types (LaJeunesse et al. 2004a; LaJeunesse et al. 2004b; Warner et al. 2006; Pochon et al. 2007; Sampayo et al. 2007). Differences in physiological traits among Symbiodinium include cell size (LaJeunesse et al. 2005), pigment composition (Frade et al. 2008), photosynthetic performance (Iglesias-Prieto et al. 2004); photosynthate composition (Loram et al. 2007) and tolerance to heat stress (Rowan and Knowlton 1995; Iglesias-Prieto et al. 2004; Berkelmans and van Oppen 2006; Sampayo et al. 2008). In particular, Symbiodinium types D1 and D1-4 (D1-4 provisionally named Symbiodinium trenchi) have been related to increased thermal tolerance of the coral holobiont, while certain C types are more sensitive to heat stress than other types (Berkelmans and van Oppen 2006; Wham et al. 2011). In contrast, the differential response of the coral holobiont to the effects of OA in relation to the type of Symbiodinium they harbour is largely unknown [except see Noonan et al. (2013)]. While symbiont community shifts may occur in some adult reef-building corals in response to stress (Berkelmans and van Oppen 2006; Jones et al. 2008), the population reverts back to the original symbiont community when the stress subsides and is generally stable over time (Thornhill et al. 2006; Finney et al. 2010). There is, however, greater opportunity for selective pressures to act during early ontogeny due to significant flexibility in the establishment of endosymbiosis (Gómez-Cabrera et al. 2008; Abrego et al. 2009; Cumbo et al. 2013). In the present study, we investigate the potential for the chronic stressor, OA, to exhibit selective pressure on A. millepora through association with Symbiodinium.

Methods

Coral spawning and settlement

Five gravid adult colonies of *Acropora millepora* were collected from Heron Island reef flat (23°26.73' S, 151°54.77' E) and placed in separate aquaria with ambient flow-through seawater. Spawning occurred seven nights after the full moon on 29 November 2010 and egg-sperm bundles were collected from each colony. Gametes were released by gentle agitation of the bundles and cross-fertilised for two hours. Embryos were then placed in 200 L tubs with aeration and for the first day approximately half the ambient seawater was changed every four hours to avoid fouling. Thereafter, water changes occurred as needed every 6 - 12 hours.

Five nights after spawning, terracotta and limestone tiles were added to the tubs for settlement. The tiles measured approximately 5 x 5 x 0.5 cm, were stacked in pairs and separated by a 0.5 cm spacer. In the six months prior to the experiment, those tiles had been pre-conditioned on Heron Island reef flat to develop a community of crustose coralline algae (CCA) and microbial biofilms necessary to cue settlement and metamorphosis of the planulae larvae (Negri et al. 2001; Webster et al. 2004). Fouling organisms had been carefully removed from the tiles using plastic scrapers, toothbrushes and tweezers. Following three days settlement time, we scored the number of recruits on each tile using a dissecting microscope, and the tiles were randomly allocated to CO₂ treatments. Each CO₂ treatment had seven replicate flowthrough aquaria randomly placed on an outdoor table.

Experimental Protocol

Seawater carbonate chemistry was maintained via a computer-controlled solenoid valve system (Aquatronica-AEB Technologies, Cavriago, Italy), which bubbled CO_2 to levels representing IPCC (Intergovernmental Panel on Climate Change) CO_2 scenarios for the year 2100. The water temperature remained ambient (mean 27 \pm 1°C) throughout the experiment. The A1FI treatment was pH 7.60 \pm 0.10, the B2 treatment was pH 7.81 \pm 0.10 and the control measured pH 8.04 \pm 0.14 (Table 1). The pH sensors (InPro4501VP, Mettler Toledo, Langacher, Switzerland), calibrated with NBS scale standard buffers (Mettler Toledo), were used to continuously monitor the

pH in the 200 L tubs and the seven replicate aquaria. The aquaria were maintained outdoors and covered with neutral density filter (Lee 298 ND 0.15; LEE Filters Limited, Andover, UK) and thin shade cloth which resulted in light averaging 140 µmol quanta m⁻² s⁻¹ during a 12 hour photoperiod. These low light levels were maintained as corals commonly recruit to cryptic, low-light habitats (Babcock and Mundy 1996; Baird and Hughes 2000).

In order to fully characterise the carbonate chemistry, the treatment water was sampled every six hours for 48 hours during a spring and neap tidal regime for Total Alkalinity (A_T) measurement via Gran titration (T50 titrator, Mettler Toledo) (Gran 1952; Dickson et al. 2007). The A_T and pH were entered, along with ambient temperature (27 °C), pressure (10.16 dbars) and salinity (35 PSU), into the CO₂calc program (Robbins et al. 2010) (Table 1). For the calculations, we specified the total hydrogen ion scale, the carbonic acid dissociation constants of Mehrbach (1973) as refit by Dickson and Millero (1987) and the sulfonic acid dissociation constant of Dickson (1990). We tested the statistical significance of the treatments using a one-way PERMANOVA in PRIMER-e (v 6.1.13), which indicated that A_T was not statistically significant between CO_2 treatments (Pseudo $F_{2,47} = 0.0005$, p = 0.999) and all other parameters were statistically distinct between CO_2 treatments (Pseudo $F_{2,47} > 43.25$, p = 0.0001) (Anderson et al. 2008).

To determine whether CO₂ treatment would change *Symbiodinium* uptake, the recruits were offered a variety of cultured *Symbiodinium* to supplement those that would naturally reside in the environment (Coffroth et al. 2006; Porto et al. 2008) (Table 2). *A. millepora* recruits have been known to establish symbiosis within 5 - 13 days of settlement (Babcock 1985; Babcock and Heyward 1986). Beginning five days after settlement, we isolated the flow-through system for one hour to disperse approximately 1.2 x 10⁶ cells of each *Symbiodinium* clade diluted in 5 mL culture media [Guillard's (F/2) Marine Water Enrichment Solution, Sigma-Aldrich, Australia] throughout the 15 L aquaria. This process was repeated for three days and then 23 days after settlement we randomly selected one tile per tank and scored the percentage of recruits with endosymbionts using the dissecting microscope. These tiles were subsequently returned to their respective treatments. For all tile analyses,

we only used the underside of the tile pairs to avoid confounding effects due differential to light levels.

Response Variables

After two months in the OA treatments we randomly selected three tiles per tank and five recruits per tile and to assess photochemical efficiency of Photosystem II (PSII) in Symbiodinium with the imaging Pulse Amplitude Modulating Fluorometer (iPAM, Walz, Effeltrich, Germany). Prior to commencement of the induction curves, the tiles were placed in a dark container with treatment water to dark-acclimate for 15 minutes. An initial saturating pulse enabled calculation of the maximum quantum yield of PSII, F_{ν}/F_{m} . The actinic light was then set for one minute at 110 µmol quanta m⁻² s⁻¹, which was similar to the average ambient light in the experimental aquaria. Another saturating pulse applied after this illumination enabled partitioning of the absorbed excitation energy between the light used for photochemistry, and excitation energy dissipated either through non-photochemical quenching (Φ_{NPO}) or fluorescence (Φ_{NO}). Furthermore, we calculated $\Delta F/F_m$ ' and used this to derive Excitation Pressure ($Q_m =$ $1 - [(\Delta F/F_m') \cdot (F_v/F_m)^{-1}])$ (Iglesias-Prieto et al. 2004). Thereafter, the recruits recovered for five minutes in the dark and a final saturating pulse was applied. The recruit positions were mapped from high-resolution digital images of the tiles and aligned with the iPAM images in order to define the areas of interest in the iPAM software (ImagingWin, Walz).

Using the same recruits, bleaching was quantified as a reduction in luminescence relative to the maximum, signifying symbiont or chlorophyll density, following the method of Anthony et al. (2008). We used the C1 - C6 scale of luminescence on the CoralWatch Coral Health Chart (CoralWatch 2001), with C1 being 100% bleached and C6 being 0% bleached. The luminescence measurement was performed for areas of interest defined at the centre of the polyp and the whole recruit from the digital images of the tiles in Photoshop (Adobe Systems Software).

In addition, the benthic community at the edge of these recruits was quantified using the point-intercept method by placing an eight-point circular grid around the recruits on the digital images of the tiles in Photoshop (Adobe Systems Software, Ireland). We classified the substrata into the following benthic functional groups: bare tile; live CCA; dead CCA; encrusting fleshy algae; turf + endolithic algae (where endoliths were within the surrounding substrata as opposed to the coral recruit skeleton); shell (including empty polychaete tubes and gastropod shells) and coral recruit. For images of the recruits surrounded by examples of the benthic functional groups see Figure 1.

Genotyping

Following the iPAM assays, the recruits were scraped from the tiles using a sterile surgical blade and preserved in salt-saturated 20% dimethyl sulfoxide (SS-DMSO) (Seutin et al. 1991). DNA was extracted using the PowerBiofilm DNA isolation kit (Mo Bio Laboratories Inc, California, USA) according to the manufacturers protocol Symbiodinium specific primers and amplified using 'ITS1CLAMP' (5' CGCCCGCCGCCCCGCCCCGCCCGCCCGCCCGGGATCCGTTTC CGTAGGTGAACCTGC 3') and 'ITS1intrev2' (5' TTCACGGAGTTCTGCAAT 3') targeting the internal transcribed spacer region 1 (ITS1) of the ribosomal genes (LaJeunesse et al. 2008). PCR amplification (initialization: 94°C, 3 min; denaturation of 35 cycles: 94°C, 40 sec: annealing: 62°C, 40 sec; elongation: 72°C, 30 sec; final step: 72°C, 10 min) was followed by polymorphism screening using denaturing gradient gel electrophoresis (DGGE) in a vertical system (CBS Scientific, San Diego, USA) (LaJeunesse 2001; LaJeunesse et al. 2003) and run for 14 hours on 8% polyacrylamide gels (37.5:1 acrylamide/bis) with a gradient of 35-60% denaturants (formamide and urea). DGGE was used to give resolution at the level of Symbiodinium types rather than at clade level although it is unable to detect background types < 10% of the population (Fabricius et al. 2004; Thornhill et al. 2006). Up to five representative samples of each characteristic ITS1-DGGE fingerprint were used to identify the Symbiodinium type by stabbing the dominant DNA bands from the denaturing gel using a 10 µl pipette tip and placing the tip in 30 ul H₂O for two hours. The DNA was then re-amplified using the 'ITS1intfor' primer (without the GC-clamp) and reverse primer 'ITS1 intrev' using the same PCR thermal cycle profile as mentioned above. PCR products were cleaned using ExoSap-IT (GE Healthcare Limited, Buckinghamshire, UK) according to the manufacturers' protocol and sequenced at the Macrogen (Seoul, Korea) using an ABI 3730xl sequencer. Sequence chromatograms were visually checked using Codoncode Aligner version

3.5.3. (Codoncode Corporation). The resulting sequences were blasted against previously recorded *Symbiodinium* types on GenBank (http://www.ncbi.nih.gov).

