

THE ROLE OF NATURAL ORGANIC LIGANDS IN
TRANSFORMATIONS OF IRON CHEMISTRY IN
SEAWATER AND THEIR EFFECT ON THE
BIOAVAILABILITY OF IRON TO MARINE
PHYTOPLANKTON.

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of Doctor of Philosophy in Science,
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CERTIFICATE OF AUTHORSHIP/ORIGINALITY

I certify that the work in this thesis has not been submitted for a degree nor has it been submitted as part of the requirements for a degree except as fully acknowledged within the text.

I also certify that the thesis has been written by me. Any help that I have received in my research work and the preparation of the thesis itself has been acknowledged. Furthermore, I certify that all information sources and literature used are indicated in the thesis.

Louisa Norman

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Figure 1.1 Schematic of the links between iron (Fe) and Carbon (C) cycling. Iron (Fe) enters the oceans via a number of sources, i.e. aerosol input (dust, ash), advective processes (horizontal transport of coastal water masses), upwelling of sediments. Fe is a vital micronutrient for phytoplankton as it is involved in the processes of photosynthesis and primary productivity. During photosynthesis phytoplankton fix atmospheric CO₂, thereby transforming inorganic carbon into organic forms which are transferred through the entire marine food web. Some of the organic carbon is respired by phytoplankton and bacteria, recycled through the food web, and exported to the sediments. During these processes Fe will be recycled and exported. Processes in bold black, iron inputs in blue, carbon processes in green, biological interactions in italics (From Norman et al., 2014).

Figure 1.2 The various size fractions, species, and associated biology and NOM of iron that exists in marine waters (From Norman et al., 2014).

Figure 1.3 Iron exists in the ocean mainly as Fe(III), either as inorganic Fe(III)', or bound to organic ligands (Fe(III)L). Organically bound Fe(III) is the predominant form (> 99%). Both Fe(III)' and Fe(III)L can be reduced by the action of sunlight (photoreduction, production of superoxide by NOM), or by biological activity (biological reduction, i.e. ferrireductase, and biological production of superoxide). Iron reduction can induce the dissociation of Fe(III)L (e.g dissociative reduction, DR), or generate Fe(II)L (e.g. non-dissociative reduction, NDR). The Fe(II)L complexes are weaker than Fe(III)L complexes and will easily dissociate to Fe(II)'. In oxygenated water the Fe(II)' is then rapidly reoxidised by O₂ to Fe(III)' (From Norman et al., 2014).

Figure 1.4 Schematic of the complex interplay between iron (Fe) chemistry and biology in defining its bioavailability to marine microorganisms. In surface water, Fe is mainly associated with particles (Partic), and with dissolved or colloidal organic ligands (L₂, e.g. exopolysaccharides, EPS; L₁ Sid, siderophores). Association with these compounds will define Fe chemical speciation and its reactivity towards the biota. Fe binding strength and reactivity is also affected by its redox chemistry (Red for reduction and Ox for oxidation), with Fe(II) usually forming the weakest complexes. Both biology (via surface reductase protein, ORProt) and light (λ) favour Fe reduction and subsequent transport with Fe(II) or Fe(III) transporters (FeTr) mainly present in eukaryotic phytoplankton. Highly specific transporter associated with siderophore uptake strategy, commonly present in

bacterioplankton, is represented separately (FeSidTr). Other non-specific uptake pathways (endocytosis, direct permeation and an ion channel) are shown. Once inside the microorganism, Fe (Fe_{int}) reacts with intracellular biological ligands (L_{bio} , e.g. Chlorophyll-*a*), is stored (e.g. vacuole, ferritin) or is involved in cellular homeostasis via gene regulation (grey arrow with \pm symbol). Release of Fe biological organic ligands (L_{rel} , such as EPS and siderophores) can exert a feedback in the control of both Fe chemistry and bioavailability. Dotted, dashed and full arrows represent aggregation/disaggregation, transfer, and chemical reaction (complexation, redox), respectively. (From Hassler et al, 2011b).

Fig. 2.1. Sea surface temperature (SST) and Chlorophyll-*a* ($\mu\text{g L}^{-1}$) plots showing the study area and sampling locations for natural humic substance-like material and nutrient enrichment experiments. Natural samples were collected from a variety of water mass types (river plume, inner shelf, outer shelf, and oceanic (cold-core cyclonic eddy (CCE) and East Australia Current (EAC)), and seawater collected for the nutrient experiments was sampled from the EAC and CCE.