Statistical Analysis

To determine the effect of CO₂ treatment, multivariate analyses were run on the groups of response variables (iPAM, bleaching, benthic functional groups, Symbiodinium types) in PRIMER-e (v 6.1.13) with PERMANOVA add-on (v 1.03) (Anderson et al. 2008). A mixed effects model was used in which CO₂ treatment was specified as a fixed factor, tank as a random factor nested in treatment and tile as a random factor nested in tank (Anderson 2001). The iPAM and bleaching data were normalised and resemblance matrices were based on Euclidean distance due to the different scales used for response variables. Benthic substrate was square root transformed and Bray Curtis similarity with dummy variable (+1) added to define joint zeros was used for the resemblance matrix. Symbiodinium data consisted of type presence/absence data and type dominance data, which were used to create a Jaccard resemblance matrix. The *Symbiodinium* types were designated dominant when the band intensity was > 30% luminescence relative to the maximum as quantified from the gel images in Photoshop (Adobe Systems Software), using the same method described above for the bleaching index (Anthony et al. 2008). The effect of CO₂ treatment on individual variables was also assessed using a univariate PERMANOVA using the same mixed effect model previously described. Type (III) Partial Sums of Squares were specified and the Pseudo-F distribution was created by 9999 permutations of the residuals under a reduced model for the main test, whereas subsequent pairwise tests used Monte Carlo simulations (Anderson and Robinson 2001).

Distance-based linear models (DISTLM) and distance-based redundancy analyses (dbRDA) were performed in PRIMER with PERMANOVA to assess the relationship between groups of predictor variables and groups of response variables (Legendre and Anderson 1999; McArdle and Anderson 2001). In this way, the variation in photophysiology (iPAM and bleaching response) was partitioned in response to benthic substrate and CO₂ level. In addition, we assessed the significance of *Symbiodinium* types present (detectable at any band intensity), *Symbiodinium* type dominance (detectable at band intensities > 30% luminescence threshold), benthic substrate and

 CO_2 level on the recruit bleaching response. Furthermore, we tested the hypothesis that CO_2 level and benthic substrate may influence *Symbiodinium* types present and/or *Symbiodinium* type dominance.

Results

Photo-physiology

The photochemical efficiency of PSII was not significantly affected by the CO_2 treatment when assessed as a multivariate (Pseudo $F_{2,228} = 0.72$; p = 0.67) or univariate response (See Supplementary Material Table S1). However, at the centre of the polyp, bleaching significantly increased more than 2.5 fold under the B2 and A1FI IPCC CO_2 scenarios in comparison to controls (Pseudo $F_{2,232} = 5.84$, p = 0.01). Similarly, bleaching of the whole recruit significantly increased by 30% under the B2 IPCC CO_2 scenario (Pseudo $F_{2,232} = 4.36$, p = 0.03) but there was no significant increase under the A1FI CO_2 scenario, although there was an increasing trend relative to the control treatment (Figure 2).

Symbiodinium

The percentage of recruits that had established endosymbiosis 23 days after settlement was not significantly affected by the CO₂ treatments, although there was a declining trend with increasing CO_2 (Pseudo $F_{2,18} = 1.48$, p = 0.25; See Supplementary Material Table S2) (Figure 3). After two months in the CO₂ treatments, Symbiodinium from clades A, C and D were detected in the recruits with 64.7% hosting multiple clades. At the clade level, the percentage of Symbiodinium present in the coral recruits was Clade A at 92.5%, Clade C at 24.8% and Clade D at 59.4%. The Symbiodinium types detected were A3, A5 and two unofficial clade A types that, for the purpose of this study were named A EF455526 and A AF333508 according to their Genbank accession numbers (EF455526 and AF333508 respectively), C3, D1 and D1-4 (provisionally named Symbiodinium trenchi and also known as D1a) (LaJeunesse et al. 2010a; LaJeunesse et al. 2010b). Due to a high degree of sequence homology in A5 and A_AF333508 these types were pooled and as type D1-4 was only present in two recruits we pooled this data with type D1 for subsequent analyses. The recruits may have taken up the cultured Symbiodinium type D1 but did not take up types A2, C_AF360576 (Genbank accession number AF360576) or C1 (within the detection limits of DGGE). There were no significant differences in the *Symbiodinium* types present in the recruits (Pseudo $F_{2,103} = 0.51$, p = 0.83) and there was no change in the dominant *Symbiodinium* type due to CO_2 treatment (Pseudo $F_{2,103} = 1.58$, p = 0.15) although there were phenotypic differences of the types within treatment (See Combined Effects).

Benthic Substrate

 CO_2 treatment significantly affected the composition of benthic functional groups around the edge of the recruits (Pseudo $F_{2,232} = 3.27$, p = 0.02; See Supplementary Material Table S3). Upon further analysis of the individual functional groups, we found that this response was characterised by a 33% decline in CCA at the edge of the recruits in the A1FI CO_2 scenario in comparison to the controls (Pseudo $F_{2,232} = 2.99$, p = 0.07) (Figure 4).

Combined Effects

Together, the benthic functional group surrounding the recruits and CO₂ treatments explained 10% of the variation in photo-physiology data (iPAM and bleaching response) (Figure 5). The dbRDA ordination clearly shows separation of the control group from the IPCC B2 and A1FI CO₂ treatments. When assessed separately in the marginal tests of the DISTLM analysis, the benthic functional groups bare tile, CCA, encrusting fleshy algae, and turf + endolithic algae along with the CO₂ treatments significantly explained the photo-physiological response (p < 0.05; See Supplementary Material Table S4). The vectors on the dbRDA ordination indicate that a similar photo-physiological response was found among recruits surrounded by CCA and encrusting fleshy algae or maintained under control conditions. In contrast, recruits in the IPCC B2 or A1FI CO₂ treatments displayed a similar photo-physiological response as recruits surrounded by bare tile or turf + endolithic algae, respectively (Figure 5).

Overall, the *Symbiodinium* type ('presence' indicating the type was detectable at any band intensity and 'dominance' indicating detectable at band intensities > 30% luminescence threshold), benthic substrate and CO₂ treatments explained 25% of the variation in the bleaching response (Figure 6). The marginal tests of the DISTLM analysis indicated that statistically significant predictors of the bleaching response

were *Symbiodinium* type D1 + D1-4 dominance along with the A1FI CO₂ scenario and Control conditions (p < 0.05; See Supplementary Material Table S5). On average, recruits that were dominant with *Symbiodinium* types D1 + D1-4 were 56% (n = 22) and 47% (n = 18) more bleached at the centre of the polyp in the B2 and A1FI CO₂ treatments, respectively. Similarly, the D1 + D1-4 dominant whole recruits were 28% and 17% more bleached in the B2 and A1FI CO₂ treatments, respectively. *Symbiodinium* type C3 presence also explained a relatively high proportion of the variation in the bleaching response but was not statistically significant (Pseudo F = 2.67, p = 0.10). An additional DISTLM analysis indicated that the substrate and CO₂ levels were not significant predictors of the uptake of *Symbiodinium* types (presence and dominance) (See Supplementary Material Figure S1 and Table S6).

Discussion

In this study, we set out to investigate the potential for OA to exhibit selective pressure on Acropora millepora through establishment of endosymbiosis with Symbiodinium during early ontogeny. As a result of increasing CO₂, the recruits were bleached at the centre of the polyp > 2.5 fold relative to controls, which may represent changes in light-harvesting pigment density or changes in Symbiodinium population density. Similarly, the whole recruit was 30% more bleached under IPCC B2 pCO₂ conditions (pH 7.80 ± 0.10) in comparison to the controls yet not under the IPCC A1FI pCO₂ conditions (pH 7.61 \pm 0.10), which may be related to the benthic functional groups surrounding the recruits (see below). Recruits containing a dominant population of Symbiodinium type D1 or D1-4 were more bleached than recruits containing Symbiodinium type C3 and visual bleaching increased under the CO₂ treatments. Interestingly, these results are the opposite of the effect of increased SST on adult colonies harbouring the more thermo-tolerant D1 versus the bleaching susceptible C3 (Jones et al. 2008). Together, these findings indicate that OA does exhibit a phenotypic selective pressure on A. millepora recruits, which may have implications for post-settlement survivorship and population dynamics. In addition, the benthic functional groups surrounding the recruits were significant in explaining the coral photo-physiological response with certain substrata (turf + endolithic algae) correlating with the response of recruits under the A1FI pCO₂ conditions while other substrata (live CCA) correlated with control conditions. Unfortunately, under the A1FI pCO₂ conditions, the percentage of live CCA surrounding the edge of the recruits significantly declined. These findings highlight the importance of benthic functional groups, which could potentially exacerbate OA conditions in the DBL of the coral recruits.

The phenomenon of coral bleaching, caused by the loss of Symbiodinium cells or degradation of their photosynthetic pigments, is generally associated with changes in sea surface temperature (SST) and light intensity (Hoegh-Guldberg 1999). More recently, OA has been shown to contribute to the bleaching response under high irradiance in adult reef-building corals, although the change in luminescence may be due to changes in either pigment or cell density (Anthony et al. 2008). Similarly, the bleaching response observed in the present study may be a result of changes in lightharvesting pigments such as chlorophyll a (chl a) and chlorophyll c2, or changes in pigments associated with thermal dissipation such as the xanthophyll cycle pigments, diadinoxanthin and diatoxanthin (Demmig-Adams and Adams 1996). Yet, in a study conducted at sub-saturating light intensities, the chl a and the xanthophyll pool per symbiont cell in Acropora formosa increased in response to increasing CO₂ (Crawley et al. 2010). This was presumably due to an increased capacity of the symbiont to process products of the photosynthetic "light-reactions" through the carboxylation of Rubsico and decreased photoprotection through photorespiration (Crawley et al. 2010). Similarly, as the present study was conducted under low irradiance, unlike Anthony et al. (2008), the bleaching response is more likely to be due to decreased symbiont cell density, rather than pigments. A decrease in Symbiodinium cells may have significant implications for the growth rate of A. millepora as there could be less carbon transferred to the host, although this depends on the initial symbiont density and whether the coral can continue to operate within a range of optimal energy acquisition (Hoogenboom et al. 2010). If the bleaching observed led to a shift in the dominant symbiont type, then this may have additional consequences for the energy budget. Although Symbiodinium D1 and D1-4 are generally regarded as more thermally tolerant, there are energetic trade-offs associated to harbouring these types such as reduced growth rates, with downstream effects on coral lipids, reproductive capacity (Little et al. 2004; Jones and Berkelmans 2011) and disease susceptibility (Bruno et al. 2007). In fact, Symbiodinium type D1-4 has been described as an opportunistic endosymbiont that outcompetes the optimal symbiont under acute or

chronic stressors (Stat and Gates 2011). These trade-offs could explain that whilst *Symbiodinium* D1 and D1-4 are known to perform better under increased SST conditions, this advantage is not necessarily sustained under other stressful conditions, such as increased pCO₂.

Noteworthy in the context of the juvenile growth rate is the need to exceed a size threshold in order to avoid mortality from disturbances such as fish grazing (Doropoulos et al. 2012b). It has previously been demonstrated that the risk of juvenile mortality is inversely proportional to growth rate and colony size (Babcock 1985; Hughes and Jackson 1985; Gosselin and Qian 1997). The delayed uptake of *Symbiodinium*, while not statistically significant in this study (Figure 3), has previously been demonstrated in *Acropora* spp. (Suwa et al. 2010) and may contribute to reduced growth rates. While the dominance of *Symbiodinium* D1 and D1-4 may allow recruits to survive OA-induced bleaching, reduced growth rates (Cohen et al. 2009; Albright et al. 2010; Doropoulos et al. 2012b) may render those recruits vulnerable to increased mortality due to predation (Doropoulos et al. 2012b) and presumably other acute disturbances such as increased SSTs (Stat and Gates 2011).

Horizontal transmission of Symbiodinium creates a window of opportunity for new recruits to harbour diverse symbionts, yet in many cases adult colonies exhibit high specificity with strong influence from the local environment (LaJuenesse et al. 2004a, LaJuenesse et al. 2004b; Abrego et al. 2009; LaJuenesse et al. 2010a). In this study, the recruits did not take up the cultured Symbiodinium types A2, C AF360576 or C1 (above the detection limits of DGGE), as these types were not likely adapted to the local environment due to their distant geographical origin (van Oppen 2004). Adult A. millepora colonies predominantly harbour Symbiodinium ITS2 types C1, C3, D1 and more rarely types A, C3k and D1-4 across the Great Barrier Reef (Cooper et al. 2011; Tonk et al. 2013) while at Heron Island the dominant type is C3 (Tonk et al. 2013). The A. millepora juveniles in this study contained a high diversity of Symbiodinium types within Clade A and although this type was not significant in explaining the photo-physiological response, it may have been an opportunistic symbiont. A recent study has shown that A. millepora adults hosted types C1 and C3 near a volcanic seep site and at a nearby low pCO2 site, while other corals hosted type D1 signifying its potential availability in the local environment (Noonan et al. 2013). This aligns with

the results of this study indicating that type D1 may not suitable for *A. millepora* under high pCO₂ conditions and also suggests that local adaptations will play an important role under future environmental perturbations (Howells et al. 2012). Long-term research is required to verify whether environmental factors such as increased pCO₂ can lead to local adaptation or a persistent change in *Symbiodinium* types.

A number of environmental factors have been shown to drive the distribution of Symbiodinium in A. millepora across the Great Barrier Reef, including SST anomalies, mean summer SST, mud and carbonate content of the sediments and distance from coastline (Cooper et al. 2011; Tonk et al. 2013). Both studies indicate additional environmental variables are at play, accounting for the unexplained variation in the *Symbiodinium* distribution data in their analyses. Interestingly, shallow reef flats that become isolated from oceanic water on low tide exhibit large fluctuations in SST, as well as large diel oscillations in carbonate chemistry (Shaw et al. 2012). In the present study, we have shown that OA alone can exert phenotypic selective pressure on Symbiodnium spp. in hospite. Conceivably, this factor may already be acting *in situ*, potentially masked by a correlation with SST anomalies. Future studies are needed to fully characterise site carbonate chemistry in order to confirm this hypothesis. Additionally, carbonate chemistry should be included as a potential predictor in multivariate analyses that aim to assess factors driving Symbiodinium distribution and the vulnerability of reef-building corals to a Symbiodinium D1 or D1-4 community shift.

The impact of community composition on carbonate chemistry presents an additional challenge for growth and survivorship of newly settled recruits under OA conditions. In our model, the photo-physiological response of recruits surrounded by turf + endolithic algae was significantly different from recruits surrounded by bare tile, as exhibited by the polarisation of these vectors in the dbRDA (Figure 5). It is not surprising that this axis only explained a small proportion of the total variation as there were no statistical differences in the fluorescence parameters and stochastic variation in recruit survival may be high during post-settlement phases (Gosselin and Qian 1997; Vermeij and Sandin 2008). Over time, however, the community composition of settlement substrate may have ecologically significant implications on the growth and survival of coral juveniles (Gosselin and Qian 1997). Interestingly,

turf + endolithic algae aligns with the highest pCO₂ treatment, the IPCC A1FI scenario, implying that this substrate produces a similar photo-physiological response. Indeed, the diel photosynthesis and respiration cycles could exacerbate the pCO₂ conditions creating even lower pH conditions at night but increasing the pH during the day (Shaw et al. 2013). A recent short-term study has demonstrated oscillating carbonate chemistry, as opposed to stable conditions, can improve growth and survivorship of recruits of the brooding coral, *Seriatopora caliendrum* presumably due to daytime relief from low pH (Dufault et al. 2012). The incidence of turf + endolithic algae surrounding recruits in the IPCC A1FI pCO₂ scenario may have improved the daytime pCO₂ conditions and may explain the slight decrease in the bleaching response for the whole recruit (Figure 2B). Dufault et al. (2012) hypothesised that nocturnal increases in pCO₂ due to respiration can accumulate and increase the intracellular dissolved inorganic carbon (DIC) pool available for calcification and photosynthesis the following day.

A. millepora recruits surrounded by live CCA demonstrated a similar photophysiological response to recruits under control conditions, as indicated by the similar direction of the vectors in the dbRDA (Figure 5). CCA provides an important settlement cue for A. millepora planulae larvae (Heyward and Negri 1999; Negri et al. 2001), which has previously been shown to perform poorly in high pCO₂ conditions (Albright et al. 2010; Doropoulos et al. 2012a). Our results align with previous studies, as after two months in the IPCC A1FI pCO₂ conditions, the percentage of live CCA surrounding the edge of the recruits significantly declined (Figure 4). The impact of OA on CCA represents an indirect mechanism by which OA may reduce A. millepora settlement. Furthermore, as coral recruits are less likely to exhibit a bleaching response when surrounded by CCA, the loss of CCA substrate due to OA may reduce the growth and survivorship of A. millepora post-settlement.

This research demonstrates that the chronic stressor, OA, causes bleaching in *A. millepora* recruits and may exhibit selective pressure through *Symbiodnium* association. Reduced symbiont density or association with particular *Symbiodinium* types may diminish growth rates thereby rendering the recruits more susceptible to acute stressors (Stat and Gates 2011; Doropoulos et al. 2012b). This will have implications for subsequent life history stages with less capacity to store lipids and

reduced reproductive output (Little et al. 2004). In addition, this study has also identified critical interactions with the benthic community, which warrant further investigation. In the field, a fine line exists between the benefits obtained from favourable manipulation of carbonate chemistry by algal photosynthesis and the potential disadvantage due to competitive interactions. Overall, the impact of OA represents a substantial threat to *A. millepora* recruitment with significant implications for the resilience of reef-building corals to both acute and chronic stressors.

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Table 1: Carbonate chemistry of the seawater in the CO_2 treatments. Total Alkalinity (A_T) and pH were measured during the experiment and are reported as mean \pm standard deviation. These values along with ambient temperature (27°C), salinity (35 PSU) and pressure (10.16 dbars) were input into the CO_2 calc program (Robbins et al. 2010) to derive the remaining parameters. We used the total hydrogen ion scale, the carbonic acid dissociation constants of Mehrbach (1973) as refit by Dickson and Millero (1987) and the sulfonic acid dissociation constant of Dickson (1990). A_T was not statistically significant between CO_2 treatments (Pseudo $F_{2,47} = 0.0005$, p = 0.999) and all other parameters were statistically distinct between CO_2 treatments (Pseudo $F_{2,47} > 43.25$, p = 0.0001).

		\mathbf{A}_{T}	$p\mathrm{CO}_2$	HCO ₃ -	CO_3^{2-}	
Treatment	pН	μmol kg ⁻¹	μatm	μmol kg ⁻¹	μmol kg ⁻¹	Ω arag
Control	8.04 ± 0.14	2291 ± 34	399 ± 169	1743 ± 162	223 ± 51	3.56 ± 0.81
B2	7.81 ± 0.10	2291 ± 33	752 ± 214	1934 ± 98	145 ± 26	2.33 ± 0.42
A1FI	7.60 ± 0.10	2290 ± 33	1297 ± 352	2056 ± 79	95 ± 18	1.53 ± 0.29

Table 2: Details of the *Symbiodinium* cultures offered to the *Acropora millepora* recruits five - eight days after settlement.

Culture		Type / Genbank		Origin of
ID	Clade	Accession no.	Isolated from Host	Culture
A001	D	D1	Acropora sp.	Okinawa
A013	D	D1	Porites annae	Okinawa
Sin	C	AF360576	Sinularia sp.	Guam
A2	A	A2	Zoanthus sociatus	Jamaica
C1	C	C1	Diccospina sanctithomae	Jamaica

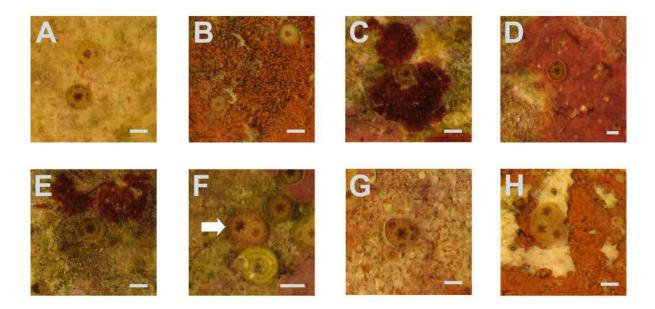


Figure 1: Examples of the benthic functional groups surrounding the *Acropora millepora* recruits (indicated by the arrow where necessary). Scale bar = 1.0 mm. A. turf + endolithic algae; B. bare tile (also some turf + endolithic algae); C. encrusting fleshy algae (also some dead crustose coralline algae [CCA]); D. live CCA; E. turf + endolithic algae (also some encrusting fleshy algae); F. shell and coral (also some turf + endolithic algae); G. dead CCA; H. dead CCA and bare tile.

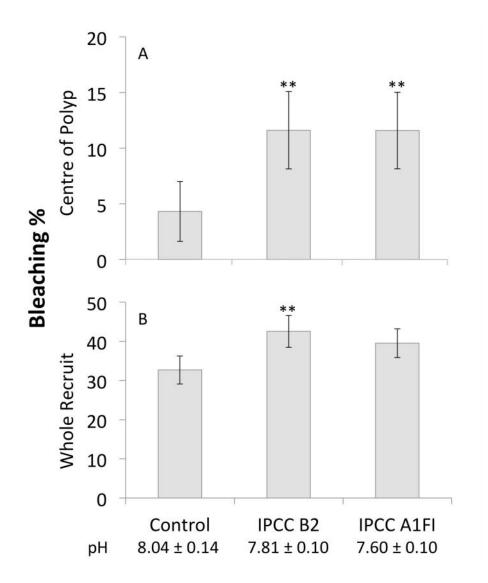


Figure 2: Bleaching % of *Acropora millepora* assessed A. at the centre of the polyp and B. for the whole recruit. Bleaching was quantified as a reduction in luminescence relative to the maximum, signifying symbiont or chlorophyll density, following the method of Anthony et al. (2008). Error bars indicate standard error (n = 15) and ** denotes significantly different to the control (p < 0.01).