Fig. 2.2 Calibration curve used for the comparison of methods to determine the concentration of electrochemically detected humic substance-like (HS-like) material. Suwannee River Fulvic Acid (SRFA) was used as the HS-like standard in concentrations between 20 and 480 $\mu\text{g L}^{-1}$. i_p represents the peak height in nA of electrochemically detected Fe'-reactive organic material. Errors = SD of triplicate samples.

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Fig. 2.6 Concentration of dissolved Fe (dFe, nM) measured at the conclusion of two 72-h shipboard nutrient-enrichment experiments undertaken during the SS2010-V09 Tasman Sea voyage (*RV Southern Surveyor*, 15th to 31st October 2010, austral spring). The experiments were conducted in 200–210- μm filtered seawater collected from the depth of the chlorophyll maximum at two sites in the Tasman Sea; East Australia Current (EAC, 29 1 °S 154 3°E), and a cold-core eddy (CCE, 32 2°S 153 8°E). Treatments were as per Fig. 2.5. Samples for the analysis of dFe were taken from replicates 1 and 2 of each treatment, therefore duplicate data points are shown for each treatment and experiment.

Fig. 2.7 Cell abundance (cells mL^{-1}) of bacteria (A), and picophytoplankton *Prochlorococcus* (B), *Synechococcus* (C), small eukaryotes (D) and large eukaryotes (E) measured by flow cytometry at T0 and at the conclusion of two 72-h shipboard nutrient-enrichment experiments undertaken during the SS2010-V09 Tasman Sea voyage (*RV Southern Surveyor*, 15th to 31st October 2010, austral spring). The experiments were conducted in 200–210- μm seawater collected from the depth of the chlorophyll maximum at two sites East Australia Current (EAC, 29 1 °S 154 3°E), and a cold-core eddy (CCE, 32 2°S 153 8°E). T0 = unamended seawater at the start of the experiment. Treatments were as per Fig. 2.5. Error = SD of

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Fig. 2.11 Concentration of electrochemically detected humic substance-like (HS-like) substances measured at the conclusion of two 72-h shipboard nutrient-enrichment experiments undertaken during the SS2010-V09 Tasman Sea voyage (*RV Southern Surveyor*, 15th to 31st October 2010, austral spring). The experiments were conducted in 200–210- μ m filtered seawater collected from the depth of the chlorophyll maximum at two sites, East Australia Current (EAC; 29 1 °S 154 3°E), and a cold-core eddy (CCE; 32 2°S 153 8°E). T0 = unamended seawater at the start of the experiment. Treatments were as per Fig. 2.7. The concentration of HS-like material is expressed as Suwannee River Fulvic Acid equivalents

(SRFA eq) in $\mu\text{g L}^{-1}$. Error = SD of triplicate incubations except for EAC FAD where errors represent half interval (range) of duplicates incubations.

* Significantly higher HS-like concentration compared to all other EAC treatments, except FAL and FAD ($p = 0.003$).

† Significantly higher HS-like concentration compared to CCE T0 and control ($p = 0.007$).

‡ Significantly lower HS-like concentration compared to CCE T0 and control ($p = 0.014$).

Fig. 2.12 Relationships between the concentration of humic substance-like (HS-like) material and Silicic acid (Si(OH)_4), phosphate (PO_4) and dissolved Fe (dFe) at the conclusion of a 72-h shipboard nutrient-experiment undertaken during the SS2010-V09 Tasman Sea voyage (*RV Southern Surveyor*, 15th to 31st October 2010, austral spring). The experiment was conducted in 200–210- μm filtered seawater collected from the depth of the chlorophyll maximum in the East Australia Current (EAC; 29 1 °S 154 3°E). Treatments were as per Fig. 2.5. Panel A = Si(OH)_4 all data; Panel B = treatments where $\text{Si(OH)}_4 < 0.7 \mu\text{mol L}^{-1}$; Panel C = treatments where $\text{Si(OH)}_4 > 20 \mu\text{mol L}^{-1}$; Panel D = PO_4 all data; Panel E = PO_4 enrichment treatment (Mix) excluded; Panel F = dFe all data; Panel G = dFe-enrichment $> 10 \text{ nM}$ (Mix treatment) excluded. High concentrations, subsequently excluded, are circled to highlight (panels A, D, and F).