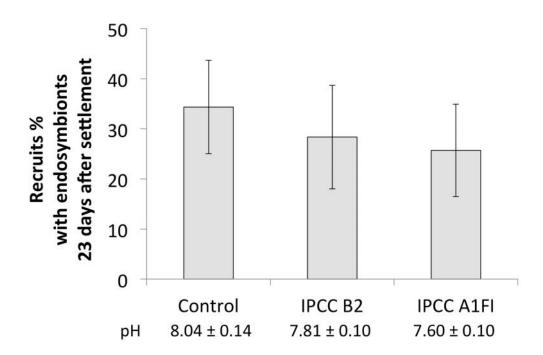


Figure 3: The percentage of *Acropora millepora* recruits that had established endosymbiosis with *Symbiodinium* 23 days after settlement. Error bars indicate standard deviation (n = 7).

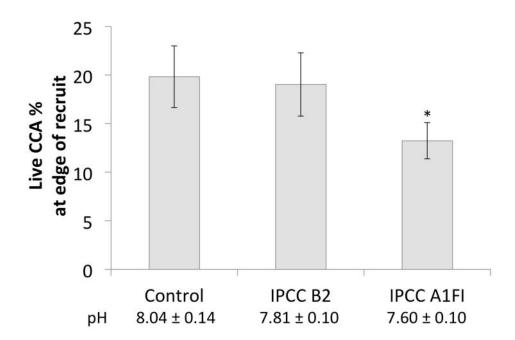


Figure 4: Percentage of live crustose coralline algae (CCA) at the edge of the *Acropora millepora* recruits after 2 months in IPCC CO_2 treatments. Error bars indicate standard error (n = 15) and * denotes significantly different from the control (p < 0.05).

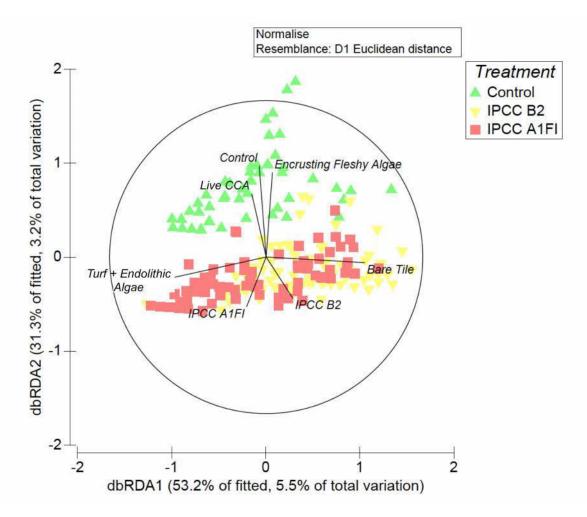


Figure 5: Distance-based redundancy analysis (dbRDA) showing ordination of the fitted values from the model of variation in photo-physiological response variables (dark-adapted Fv/Fm, at 110 μmol quanta m⁻² s⁻¹: Φ_{NPQ} , Φ_{NO} , $\Delta F/F_m$ ', Excitation Pressure (Q_m), Bleaching % at Centre of Polyp and Bleaching % for the Whole Recruit) explained by the benthic functional groups and CO₂ treatments. Vectors are predictor variables explaining a statistically significant proportion of the response (p < 0.05), as assessed by the marginal tests of distance-based linear models (DISTLM) (See Supplementary Material Table S4 for Full Statistics).

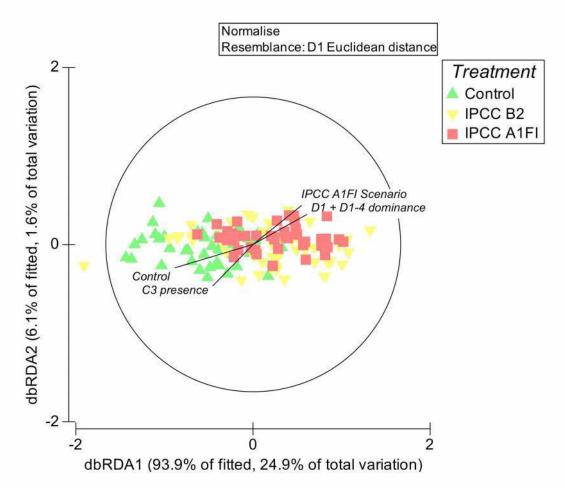


Figure 6: Distance-based redundancy analysis (dbRDA) showing ordination of the fitted values from the model of variation in bleaching response variables (Bleaching % at Centre of Polyp and Bleaching % for the Whole Recruit) explained by the *Symbiodinium* type (presence and dominance), benthic functional groups and CO_2 treatments. Vectors are predictor variables explaining a large proportion of the response (p < 0.10), as assessed by the marginal tests of distance-based linear models (DISTLM) (See Supplementary Material Table S5 for Full Statistics).

Table S1: PERMANOVA results of the photochemical efficiency of Photosystem II in *Symbiodinium* within *Acropora millepora* recruits after two months in CO₂ treatments. All models are univariate except the final model, iPAM multivariate, which includes all variables.

Variables	Source of Variation	df	Pseudo F	P (Permutation)
Fv/Fm	Treatment	2	0.87947	0.4379
dark-adapted	Tank (Treatment)	17	2.4578	0.0152
	Tile (Tank(Treatment))	38	21.145	0.0001
Δ F/Fm'	Treatment	2	0.31794	0.7383
110 µmol quanta s ⁻¹	Tank (Treatment)	17	1.2415	0.2811
	Tile (Tank(Treatment))	38	7.2193	0.0001
Q_{m}	Treatment	2	0.20475	0.8134
1 - (ΔF/Fm') · (Fv/Fm) ⁻¹	Tank (Treatment)	17	1.0497	0.4253
	Tile (Tank(Treatment))	38	9.8168	0.0001
Fv/Fm	Treatment	2	0.015579	0.986
dark recovery	Tank (Treatment)	17	0.80958	0.6851
	Tile (Tank(Treatment))	38	5.4629	0.0001
Φ_{NPQ}	Treatment	2	0.86704	0.4328
110 µmol quanta s ⁻¹	Tank (Treatment)	17	1.0805	0.4009
	Tile (Tank(Treatment))	38	2.5335	0.0001
Φ_{NPQ}	Treatment	2	1.8009	0.1936
dark recovery	Tank (Treatment)	17	0.55684	0.9077
	Tile (Tank(Treatment))	38	3.1206	0.0001
Φ_{NO} minimum	Treatment	2	0.888	0.4376
dark-adapted	Tank (Treatment)	17	2.3921	0.0178
	Tile (Tank(Treatment))	38	21.575	0.0001
Φ_{NO}	Treatment	2	1.2185	0.3196
110 μmol quanta s ⁻¹	Tank (Treatment)	17	1.3991	0.187
	Tile (Tank(Treatment))	38	3.9426	0.0001
Φ_{NO}	Treatment	2	0.43264	0.6732
dark recovery	Tank (Treatment)	17	0.73565	0.7689
	Tile (Tank(Treatment))	38	7.3319	0.0001
iPAM multivariate	Treatment	2	0.72297	0.6717
(all of the above)	Tank (Treatment)	17	1.2935	0.0992
	Tile (Tank(Treatment))	38	6.1931	0.0001

Table S2: PERMANOVA results for the percentage of *Symbiodinium* uptake 23 days after settlement and the *Symbiodinium* type (presence and dominance) within *Acropora millepora* recruits after two months in CO₂ treatments.

Variables	Source of Variation	df	Pseudo F	P (Permutation)
Symbiodinium uptake	Treatment	2	1.4762	0.2536
23 days after settlement		18		
Symbiodinium type	Treatment	2	0.51402	0.8308
presence	Tank (Treatment)	14	0.8878	0.645
	Tile (Tank(Treatment))	13	1.6868	0.0039
Symbiodinium type	Treatment	2	1.5766	0.1466
dominance	Tank (Treatment)	14	1.4751	0.1229
	Tile (Tank(Treatment))	13	1.0946	0.3056
Symbiodinium type	Treatment	2	0.90752	0.5122
presence + dominance	Tank (Treatment)	14	1.1207	0.3465
	Tile (Tank(Treatment))	13	1.4189	0.0256

Table S3: PERMANOVA results for the benthic functional groups surrounding the edge of the *Acropora millepora* recruits after two months in CO₂ treatments.

Variables	Source of Variation	df	Pseudo F	P (Permutation)	Pairwise P (Monte Carlo)
Benthic Substrate	Treatment	2	3.272	0.02	Control ≠ B2; p = 0.0035
	Tank (Treatment)	17	1.4668	0.0705	Control & A1FI; p = 0.0664
	Tile (Tank(Treatment))	38	3.6818	0.0001	B2 & A1FI; p = 0.2145
Live Crustose	Treatment	2	2.993	0.0703	
Coralline Algae	Tank (Treatment)	17	1.1144	0.376	
	Tile (Tank(Treatment))	38	2.7048	0.001	
Bare Tile	Treatment	2	4.4939	0.0196	Control < B2; $p = 0.0017$
	Tank (Treatment)	17	1.5412	0.1283	Control & A1FI; p = 0.0545
	Tile (Tank(Treatment))	38	5.3557	0.0001	B2 & A1FI; p = 0.426
Turf +	Treatment	2	1.5623	0.2296	
Endolithic Algae	Tank (Treatment)	17	1.9667	0.0267	
	Tile (Tank(Treatment))	38	2.8185	0.0001	
Encrusting Fleshy	Treatment	2	0.27728	0.7962	
Algae	Tank (Treatment)	17	0.77459	0.7266	
	Tile (Tank(Treatment))	38	1.3165	0.0986	
Shell	Treatment	2	2.4835	0.1033	
	Tank (Treatment)	17	0.69438	0.8044	
	Tile (Tank(Treatment))	38	1.9882	0.0012	
Coral	Treatment	2	2.4835	0.1033	
	Tank (Treatment)	17	0.69438	0.8044	
	Tile (Tank(Treatment))	38	1.9882	0.0012	
Dead Crustose	Treatment	2	0.51146	0.6097	
Coralline Algae	Tank (Treatment)	17	1.2942	0.2519	
	Tile (Tank(Treatment))	38	2.0434	0.0005	

Table S4: Distance-based linear model (DISTLM) marginal tests indicating the proportion of photo-physiology variation (dark-adapted Fv/Fm, at 110 µmol quanta m⁻² s⁻¹: Φ_{NPQ} , Φ_{NO} , $\Delta F/F_m$ ', Excitation Pressure (Q_m), Bleaching % at Centre of Polyp and Bleaching % for the Whole Recruit) explained by the benthic functional groups and CO_2 treatments.

Predictor Variables	Pseudo F	P	Proportion of variation explained
Bare Tile	11.983	0.0001	0.0406
Shell	0.97681	0.4066	0.0034
Coral	0.36075	0.8321	0.0013
Live CCA	2.7931	0.0283	0.0098
Dead CCA	1.5021	0.2014	0.0053
Encrusting Fleshy Algae	2.4852	0.0457	0.0087
Turf + Endolithic Algae	7.6351	0.0001	0.0263
Control	7.5963	0.0001	0.0261
IPCC B2	4.552	0.0017	0.0158
IPCC A1FI	3.9961	0.0048	0.0139

Table S5: Distance-based linear model (DISTLM) marginal tests indicating the proportion of bleaching variation (Bleaching % at Centre of Polyp and Bleaching % for the Whole Recruit) explained by the *Symbiodinium* type (presence and dominance), benthic functional groups and CO₂ treatments.

			Proportion of variation
Predictor Variables	Pseudo F	P	explained
A3 presence	0.061096	0.8851	0.0005
A5 + A_AF333508 presence	0.55502	0.4750	0.0042
A_EF455526 presence	0.36332	0.5694	0.0028
C3 presence	2.6674	0.0986	0.0200
D1 + D1-4 presence	0.55423	0.4736	0.0042
A3 dominance	2.2815	0.1266	0.0171
A5 + A_AF333508 dominance	0.23219	0.6741	0.0018
A_EF455526 dominance	1.9345	0.1553	0.0146
C3 dominance	1.2674	0.2567	0.0096
D1 + D1-4 dominance	4.5884	0.0313	0.0338
Bare Tile	1.6106	0.2020	0.0121
Shell	0.81592	0.3745	0.0062
Coral	2.488	0.1086	0.0186
Live CCA	0.91213	0.3449	0.0069
Dead CCA	1.47	0.2232	0.0111
Encrusting Fleshy Algae	0.33804	0.5962	0.0026
Turf + Endolithic Algae	2.6126	0.1050	0.0196
Control	10.042	0.0028	0.0712
IPCC B2	1.2166	0.2681	0.0092
IPCC A1FI	3.7872	0.0495	0.0281

Table S6: Distance-based linear model (DISTLM) marginal tests indicating the proportion of *Symbiodinium* type (presence and dominance) variation explained by the benthic functional groups and CO₂ treatments.

Predictor Variables	Pseudo F	P	Proportion of variation explained
Bare Tile	1.4541	0.2065	0.0110
Shell	1.1028	0.3493	0.0083
Coral	2.8031	0.0212	0.0210
Live CCA	0.62605	0.6810	0.0048
Dead CCA	1.5633	0.1715	0.0118
Encrusting Fleshy Algae	0.88448	0.4731	0.0067
Turf + Endolithic Algae	1.4721	0.1961	0.0111
Control	0.97031	0.4256	0.0074
IPCC B2	1.1851	0.3119	0.0090
IPCC A1FI	0.62431	0.6784	0.0047

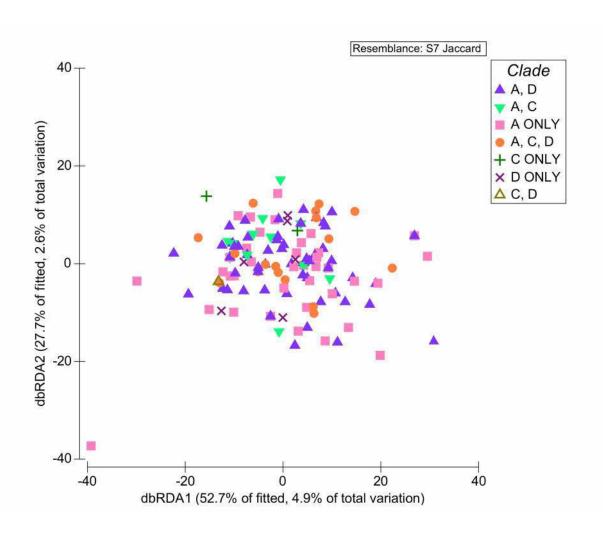


Figure S1: Distance based redundancy analysis (dbRDA) showing ordination of the fitted values from the model of variation in *Symbiodinium* type (presence and dominance) explained by the benthic functional groups and CO₂ treatments.

References

- Abrego D, van Oppen MJ, Willis BL (2009) Highly infectious symbiont dominates initial uptake in coral juveniles. Molecular Ecology 18: 3518-3531
- Albright R, Langdon C (2011) Ocean acidification impacts multiple early life history processes of the Caribbean coral *Porites asteroides*. Global Change Biology doi: 10.1111/j.1365-2486.2011.02404.x
- Albright R, Mason B, Miller M, Langdon C (2010) Ocean acidification compromises recruitment success of the threatened Caribbean coral *Acropora palmata*. Proceedings of the National Academy of Science 107: 20400 20404
- Anderson MJ (2001) A new method for non-parametric multivariate analysis of variance. Austral Ecology 26: 32-46
- Anderson MJ, Gorley RN, Clarke KR (2008) PERMANOVA+ for PRIMER: Guide to software and statistical methods. PRIMER-E Ltd, Plymouth, UK
- Anderson MJ, Robinson J (2001) Permutation tests for linear models. Australian and New Zealand Journal of Statistics 43: 75-88
- Anthony KRN, Connolly SR, Hoegh-Guldberg O (2007) Bleaching, energetics, and coral mortality risk: Effects of temperature, light, and sediment regime. Limnology and Oceanography 52: 716-726
- Anthony KRN, Diaz-Pulido G, Verlinden N, Tilbrook B, Andersson AJ (2013)

 Benthic buffers and boosters of ocean acidification on coral reefs.

 Biogeosciences 10: 4897-4909
- Anthony KRN, Kleypas JA, Gattuso J-P (2011) Coral reefs modify their seawater carbon chemistry implications for impacts of ocean acidification. Global Change Biology 17: 3655-3666
- Anthony KRN, Kline DI, Diaz-Pulido G, Dove S, Hoegh-Guldberg O (2008) Ocean acidification causes bleaching and productivity loss in coral reef builders. Proceedings of the National Academy of Science 105: 17442-17446
- Babcock R, Mundy C (1996) Coral recruitment: consequences of settlement choice for early growth and survivorship in two scleractinians. Journal of Experimental Marine Biology and Ecology 206: 179-201
- Babcock RC (1985) Growth and mortality in juvenile corals (Goniastrea, Platygyra and Acropora): the first year Proceedings of the Fifth International Coral Reef Congress, Tahiti, French Polynesia, pp 355-360

- Babcock RC, Heyward AJ (1986) Larval development of certain gamete-spawning scleractinian corals. Coral Reefs 5: 111-116
- Baird AH, Hughes T (2000) Competitive dominance by tabnular corals: an experimental analysis of recruitment and survival of understory assemblages.

 Journal of Experimental Marine Biology and Ecology 251: 117-132
- Berkelmans R, van Oppen MJH (2006) The role of zooxanthellae in the thermal tolerance of corals: a 'nugget of hope' for coral reefs in an era of climate change. Proceedings of the Royal Society B 273: 2305-2312
- Bruno JF, Selig ER, Casey KS, Page CA, Willis BL, Harvell CD, Sweatman H, Melendy AM (2007) Thermal stress and coral cover as drivers of coral disease outbreaks. PLOS Biology 5: e124
- Caldeira K, Wickett ME (2003) Anthropogenic carbon and ocean pH. Nature 425: 365
- Canadell JG, Quere CL, Raupacha MR, Fielde CB, Buitenhuisc ET, Ciaisf P, Conwayg TJ, Gillettc NP, Houghtonh RA, Marlandi G (2007) Contributions to accelerating atmospheric CO₂ growth from economic activity, carbon intensity, and efficiency of natural sinks. Proceedings of the National Academy of Science 104: 18866-18870
- Coffroth MA, Lewis CF, Santos SR (2006) Environmental populations of symbiotic dinoflagellates in the genus *Symbiodinium* can initiate symbioses with reef cnidarians. Current Biology 16: 987-987
- Cohen AL, McCorkle DC, dePutron S, Gaetani GA, Rose KA (2009) Morphological and compositional changes in the skeletons of new coral recruits reared in acidified seawater: Insights into the biomineralization response to ocean acidification. Geochemistry, Geophysics, Geosystems 10: Q07005
- Cooper TF, Berkelmans R, Ulstrup KE, Weeks S, Radford B, Jones AM, Doyle J, Canto M, O'Leary RA, van Oppen MJH (2011) Environmental factors controlling the distribution of *Symbiodinium* harboured by the coral *Acropora millepora* on the Great Barrier Reef. PLoS One 6
- CoralWatch (2001) Coral Health Chart. Queensland Brain Institute, The University of Queensland, St Lucia
- Crawley A, Kline DI, Dunn S, Anthony K, Dove S (2010) The effect of ocean acidification on symbiont photorespiration and productivity in *Acropora formosa*. Global Change Biology 16: 851-863

- Cumbo VR, Baird AH, van Oppen MJH (2013) The promiscuous larvae: flexibility in the establishment of symbiosis in corals. Coral Reefs 32: 111-120
- Demmig-Adams B, Adams WI (1996) The role of xanthophyll cycle carotenoids in the protection of photosynthesis. Trends in Plant Science 1: 21-26
- Dickson AG (1990) Standard potential of the reaction AgCl(s) + .5H2(g) = Ag(s) + HCl(aq) and the standard acidity constant of the ion HSO4– in synthetic sea water from 273.15 to 318.15 K. The Journal of Chemical Thermodynamics 22: 113-127
- Dickson AG, Millero F (1987) A comparison of the equilibrium-constants for the dissociation of carbonic-acid in seawater media Deep-Sea Research Part A Oceanographic Research Papers 34: 1733-1743
- Dickson AG, Sabine CL, Christian J (2007) Guide to best practices for ocean CO₂ measurements. PICES Special Publication 3, pp 191
- Doropoulos C, Diaz-Pulido G (2013) High CO₂ reduces the settlement of a spawning coral on three common species of crustose coralline algae. Marine Ecology Progress Series 475: 93-99
- Doropoulos C, Ward S, Diaz-Pulido G, Hoegh-Guldberg O, Mumby PJ (2012a)