Fig. 3.1 Cruise track from the Primary productivity induced by Iron and Nitrogen in the Tasman Sea (PINTS) voyage (*RV Southern Surveyor*, Jan-Feb 2010). Transect stations are shown as circles and process stations as diamonds. Profiles presented in this chapter were from two process stations P1 (30.0 °S, 156.0 °E, also Stn 5) and P3 (46.2 °S, 159.5 °E, also Stn 12) and from Stn 14, 44.6 °S, 149.4 °E. Stn 14 was a reoccupation of process station 3 from the SAZ-Sense expedition (*Aurora Australis*, January–February 2007). Water for the Fe-enrichment experiments was collected stations P1 and P3. Thicker solid lines indicate the East Australian Current (EAC), Tasman Front (TF), and EAC Extension. The dashed line represents the path of the subtropical front (STF) (From Hassler et al., 2014).

Fig. 3.2 Seawater depth profiles of dissolved nutrients nitrate + nitrite (NO_x ; panel A), reactive phosphorus (PO_4 ; panel B), and silicic acid (Si(OH)_4 ; panel C), measured at stations P1 (30.0 °S, 156.0 °E), P3 (46.2 °S, 159.5 °E) and Stn 14 (44.6 °S, 149.4 °E) collected during the PINTS voyage (*RV Southern Surveyor*, Jan-Feb 2010).

Fig. 3.3 Seawater depth profiles of total chlorophyll-*a* (TChl-*a* $\mu\text{g L}^{-1}$) measured at stations P1 (30.0 °S, 156.0 °E, depths 15 to 125 m), P3 (46.2 °S, 159.5 °E, depths 15 to 80 m) and Stn 14 (44.6 °S, 149.4 °E, depths 15 to 50 m) collected during the PINTS voyage (RV *Southern Surveyor*, Jan-Feb 2010).

Fig. 3.4 Seawater depth profiles (15 to 1000m) from process stations P1 (30.0 °S, 156.0 °E), and P3 (46.2 °S, 159.5 °E) and Stn 14 (44.6 °S, 149.4 °E) collected during the PINTS voyage (RV *Southern Surveyor*, Jan-Feb 2010). Concentration of dissolved Fe (dFe, nM) and the concentration of electrochemically detected Fe'-binding organic ligands (SumL, nM) and their calculated conditional stability constant (Log $K_{\text{Fe'L}}$) are presented together with the concentration of humic substance-like (HS-like) material. HS-like material is expressed as Suwannee River Fulvic Acid (SRFA) $\mu\text{g L}^{-1}$.

Fig. 3.5 Relationship between the concentration of Fe-binding organic ligands (ΣL) and the conditional stability constant (Log $K_{\text{Fe'L}}$) for process station P1 (30.0 °S, 156.0 °E), process station P3 (46.2 °S, 159.5 °E) and Stn 14 (44.6 °S, 149.4 °E). Samples were collected during the PINTS voyage (RV *Southern Surveyor*, Jan-Feb 2010; Hassler et al., 2014).

Fig. 3.6 Relationships between the concentration of total chlorophyll-*a* (TChl-*a*) and Fe-binding organic ligands (ΣL), TChl-*a* and ligand conditional stability constant (log $K_{\text{Fe'L}}$) at depths between 15 and 125 m at process station P1 (30.0 °S, 156.0 °E), and TChl-*a* and ΣL at depths between 15 and 50 m, humic substance-like (HS-like) material and SumL at depths between 15 and 300 m and at Stn 14 (44.6 °S, 149.4 °E). Samples were collected during the PINTS voyage (RV *Southern Surveyor*, Jan-Feb 2010; Hassler et al., 2014). HS-like material is expressed as Suwannee River Fulvic Acid (SRFA) equivalent in $\mu\text{g L}^{-1}$.

Fig. 3.7 Relationships between dissolved Fe (dFe) concentration (nM) and macronutrients nitrate + nitrite (NO_x), phosphate (PO_4) and silicic acid ($\text{Si}(\text{OH})_4$) ($\mu\text{mol L}^{-1}$) at process station P3 (46.2 °S, 159.5 °E) at depths between 15 and 300 m. Samples were collected during the PINTS voyage (RV *Southern Surveyor*, Jan-Feb 2010; Hassler et al., 2014).

Figure 3.8 Concentrations of dissolved Fe (dFe, nM) and relative concentration (%) of labile Fe ($\text{Fe}_{\text{Labile}}$) associated with Fe enrichment experiments using phytoplankton communities collected from two sites in the Tasman Sea, P1 (30.0 °S, 156.0 °E, panels A and C) and P3 (46.2 °S, 159.5 °E, panels B and D) during the PINTS voyage (RV *Southern Surveyor*, Jan-Feb 2010). The data presented comes from unamended seawater (T0) and after 4-d incubation in samples with and without the addition of Fe and organic ligands. Treatments measured after 4-d incubation comprised an unamended control (Con), inorganic Fe only (2 nM, Fe),