 Ocean acidification reduces coral recruitment by disrupting intimate larvalalgal settlement interactions. Ecology Letters 15: 338-346
- Doropoulos C, Ward S, Marshell A, Diaz-Pulido G, Mumby PJ (2012b) Interactions among chronic and acute impacts on coral recruits: the importance of size-escape thresholds. Ecology 93: 2131-2138
- Dufault AM, Cumbo VR, Fan T-Y, Edmunds PJ (2012) Effects of diurnally oscillating pCO₂ on the calcification and survival of coral recruits. Proceedings of the Royal Society B-Biological Sciences 279: 2951-2958
- Fabricius KE, Mieog JC, Colin PL, Idip D, van Oppen MJH (2004) Identity and diversity of coral endosymbionts (zooxanthellae) from three Palauan reefs with contrasting bleaching, temperature and shading histories. Molecular Ecology 13:2445-2458
- Fairfull SJL, Harriott VJ (1999) Succession, space and coral recruitment in a subtropical fouling community. Marine and Freshwater Research 50: 235-242
- Falkowski PG, Dubinsky Z (1981) Light-shade adaptation of *Stylophora pistillata*, a hermatypic coral from the Gulf of Eilat. Nature 289: 172-174

- Falkowski PG, Dubinsky Z, Muscatine L, Porter JW (1984) Light and bioenergetics of a symbiotic coral. BioScience 34: 705-709
- Finney JC, Pettay DT, Sampayo EM, Warner ME, Oxenford HA, LaJeunesse TC (2010) The relative significance of host-habitat, depth, and geography on the ecology, endemism, and speciation of coral endosymbionts in the genus *Symbiodinium*. Microbial Ecology 60: 250-263
- Frade PR, Bongaerts P, Winkelhagen AJS, Tonk L, Bak RPM (2008) In situ photobiology of corals over large depth ranges: A multivariate analysis on the roles of environment, host, and algal symbiont. Limnology and Oceanography 53: 2711-2723
- Gómez-Cabrera MdC, Ortiz JC, Loh WKW, Ward S, Hoegh-Guldberg O (2008) Acquisition of symbiotic dinofagellates (*Symbiodinium*) by juveniles of the coral *Acropora longicyathus*. Coral Reefs 27: 219-226
- Gosselin LA, Qian PY (1997) Juvenile mortality in benthic marine invertebrates.

 Marine Ecology Progress Series 146: 265-282
- Gran G (1952) Determination of the equivalence point in potentiometric titrations:

 Part II. Analyst 77: 661-670
- Heyward AJ, Negri AP (1999) Natural inducers for coral larval metamorphosis. Coral Reefs 18: 273-279
- Hill R, Ralph PJ (2008) Impact of bleaching stress on the function of the oxygen evolving complex of zooxanthellae from scleractinian corals. Journal of Phycology 44: 299-310
- Hoegh-Guldberg O (1999) Climate change, coral bleaching and the future of the world's coral reefs. Marine and Freshwater Research 50: 839-866
- Hoegh-Guldberg O, Smith G (1989) The effect of sudden changes in temperature, light and salinity on the population density and export of zooxanthellae from the reef corals *Stylophora pistillata* Esper and *Seriatopora hystrix* Dana. Journal of Experimental Marine Biology and Ecology 129: 279-303
- Hoogenboom M, Beraud E, Ferrier-Pages C (2010) Relationship between symbiont density and photosynthetic carbon acquisition in the temperate coral *Cladocora caespitosa*. Coral Reefs 29: 21-29
- Howells EJ, Beltran VH, Larsen NW, Bay LK, Willis BL, van Oppen MJH (2012) Coral thermal tolerance shaped by local adaptation of photosymbionts. Nature Climate Change 2: 116-120

- Hughes TP, Jackson JBC (1985) Population Dynamics and Life Histories of Foliaceous Corals. Ecological Monographs 55: 142-166
- Hurd CL, Cornwall CE, Currie K, Hepburn CD, McGraw CM, Hunter KA, Boyd PW (2011) Metabolically induced pH fluctuations by some coastal calcifiers exceed projected 22nd century ocean acidification: a mechanism for differential susceptibility? Global Change Biology 17: 3254-3262
- Iglesias-Prieto R, Beltran V, LaJeunesse T, Reyes-Bonilla H, Thome P (2004)

 Different algal symbionts explain the vertical distribution of dominant reef corals in the eastern Pacific. Proceedings of the Royal Society of London. B 271: 1757-1763
- Jokiel PL (2011) Ocean Acidifcation and control of reef coral calcification by boundary layer limitation of proton flux. Bulletin of Marine Science 87: 639–657
- Jokiel PL, Coles SL (1990) Response of Hawaiian and other Indo-Pacific reef corals to elevated temperature. Coral Reefs 8: 155-162
- Jones AM, Berkelmans R (2011) Tradeoffs to thermal acclimation: energetics and reproduction of a reef coral with heat tolerant *Symbiodinium* Type-D. Journal of Marine Biology 2011: 185890
- Jones AM, Berkelmans R, van Oppen MJH, Mieog JC, Sinclair W (2008) A community change in the algal endosymbionts of a scleractinian coral following a natural bleaching event: field evidence of acclimatization. Proceedings of the Royal Society of London: B 275: 1359-1365
- Kleypas JA, Anthony KRN, Gattuso J-P (2011) Coral reefs modify their seawater carbon chemistry case study from a barrier reef (Moorea, French Polynesia). Global Change Biology 17: 3667-3678
- Kleypas JA, Feely RA, Fabry VJ, Langdon C, Sabine CL, Robbins LL (2006) Impacts of Ocean Acidification on Coral Reefs and Other Marine Calcifiers: A Guide for Future Research Workshop held 18–20 April 2005, sponsored by NSF, NOAA, and the U.S. Geological Survey, St. Petersburg, FL, pp 88
- Kroeker KJ, Kordas RL, Crim RN, Singh GG (2010) Meta-analysis reveals negative yet variable effects of ocean acidification on marine organisms. Ecological Letters 13: 1419-1434
- Kump LR, Bralower TJ, Ridgwell A (2009) Ocean acidification in deep time.

 Oceanography 22: 94-107

- LaJeunesse TC (2001) Investigating the biodiversity, ecology, and phylogeny of endosymbiotic dinoflagellates in the genus *Symbiodinium* using the its region: In search of a "species" level marker. Journal of Phycology 37: 866-880
- LaJeunesse TC, Bhagooli R, Hidaka M, DeVantier L, Done T, Schmidt GW, Fitt WK, Hoegh-Guldberg O (2004a) Closely related *Symbiodinium* spp. differ in relative dominance in coral reef host communities across environmental, latitudinal and biogeographic gradients. Marine Ecology-Progress Series 284: 147-161
- LaJeunesse TC, Bonilla HR, Warner ME, Wills M, Schmidt GW, Fitt WK (2008) Specificity and stability in high latitude eastern Pacific coral-algal symbioses. Limnology and Oceanography 53: 719-727
- LaJeunesse TC, Lambert G, Andersen RA, Coffroth MA, Galbraith DW (2005)

 Symbiodinium (Pyrrhophyta) genome sizes (DNA content) are smallest among dinoflagellates. Journal of Phycology 41: 880-886
- LaJeunesse TC, Loh WKW, Woesik Rv, Hoegh-Guldberg O, Schmidt GW, Fitt WK (2003) Low symbiont diversity in southern Great Barrier Reef corals, relative to those of the Caribbean. Limnology & Oceanography 48: 2046-2054
- LaJeunesse TC, Pettay DT, Sampayo EM, Phongsuwan N, Brown B, Obura DO, Hoegh-Guldberg O, Fitt WK (2010a) Long-standing environmental conditions, geographic isolation and host-symbiont specificity influence the relative ecological dominance and genetic diversification of coral endosymbionts in the genus *Symbiodinium*. Journal of Biogeography 37:785-800
- LaJeunesse TC, Smith RT, Walther M, Pinzon J, Pettay DT, McGinley M, Aschaffenburg M, Medina-Rosas P, Cupul-Magana AL, Perez AL, Reyes-Bonilla H, Warner ME (2010b) Host-symbiont recombination versus natural selection in the response of coral-dinoflagellate symbioses to environmental disturbance. Proceedings of the Royal Society of London B 277: 2925-2934
- LaJeunesse TC, Thornhill DJ, Cox EF, Stanton FG, Fitt WK, Schmidt GW (2004b)

 High diversity and host specificity observed among symbiotic dinoflagellates in reef coral communities from Hawaii. Coral Reefs 23: 596-603
- Legendre P, Anderson MJ (1999) Distance-based redundancy analysis: testing multispecies responses in multifactorial ecological experiments. Ecological Monographs 75: 435-450

- Little AF, van Oppen MJH, Willis BL (2004) Flexibility in algal endosymbiosies shapes growth in reef corals. Science 304: 1492-1494
- Loram JE, Trapido-Rosenthal HG, Douglas AE (2007) Functional significance of genetically different symbiotic algae *Symbiodinium* in a coral reef symbiosis. Molecular Ecology 16: 4849-4857
- Marfenin NN (1997) Adaptation capabilities of marine modular organisms. Hydrobiologia 355: 153-158
- Marubini F, Ferrier-Pages C, Furla P, Allemand D (2008) Coral calcification responds to seawater acidification: a working hypothesis towards a physiological mechanism. Coral Reefs DOI 10.1007/s00338-008-0375-6
- McArdle BH, Anderson MJ (2001) Fitting multivariate models to community data: a comment on distance-based redundancy analysis. Ecology 82: 290-297
- Mehrbach C, Culberso CH, Hawley JE, Pytkowic RM (1973) Measurement of apparent dissociation-constants of carbonic acid in seawater at atmospheric-pressure. Limnology and Oceanography 1973: 897-901
- Negri AP, Webster NS, Hill RT, Heyward AJ (2001) Metamorphosis of broadcast spawning corals in response to bacteria isolated from crustose algae. Marine Ecology Progress Series 223: 121-131
- Noonan SHC, Fabricius KE, Humphrey C (2013) *Symbiodinium* community composition in scleractinian corals is not affected by life-long exposure to elevated carbon dioxide. PLoS One 8: e63985
- Pandolfi JM, Connolly SR, Marshall DJ, Cohen AL (2011) Projecting coral reef futures under global warming and ocean acidification. Science 333: 418-422
- Pochon X, Garcia-Cuetos L, Baker AC, Castella E, Pawlowski J (2007) One-year survey of a single Micronesian reef reveals extraordinarily rich diversity of *Symbiodinium* types in soritid foraminifera. Coral Reefs 26: 867-882
- Pochon X, Gates RD (2010) A new *Symbiodinium* clade (Dinophyceae) from soritid foraminifera in Hawaii. Molecular Phylogenetics and Evolution 56: 492-497
- Porto I, Granados C, Restrepo JC, Sanchez JA (2008) Macroalgal-associated dinoflagellates belonging to the genus *Symbiodinium* in Caribbean reefs. PLoS One 3: e2160
- Robbins LL, Hansen ME, Kleypas JA, Meylan SC (2010) CO2calc—A user-friendly seawater carbon calculator for Windows, Max OS X, and iOS (iPhone). U.S. Geological Survey, Florida

- Rowan R, Knowlton N (1995) Intraspecific diversity and ecological zonation in coral algal symbiosis. Proceedings of the National Academy of Science 92: 2850-2853
- Rowan R, Powers DA (1991) A molecular genetic classification of zooxanthellae and the evolution of animal-algal symbioses. Science 251: 1348-1351
- Sabine CL, Feely RA, Gruber N, Key RM, Lee K, Bullister JL, Wanninkhof R, Wong CS, Wallace DWR, Tilbrook B, Millero FJ, Peng T-H, Kozyr A, Ono T, Rios AF (2004) The oceanic sink for anthropogenic CO₂. Science 305: 367-371
- Sampayo EM, Francheschinis L, Hoegh-Guldberg O, Dove S (2007) Niche partitioning of closely related symbiotic dinoflagellates. Molecular Ecology 16: 3721-3733
- Sampayo EM, Ridgway T, Bongaerts P, Hoegh-Guldberg O (2008) Bleaching susceptibility and mortality of corals are determined by fine-scale differences in symbiont type. Proceedings of the National Academy of Sciences of the United States of America 105: 10444-10449
- Schneider K, Erez J (2006) The effect of carbonate chemistry on calcification and photosynthesis in the hermatypic coral *Acropora eurystoma*. Limnology and Oceanography 51: 1284-1293
- Seutin G, White BN, Boag PT (1991) Preservation of avian blood and tissue samples for DNA analyses. Canadian Journal of Zoology 69: 82-90
- Shashar N, Kinane S, Jokiel PL, Patterson MR (1996) Hydromechanical boundary layers over a coral reef. Journal of Experimental Marine Biology and Ecology 199: 17-28
- Shaw EC, McNeil BI, Tilbrook B (2012) Impacts of ocean acidification in naturally variable coral reef flat ecosystems. Journal of Geophysical Research 117: C03038
- Shaw EC, McNeil BI, Tilbrook B, Matear R, Bates ML (2013) Anthropogenic changes to seawater buffer capacity combined with natural reef metabolism induce extreme future coral reef CO₂ conditions. Global Change Biology 19: 1632-1641
- Stat M, Gates RD (2011) Clade D *Symbiodinium* in scleractinian corals: A "nugget" of hope, a selfish opportunist, an ominous sign, or all of the above? Journal of Marine Biology 2011: 730715

- Suwa R, Nakamura M, Morita M, Shimada K, Iguchi A, Sakai K, Suzuki A (2010) Effects of acidified seawater on early life stages of scleractinian corals (Genus Acropora). Fisheries Science 76: 93-99
- Tonk L, Sampayo EM, Weeks S, Magno-Canto M, Hoegh-Guldberg O (2013) Hostspecific interactions with environmental factors shape the distribution of *Symbiodinium* across the Great Barrier Reef. PLoS One 8: e68533
- Thornhill DJ, LaJeunesse TC, Kemp DW, Fitt WK, Schmidt GW (2006) Multi-year, seasonal genotypic surveys of coral-algal symbioses reveal prevalent stability or post-bleaching reversion. Marine Biology 148:711-722
- van Oppen MJH (2004) Mode of zooxanthella transmission does not affect zooxanthella diversity in acroporid corals. Marine Biology 144: 1-7
- Vermeij MJA, Sandin SA (2008) Density-dependent settlement and mortality structure the earliest life phases of a coral population. Ecology 89: 1994-2000
- Veron JEN (2008) A Reef In Time: The Great Barrier Reef From Beginning To End.
 Belknap Press
- Walters RG (2005) Towards an understanding of photosynthetic acclimation. Journal of Experimental Botany 56: 435-447
- Warner ME, Fitt WK, Schmidt GW (1999) Damage to photosystem II in symbiotic dinoflagellates: A determinant of coral bleaching. Proceedings of the National Academy of Science 96: 8007–8012
- Warner ME, LaJeunesse TC, Robison JD, Thur RM (2006) The ecological distribution and comparative photobiology of symbiotic dinoflagellates from reef corals in Belize: Potential implications for coral bleaching. Limnology & Oceanography 51: 1887-1897
- Webster NS, Smith LD, Heyward AJ, Watts JEM, Webb RI, Blackall LL, Negri AP (2004) Metamorphosis of a scleractinian coral in response to microbial biofilms. Applied and Environmental Microbiology 70: 1213-1221
- Wham DC, Pettay DT, LaJeunesse TC (2011) Microsatellite loci for the host-generalist "zooxanthella" *Symbiodnium trenchi* and other Clade D *Symbiodinium*. Conservation Genetics Resources 3: 541-544
- Wiedenmann J, D'Angelo C, Smith EG, Hunt AN, Legiret F-E, Postle AD, Achterberg EP (2012) Nutrient enrichment can increase the susceptibility of reef corals to bleaching. Nature Climate Change 3: 160-164

Wilson J, Harrison P (2005) Post-settlement mortality and growth of newly settled reef corals in a subtropical environment. Coral Reefs 24: 418-421

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Chapter 5: General Discussion

The survival of reef-building corals is threatened by the physiological impacts of ocean acidification (OA) operating across several life-history stages. While most of the current literature has focused on the decline in calcification, characterising the effect of OA on photo-physiological processes is essential to understanding the phenomenon of CO₂-induced bleaching and productivity loss (Anthony et al. 2008). The mechanisms of thermal coral bleaching are well understood, albeit the site of dysfunction is still debated (Jones et al. 1998; Warner et al. 2006), yet little is known about how increased pCO₂ lowers the thermal bleaching threshold. Furthermore, the ecological ramifications of the impact of OA must be considered in order to improve the predictive power of climate models on coral reef ecosystems. The resilience of adult corals to OA is fundamentally dependent on coral reef carbonate chemistry, which is shaped by the community composition and residence time of reef waters (Anthony et al. 2011; Kleypas et al. 2011; Shaw et al. 2013). Overall, coral reef communities are likely to respond to selective pressures, such as OA, with a shift towards more resilient species of the coral host (Inoue et al. 2013) or resilient species of their symbiotic dinoflagellates (genus Symbiodinium) (Jones et al. 2008). In relation thereto, coral juvenile uptake of Symbiodinium represents a life history stage amenable to adaption of the symbiosis (Baird et al. 2007; Abrego et al. 2009b; Van Oppen et al. 2011). The first objective of this thesis was to investigate the photophysiological response of reef-building corals to OA during the adult phase and at the onset of symbiosis during the juvenile phase. The second objective was to address these physiological changes in an ecological setting by assessing the coral response to future OA treatments in relation to their surrounding community composition and diurnal fluctuation in carbonate chemistry.

A review of the literature investigating the impact of OA on productivity (**Chapter 1**) examined the fact that increased pCO₂ only leads to a decline in productivity under relatively high light and/or temperature conditions (Reynaud et al. 2003; Langdon and Atkinson 2005; Anthony et al. 2008; Kaniewska et al. 2012; Anthony et al. 2013). This is confirmed by experiments in this thesis whereby CO₂-induced bleaching and productivity loss occurred in *Acropora aspera* under an average maximum daily irradiance of 860 μmol quanta m² s⁻¹ (**Chapter 2**), but not in *Acropora millepora*

exposed to an average maximum daily irradiance of 400 µmol quanta m² s⁻¹ (**Chapter 3**). In contradistinction to this analysis is the fact that juvenile *A. millepora* exhibited a bleaching response under very low light conditions (averaging 140 µmol quanta m² s⁻¹) (**Chapter 4**); here, the bleaching observed may have been partially due to the delayed uptake of *Symbiodinium* or may have been related to the difference in dominant *Symbiodinium* type. The divergence between adult and recruit responses may be associated with their differential skeletal morphology and light scattering processes. Recruits are small and less complex than adult colonies and this may partially explain their cryptic behaviour and settlement in low light areas. In contrast, adult have the ability to self-shade due to the complexity and orientation of their branches. Furthermore, adult colonies have a greater pool of energetic resources, which may result in increased capacity for photoprotection and photorepair.

Coral bleaching is a loss of symbionts or their pigments but these characteristics can also be features of the mechanism of photoacclimation (Brown 1997; Niyogi 1999). The differentiation lies in whether the bleaching represents a loss of health or fitness (Fitt et al. 2000). In *A. aspera*, the decrease in *Symbiodinium* cells per cm² led to a decrease in areal net maximum photosynthesis (Pnmax), which would have implications for the energy budget of the coral holobiont (Chapter 2). In the *A. millepora* juveniles, the photochemical efficiency of photosystem II (PSII) was not significantly affected by OA treatments despite the bleaching response (Chapter 4); future studies are needed employing oxygen microsensors to fully characterize the energy budgets of these juveniles. A decline in growth has been observed in *A. millepora* juveniles under OA conditions (Doropoulos et al. 2012), although this might be related to the increased energy cost of maintaining optimal conditions for skeletogenesis rather than decreased energy production, hence warranting the need for assays which would clearly partition photosynthetic and respiration changes in juveniles.

Photoacclimation optimises the light entering the photosystems in order to maximise the rate of photosynthesis and minimise photodamage (Niyogi 1999; Walters 2005). The potential decline of photorespiration as a photoprotective mechanism may contribute to CO₂-induced bleaching (Crawley et al. 2010) (**Chapter 2**). Indeed, this may explain the fact that CO₂-induced bleaching only occurs with concomitant high

light or temperature, as loss of photoprotection at sub-saturating light intensities would not be an issue for photosynthetic electron transport. Further molecular studies are needed to determine the expression and activation of Ribulose-1,5-bisphosphate carboxylase oxygenase (Rubisco) under increased pCO₂ conditions: the prediction from work in Chapter 2 would be that Rubisco levels may decline under increased pCO₂. In addition, the expression of enzymes downstream of phosphoglycolate phosphatase (PGPase) in the photorespiratory pathway, such as glycolate oxidase (GLO), should be assessed to confirm that the glycolate indeed enters that pathway. Furthermore, future experiments should compare the expression of these genes and proteins between *Symbiodinium* types in order to determine whether this led to the increased susceptibility of types D1 and D1-4.

Relaxation from carbon limitation may lead to dinoflagellate cell proliferation, as observed in *A. aspera* under the IPCC A2 CO₂ (pH 7.8 ± 0.1) conditions (**Chapter 2**). This would indicate that the host and symbiont carbon concentrating mechanisms (CCM) remain functional and that delivery of additional CO₂ occurs under OA conditions (Wooldridge 2013). Yet excess algal symbionts can increase the susceptibility of corals to thermal bleaching presumably due to subsequent CO₂ limitation under high light or temperature conditions (Cunning and Baker 2013). In the case of high pCO₂ conditions, the symbiont may photoacclimate to the new pCO₂ levels by reducing the use of the photorespiratory pathway, which would render the holobiont more susceptible to CO₂-induced bleaching with concomitant high light or temperature. This hypothesis requires further testing potentially through molecular analysis of carbonic anhydrase enzymes with an experiment designed to partition the variation caused by light, temperature and pCO₂.