desferrioxamine B ([15 nM], DFB), glucuronic acid ([15 nM], GLU), natural pelagic bacterial exopolymeric substances ([0.8 nM], EPS), fulvic acid ([100 $\mu\text{g L}^{-1}$], as Suwannee River Fulvic Acid, FA), and two treatments containing Australian desert dust (D1, 2009 Brisbane dust storm, and D2, red composite, both from the Buronga region, NSW) which were predicted to release ~ 2 nM Fe. DFB, EPS, GLU and FA treatments were all enriched with 2 nM inorganic Fe. Closed symbols indicate samples with phytoplankton present, open symbols indicate samples where phytoplankton were absent (0.2 μm filtered, single incubations). Error bars represent half-interval of duplicate samples; where no error bars are present the data presented is from a single sample.

Figure 3.9 Concentration of organic ligands and calculated conditional stability constants ($\log K_{\text{Fe}^{\text{L}}}$) associated with Fe-enrichment experiments using phytoplankton communities collected from two sites in the Tasman Sea, P1 (30.0 °S, 156.0 °E, panels A and C) and P3 (46.2 °S, 159.5 °E, panels B and D) during the PINTS voyage (RV *Southern Surveyor*, Jan-Feb 2010). The data presented comes from unamended seawater (T0) and after 4-d incubation for samples with and without the addition of Fe and organic ligands. Treatments were as per Fig. 3.8. Closed symbols indicate samples with phytoplankton present, open symbols indicate samples where phytoplankton were absent (0.2- μm filtered, single incubations). Where two ligand classes were detected, stronger ligands are indicated by a red symbol and weaker ligands by a blue. Error bars represent half-interval of duplicate samples; where no error bars are present the data presented is from a single sample.

Figure 3.10 Concentration of humic substance-like material (HS-like), expressed as Suwannee River Fulvic Acid equivalents (SRFA eq) in $\mu\text{g L}^{-1}$, associated with Fe enrichment experiments using phytoplankton communities collected from two sites in the Tasman Sea, P1 (30.0 °S, 156.0 °E, panel A) and P3 (46.2 °S, 159.5 °E, panels B) during the PINTS voyage (RV *Southern Surveyor*, Jan-Feb 2010). The data presented comes from unamended seawater (T0) and after 4-d incubation for samples with and without the addition of Fe and organic ligands. Treatments were as per Fig. 3.8. Closed symbols indicate samples with phytoplankton present, open symbols indicate samples where phytoplankton were absent (0.2- μm filtered, single incubations). Error bars represent half-interval of duplicate samples; where no error bars are present the data presented is from a single sample. Note difference in y-axis scale.

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and P3 (46.2 °S, 159.5 °E) during the PINTS voyage (RV *Southern Surveyor*, Jan.-Feb. 2010). Treatments were as per Fig. 3.8 Error bars represent the half interval of duplicate samples. T0 values not shown; see Table 3.1.

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SUMMARY

It is widely accepted that the complexation of iron (Fe) with organic compounds is the primary factor that regulates Fe reactivity and its bioavailability to phytoplankton in the open ocean. Despite considerable efforts to unravel the provenance of the many organic ligands present in the 'ligand soup' much of this pool remains largely unresolved and the ligands remain grouped into either strong (L_1) or weak (L_2) types. The Tasman Sea and Southern Ocean are areas of particular interest as both regions are subject to Fe limitation or co-limitation and are likely to be severely affected under climate change scenarios. The predictions of dryer conditions in central Australia suggest that the Tasman Sea may be subject to changes in the intensity and frequency of atmospheric dust deposition and, in consequence, enhanced Fe deposition into the surface waters. This thesis aims to improve our knowledge of a) how natural organic ligands affect Fe solubility, chemistry, and bioavailability, and b) which forms of Fe are available to phytoplankton.

Natural seawater samples (surface and profiles to 1000m) revealed that electrochemically detected HS-like material, which are thought to make up a proportion of the weaker L_2 class of ligands, account for a very small fraction of the Fe-binding organic ligand pool. The distribution of HS-like material in coastal, shelf and offshore regions associated with the EAC does not exhibit a nearshore to offshore (high to low) concentration gradient, likely because of low riverine HS-like input. Higher concentrations of HS-like material were generally found at, or adjacent to, the chlorophyll maximum (C_{max}). However, little correlation with chlorophyll-*a* (Chl-*a*) was observed and so these higher concentrations are more likely linked to degraded algal material and microbial activity rather than direct primary productivity. Perturbation experiments using water collected offshore in the EAC and a cold core cyclonic eddy (CCE) indicated that the *in situ* utilisation and production of HS-like material, and its character, differ depending on the phytoplankton and microbial communities present, and reflect the biological activities of these different communities, as well as photochemical transformations. The addition of a model HS (Suwannee River fulvic acid) enhanced Chl-*a* concentration in both communities, particularly in the EAC, likely due to the remineralisation of Fe and other nutrients via photochemical and bacterial transformation of this material.

Seawater depth profiles from the northern and southern Tasman Sea indicate Fe limitation (or co-limitation) at the stations sampled. Dissolved Fe (dFe), organic ligand concentrations and conditional stability constants were consistent with previous studies (showing the

presence of mostly L₂ ligands) with higher ligand concentrations and conditional stability constants close to the C_{max}. Ligand concentration, as previously reported, is in excess of dFe throughout the water column, although no correlation between dFe and ligand concentration was observed.

Fe-enrichment experiments using two contrasting phytoplankton communities investigated how the communities respond, in terms of biomass and community structure, to inorganic Fe delivered alone or bound to an organic ligand (siderophore, saccharides, bacterial exopolymeric substances (EPS)) or dust-borne Fe from two dust samples (D1 and D2) originating from the Australian continent. Overall, Fe bound to a strong Fe-binding siderophore was much less available to both phytoplankton communities; whereas, Fe bound to bacterial EPS (lowest conditional stability constant) induced the greatest increase in overall phytoplankton biomass. Dust D1 did not have the highest rate of dFe uptake, or result in the greatest increase Chl-*a*, but did induce the greatest shift in community structure. Whilst one ligand (L₂) was measured in most incubations, both L₁ and L₂ ligands were detected in the D1 and inorganic Fe incubations, indicating *in situ* biological production of Fe-binding ligands (i.e. siderophores or EPS) in response to Fe addition and an added ligand component from the dust. The greater response of the phytoplankton to the EPS and D1 led to further laboratory experiments.

Analysis of 4 EPS isolates (1 bacterial, 1 mixed natural community, and 2 microalgal laboratory cultures) showed that both bacterial and algal EPS contain functional components known to bind Fe (uronic acid, saccharides). The bacterial EPS was made up of mainly high molecular mass components, whereas the algal EPS were of low molecular mass. Most EPS contained components that were measured as both L₁ and L₂ ligands, with the L₁ ligands having an affinity for Fe close to that of bacterial siderophores. EPS greatly enhanced Fe solubility in seawater, however, it may also accelerate Fe(II) oxidation, and thus, Fe(II) removal from the system. Other trace elements and macronutrients were associated with the EPS that may be accessible to phytoplankton and could help to relieve nutrient limitation. Bioaccumulation experiments indicated that Fe bound to all EPS used was highly bioavailable to the Southern Ocean diatom *C. simplex* (50 to > 100%) relative to the bioavailability of inorganic Fe (assumed 100% bioavailable). This enhanced bioavailability was likely due to increased Fe solubility, and possible formation of more bioavailable forms of Fe.

Further experiments using dust D1, and rainwater collected in the Tasman Sea, revealed that despite low fractional solubilities (< 1%), the dust represents, potentially, an important

source of Fe and other vital macronutrients and trace elements. Both the rainwater and dust were associated with ligands in the L₂ class that helped to maintain the solubility of Fe. Light exposure, particularly UV, can a) have a substantial effect on the Fe chemistry of the Fe-laden dust, lowering the conditional stability constant and altering the size distribution of both Fe and ligands (including saccharides and HS-like material), and b) improve the bioavailability of dust-borne Fe to *C. simplex*.

The perturbation experiments in the EAC, CCE and north and south Tasman Sea demonstrated that organic ligands play an important role in regulating the nutrient dynamics of marine systems. They show that the bioavailability of Fe to phytoplankton is dependent on the various Fe species and Fe sources (i.e. inorganic Fe, organically bound, dust-borne), and that this differs between phytoplankton size fractions and from one bacterio- or phytoplankton species to another. The Tasman Sea and Southern Ocean receive, possibly increasing, periodic inputs of atmospheric dust from the source region of D1, which initiated a substantial community shift in perturbation experiments. However, the impact that dust-borne Fe will have on a natural phytoplankton community will be dependent on the duration and intensity of the dust deposition event, and the nutritive state and community structure of the resident phytoplankton. Bacterial siderophores have previously been suggested as key players in Fe biogeochemistry, however, in remote regions bacterial and algal EPS could play a significant role in the biogeochemical cycling of Fe and other nutrients, and their contribution should also be considered to further our understanding of the dynamics of Fe-limited oceans.