Just as the susceptibility of corals to thermal bleaching depends on the physiological state of the coral prior to the thermal stress (Cunning and Baker 2013), the impact of OA may differ depending on the life history strategies of the coral and this aspect must be considered in future reef models. In **Chapter 3**, the *A. millepora* branches collected outside the lagoon at Lizard Island, characterised by a less dense population of symbionts, were the only coral branches to show a significant increase in net photosynthesis under OA conditions. Furthermore, comparison with coral branches

from a lagoon site showed a physiological trade-off between protein per surface area and skeletal density or branch compactness. Although there was no effect of OA on these characteristics in this short-term aquaria-based experiment, previous work has shown that increased protein and decreased calcification occur under long-term OA conditions (Fine and Tchernov 2007). A key finding in Chapter 3 is that coral branches may be less susceptible to the detrimental effects of OA in an environment that favours historical maintenance of a relatively smaller symbiont population. Moreover, allocation of resources to protein thickness at the expense of skeletal density or branch compactness may be an early indicator of the impact of OA, given that this life-history strategy is already presently utilized under the diurnal variation in carbonate chemistry. While recent studies have highlighted the need to incorporate diurnal variation in future experiments (Dufault et al. 2012; Kline et al. 2012; Shaw et al. 2013), Chapter 3 has highlighted the importance of life-long exposure to biogeochemical cycles at the coral collection sites, with implications for both experimental results and modeling. Future mechanistic and ecological models should be developed considering these physiological trade-offs under multiple anthropogenic stressors, such as temperature and pCO₂ in line with ecologically relevant diel oscillations.

In a rapidly changing environment, reproduction represents an important mode for recombination and adaption yet the uptake of *Symbiodinium* types appears to be nonspecific during early ontogeny (Abrego et al. 2009a; Cumbo et al. 2013). In addition, differential holobiont mortality can drive community shifts both at the symbiont and the host level (Sampayo et al. 2008). In some species, juveniles may revert to the homologous adult *Symbiodinium* type after > 3 years but other species, such as *A. millepora*, may lack specificity and therefore have greater opportunity for adaptation to new environmental conditions (Abrego et al. 2009b). The fitness and survival of the juvenile holobiont is key to this process. In **Chapter 4**, we observed a correlation between juvenile bleaching and the establishment of symbiosis with *Symbiodinium* types D1 or D1-4. Several questions arise from this finding. Firstly, does this level of bleaching indicate a change in fitness of the juveniles? If so, will this change the ratio of juvenile survival with certain *Symbiodinium* types? And lastly, what is the ecological importance of this mode of adaptation in comparison to differential adult survival, which may lead to community shifts as discussed in **Chapter 3**? Future OA

coral reproduction experiments should be carried out long term in order to determine the persistence of the shift in symbiont population and the implications for holobiont fitness. The hypotheses generated in **Chapters 3 & 4** suggest that OA should have profound effects on symbiont populations of any given coral species. However, recent research has shown that symbiont populations were similar regardless of distance from a volcanic CO₂ seep site (Noonan et al. 2013) and this warrants further investigation particularly in combination with the associated temperature increases predicted under future climate change.

In conclusion, this thesis has corroborated the OA-induced physiological changes of reef-building corals at various life history stages and examined the ecological implications of these findings. The increased susceptibility of corals to the detrimental impacts of increased pCO₂ is likely to coincide with additional environmental stressors such as high light and temperature, potentially due to the impact of OA on the photoprotective mechanism, photorespiration. In relation thereto, qualification of coral bleaching is required in order to differentiate bleaching detrimental to the health and fitness of the holobiont as opposed to photoacclimation. The research in this thesis has captured the link between bleaching and loss of productivity under high pCO₂ and high light conditions, particularly in adult colonies rather than juvenile recruits. The physiological changes induced by OA are likely to lead to processes of natural selection. Noteworthy in this context is the fact that diurnal variation in carbonate chemistry plays a significant role in shaping coral reef community composition. This suggests that anthropogenic OA may similarly place selective pressure on communities and these findings must be incorporated into future reef models. The rapid climate change predicted to occur over the coming century is likely to exhibit selective pressure and potentially cause differential mortality of both juvenile and adult reef-building corals. While this represents an adaptive mechanism, the resultant community shift may hinder the critical functions and services of coral reef ecosystems.

References

- Abrego D, van Oppen MJ, Willis BL (2009a) Highly infectious symbiont dominates initial uptake in coral juveniles. Molecular Ecology 18:3518-3531
- Abrego D, van Oppen MJH, Willis BL (2009b) Onset of algal endosymbiont specificity varies among closely related species of *Acropora* corals during early ontogeny. Molecular Ecology 18:3532–3543
- Anthony KRN, Kleypas JA, Gattuso J-P (2011) Coral reefs modify their seawater carbon chemistry implications for impacts of ocean acidification. Global Change Biology 17:3655-3666
- Anthony KRN, Kline DI, Diaz-Pulido G, Dove S, Hoegh-Guldberg O (2008) Ocean acidification causes bleaching and productivity loss in coral reef builders. Proceedings of the National Academy of Science 105:17442-17446
- Anthony KRN, Diaz-Pulido G, Verlinden N, Tilbrook B, Andersson AJ (2013)

 Benthic buffers and boosters of ocean acidification on coral reefs.

 Biogeosciences 10:4897-4909
- Baird AH, Cumbo VR, Leggat W, Rodriguez-Lanetty M (2007) Fidelity and flexibility in coral symbioses. Marine Ecology Progress Series 347:307-309
- Brown B (1997) Coral bleaching: causes and consequences. Coral Reefs 16:S129-S138
- Crawley A, Kline DI, Dunn S, Anthony K, Dove S (2010) The effect of ocean acidification on symbiont photorespiration and productivity in *Acropora formosa*. Global Change Biology 16:851-863
- Cumbo VR, Baird AH, van Oppen MJH (2013) The promiscuous larvae: flexibility in the establishment of symbiosis in corals. Coral Reefs 32:111-120
- Cunning R, Baker AC (2013) Excess algal symbionts increase the susceptibility of reef corals to bleaching. Nature Climate Change 3:259-262
- Doropoulos C, Ward S, Marshell A, Diaz-Pulido G, Mumby PJ (2012) Interactions among chronic and acute impacts on coral recruits: the importance of size-escape thresholds. Ecology 93:2131-2138
- Dufault AM, Cumbo VR, Fan T-Y, Edmunds PJ (2012) Effects of diurnally oscillating pCO₂ on the calcification and survival of coral recruits. Proc R Soc B-Biol Sci 279:2951-2958
- Fine M, Tchernov D (2007) Scleractinian coral species survive and recover from decalcification. Science 315:1811

- Fitt WK, McFarland FK, Warner ME, Chilcoat GC (2000) Seasonal patterns of tissue biomass and densities of symbiotic dinoflagellates in reef corals and relation to coral bleaching. Limnology and Oceanography 45:677-685
- Inoue S, Kayanne H, Yamamoto S, Kurihara H (2013) Spatial community shift from hard to soft corals in acidified water. Nature Climate Change DOI: 10.1038/NCLIMATE1855
- Jones AM, Berkelmans R, van Oppen MJH, Mieog JC, Sinclair W (2008) A community change in the algal endosymbionts of a scleractinian coral following a natural bleaching event: field evidence of acclimatization. Proceedings of the Royal Society of London: B 275:1359-1365
- Jones RJ, Hoegh-Guldberg O, Larkum AWD, Schreiber U (1998) Temperatureinduced bleaching of corals begins with impairment of the CO₂ fixation mechanism in zooxanthellae. Plant Cell and Environment 21:1219-1230
- Kaniewska P, Campbell PR, Kline DI, Rodriguez-Lanetty M, Miller DJ, Dove S, Hoegh-Guldberg O (2012) Major cellular and physiological impacts of ocean acidification on a reef building coral. Plos One 7:e34659
- Kleypas JA, Anthony KRN, Gattuso J-P (2011) Coral reefs modify their seawater carbon chemistry case study from a barrier reef (Moorea, French Polynesia). Global Change Biology 17:3667-3678
- Kline DI, Teneva L, Schneider K, Miard T, Chai A, Marker M, Headley K, Opdyke B, Nash M, Valetich M, Caves JK, Russell BD, Connell SD, Kirkwood BJ, Brewer P, Peltzer E, Silverman J, Caldeira K, Dunbar RB, Koseff JR, Monismith SG, Mitchell BG, Dove S, Hoegh-Guldberg O (2012) A short-term in situ CO₂ enrichment experiment on Heron Island (GBR). Scientific Reports 2:413
- Langdon C, Atkinson M (2005) Effect of elevated *p*CO₂ on photosynthesis and calcification of corals and interactions with seasonal change in temperature/irradiance and nutrient enrichment. Journal of Geophysical Research 110:1-16
- Niyogi K (1999) Photoprotection revisited: genetic and molecular approaches. Annual review of plant physiology and plant molecular biology 50:333-359
- Noonan SHC, Fabricius KE, Humphrey C (2013) *Symbiodinium* community composition in scleractinian corals Is not affected by life-long exposure to elevated carbon dioxide. PLoS One 8:e63985

- Reynaud S, Leclercq N, Romaine-Lioud S, Ferrier-Pages C, Jaubert J, Gattuso J (2003) Interacting effects of CO₂ partial pressure and temperature on photosynthesis and calcification in a scleractinian coral. Global Change Biology 9:1660-1668
- Sampayo EM, Ridgway T, Bongaerts P, Hoegh-Guldberg O (2008) Bleaching susceptibility and mortality of corals are determined by fine-scale differences in symbiont type. Proceedings of the National Academy of Sciences of the United States of America 105:10444-10449
- Shaw EC, McNeil BI, Tilbrook B, Matear R, Bates ML (2013) Anthropogenic changes to seawater buffer capacity combined with natural reef metabolism induce extreme future coral reef CO₂ conditions. Global Change Biology 19:1632-1641
- van Oppen MJH, Souter P, Howells EJ, Heyward A, Berkelmans R (2011) Novel genetic diversity through somatic mutations: Fuel for adaptation of reef corals? Diversity 3:405-423
- Walters RG (2005) Towards an understanding of photosynthetic acclimation. Journal of Experimental Botany 56:435-447
- Warner ME, LaJeunesse TC, Robison JD, Thur RM (2006) The ecological distribution and comparative photobiology of symbiotic dinoflagellates from reef corals in Belize: Potential implications for coral bleaching. Limnology & Oceanography 51:1887-1897
- Wooldridge SA (2013) Breakdown of the coral-algae symbiosis: towards formalising a linkage between warm-water bleaching thresholds and the growth rate of the intracellular zooxanthellae. Biogeosciences 10:1647-1658

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